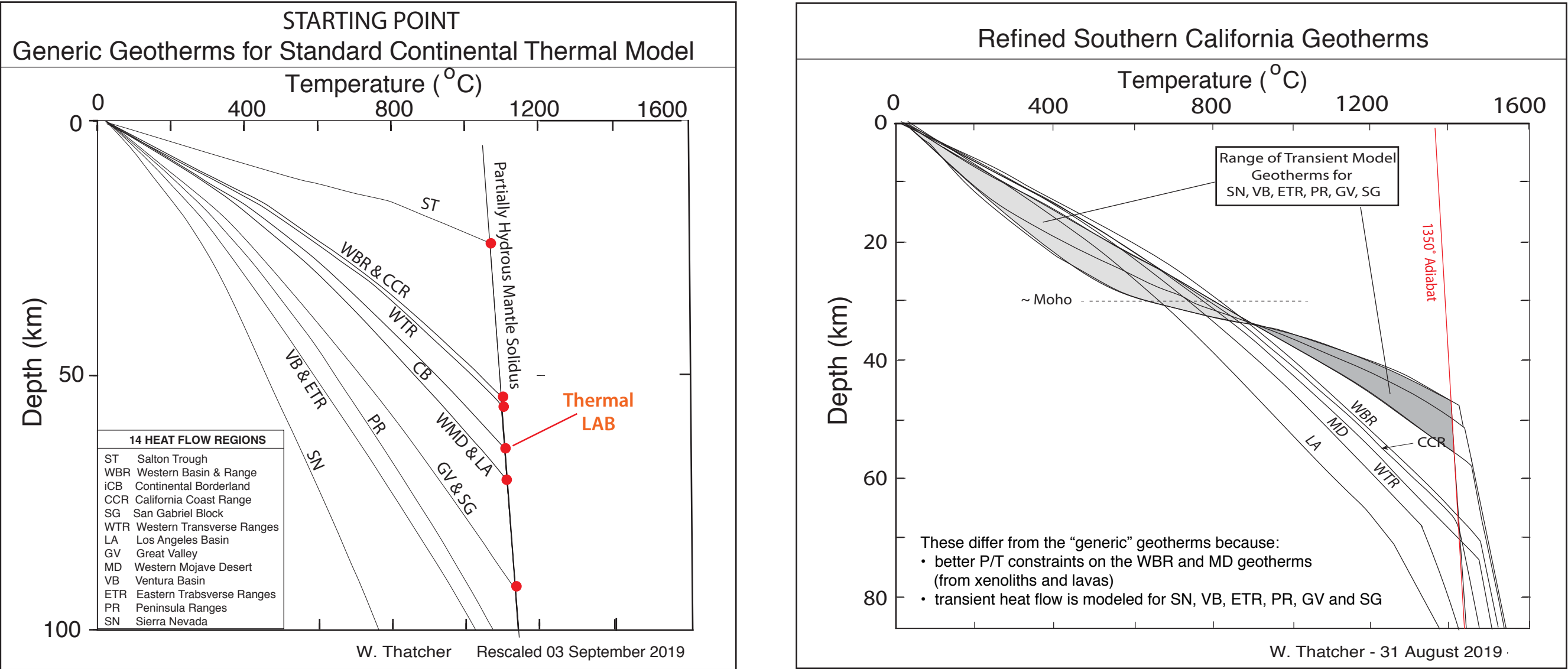


The SCEC Community Rheology Model (CRM) is a three-dimensional description of the rheology of southern California's lithosphere, based on an ongoing synthesis of data from a wide range of sources. These sources include but are not limited to seismic imaging studies, rock deformation experiments and theory, regional-scale geological mapping, detailed descriptions of rocks (petrology and fabric), and thermal modeling constrained by surface heat flow data and depth to the seismic LAB. During the SCEC5 period we have been assembling a preliminary version of the CRM that makes use of a simplified representation of the regional geologic and thermal structure. This *preliminary* CRM comprises a thermal model (CTM), a geologic framework model (GFM), and viscous rheologies for each of the GFM rock types. The preliminary CTM and GFM are depth profiles of temperature and lithology, respectively, assigned to geographic subregions. Viscous flow laws for each GFM rock type have been developed, based on mixing relationships and rheological information for the component minerals. We are in the process of integrating these components for distribution via the SCEC UVCN framework and other formats.

Community Thermal Model (CTM)

Refined Community Thermal Model (CTM) geotherms are very different and much improved from generic starting model

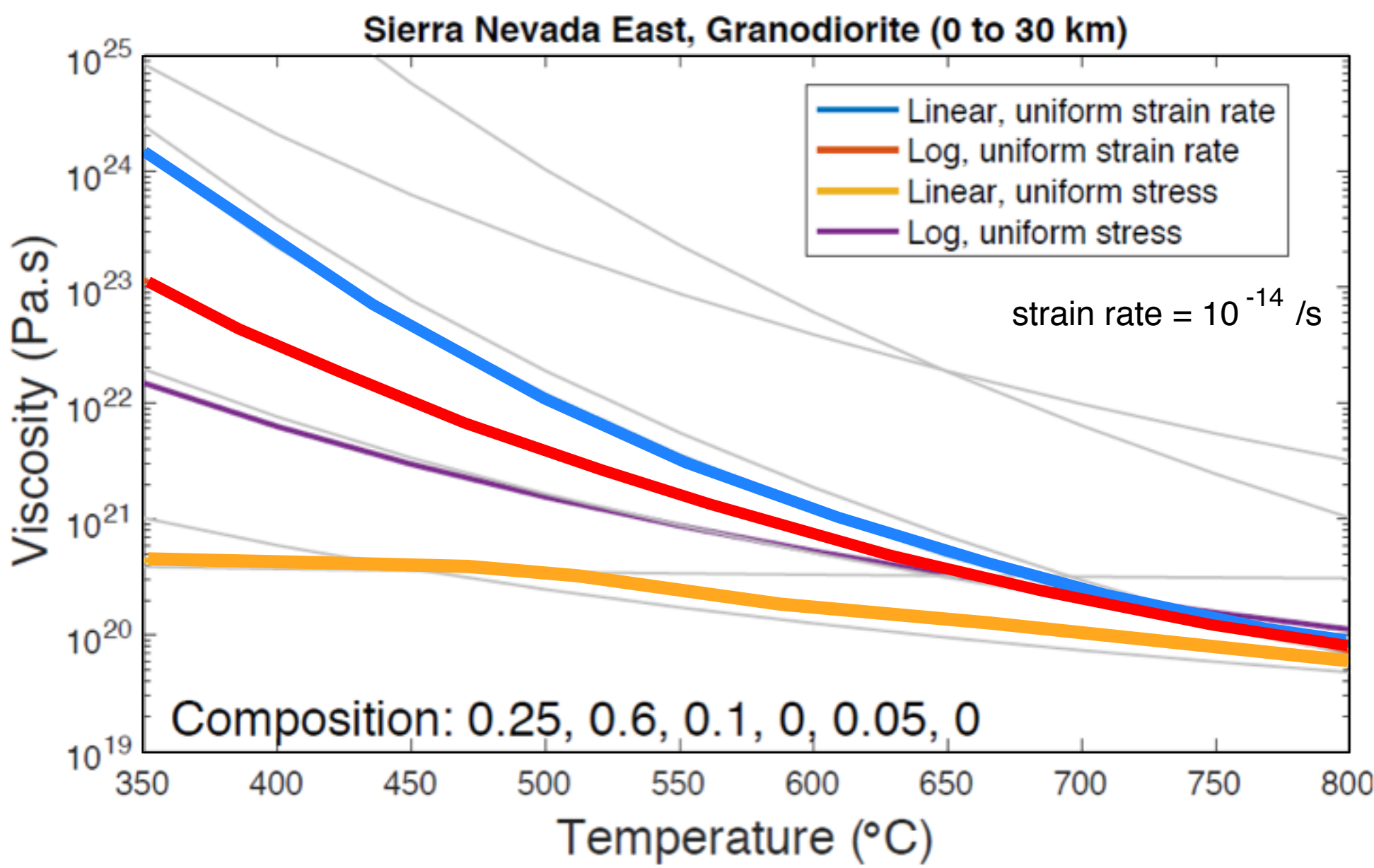


- Narrower temperature range than generic geotherms
- Almost all SoCal has warm crust and upper mantle lid
- LAB temperature is warmer too: 1200° - 1400° C

Important implications for CRM: Less lateral rheological variation

Rock Rheologies

- Mineral flow laws are specified for feldspar, quartz, pyroxene, olivine, biotite and amphibole (table below)
- GF “rocks” are defined by % of these minerals, per GFM lithology descriptions
- Aggregate viscous rheology is calculated from mixing relations, assuming dislocation creep (box below)
- Viscosity and rock type estimates based on the SCEC CVM and thermal models are also being incorporated into the gridded GFM (Shinevar et al., 2018)



Viscosity for granodiorite/ quartz diorite, calculated with four different mixing assumptions

- Quartz 25%
- Feldspar 60%
- Amphibole 5%
- Mica 10%

Effective viscosity for granodiorite exceeds 10²⁰ Pa s at temperatures typical for the lower crust (450-780°C at 25-30 km depth, per CTM). Mixing law choice matters, though less at high temperatures.

For non-shear zone rocks, we recommend the Huet et al. (2014) mixing law and assume uniform strain rate (result is similar to average of red and blue lines on figure above). For shear zone rocks, the linear, uniform stress relationship (orange line on figure above) is recommended. Note that shear zone strain rate will be about 100 times higher than assumed above, resulting in much lower effective viscosities than indicated by the orange line.

Consensus major mineral flow laws for the preliminary CRM

Mineral	Quartz	Feldspar	Biotite	Pyroxene	Amphibole	Olivine
RHEOL code	HTD01	R006wd	Kr90	HC90amp	HC90amp	HC90amp
Reference	Hirth, Thatcher and Dunlap, 2001, quartzite creep	Rybacki and Dresen 2004, at 1, 1990, wet An100 in biotite dislocation creep	Kronenberg et al., 1990, wet dislocation creep	Dimanov and Hacker, 2005, dislocation creep	Hirth and Kohlstedt, 2003, wet dislocation creep	Hirth and Kohlstedt, 2003, wet dislocation creep
Category	2 (function of T, P and fw)	6 (function of T, P and fw)	1 (function of T)	1 (function of T)	1 (function of T)	6 (function of T, P and fw)
Strain rate	$\dot{\epsilon} = \frac{A}{B} \exp\left(\frac{-Q}{RT}\right) \exp\left(\frac{-P}{K}\right) \exp\left(\frac{-V}{RT}\right)$	$\dot{\epsilon} = \frac{A}{B} \exp\left(\frac{-Q}{RT}\right) \exp\left(\frac{-P}{K}\right) \exp\left(\frac{-V}{RT}\right)$	$\dot{\epsilon} = \frac{A}{B} \exp\left(\frac{-Q}{RT}\right) \exp\left(\frac{-P}{K}\right) \exp\left(\frac{-V}{RT}\right)$	$\dot{\epsilon} = \frac{A}{B} \exp\left(\frac{-Q}{RT}\right) \exp\left(\frac{-P}{K}\right) \exp\left(\frac{-V}{RT}\right)$	$\dot{\epsilon} = \frac{A}{B} \exp\left(\frac{-Q}{RT}\right) \exp\left(\frac{-P}{K}\right) \exp\left(\frac{-V}{RT}\right)$	$\dot{\epsilon} = \frac{A}{B} \exp\left(\frac{-Q}{RT}\right) \exp\left(\frac{-P}{K}\right) \exp\left(\frac{-V}{RT}\right)$
Stress	$\sigma = \frac{A}{B} \exp\left(\frac{-Q}{RT}\right) \exp\left(\frac{-P}{K}\right) \exp\left(\frac{-V}{RT}\right)$	$\sigma = \frac{A}{B} \exp\left(\frac{-Q}{RT}\right) \exp\left(\frac{-P}{K}\right) \exp\left(\frac{-V}{RT}\right)$	$\sigma = \frac{A}{B} \exp\left(\frac{-Q}{RT}\right) \exp\left(\frac{-P}{K}\right) \exp\left(\frac{-V}{RT}\right)$	$\sigma = \frac{A}{B} \exp\left(\frac{-Q}{RT}\right) \exp\left(\frac{-P}{K}\right) \exp\left(\frac{-V}{RT}\right)$	$\sigma = \frac{A}{B} \exp\left(\frac{-Q}{RT}\right) \exp\left(\frac{-P}{K}\right) \exp\left(\frac{-V}{RT}\right)$	$\sigma = \frac{A}{B} \exp\left(\frac{-Q}{RT}\right) \exp\left(\frac{-P}{K}\right) \exp\left(\frac{-V}{RT}\right)$
n	4	3	18	5.5	3.7	3.5
Q	135000	345000	51000	54000	244000	520000
V	0	38e-6	0	0	0	22e-6
B	1.1941e+10	5.1951e-07	2.7013e-07	4.2398e-05	7.0505e-06	8.3362e+06
A	6.3096e-42	1.5849e-24	1.0000e-138	6.9405e-33	6.9405e-27	1.0095e-25
m	0	0	0	0	0	0
p	1	1	0	0	0	1.2

Example: flow law for non-shear zone granodiorite. Use Huet et al. (2014) equation for aggregate viscosity of rock with i mineral phases in proportions ϕ_i :

$$\eta_{\text{aggregate}} = \sum_i \frac{\phi_i n_i}{n_i + 1} \prod_j \left(\eta_j \frac{n_j + 1}{n_j} \right)^{\frac{\phi_j \phi_i n_j}{\phi_i \phi_j n_j + 1}}$$

After much math, this simplifies to a power law:

$$\eta_{\text{Huet}} = 1.3934 \times 10^8 \exp\left(\frac{9376.1}{T}\right) \epsilon^{-0.7246} \phi_w^{0.2551}$$

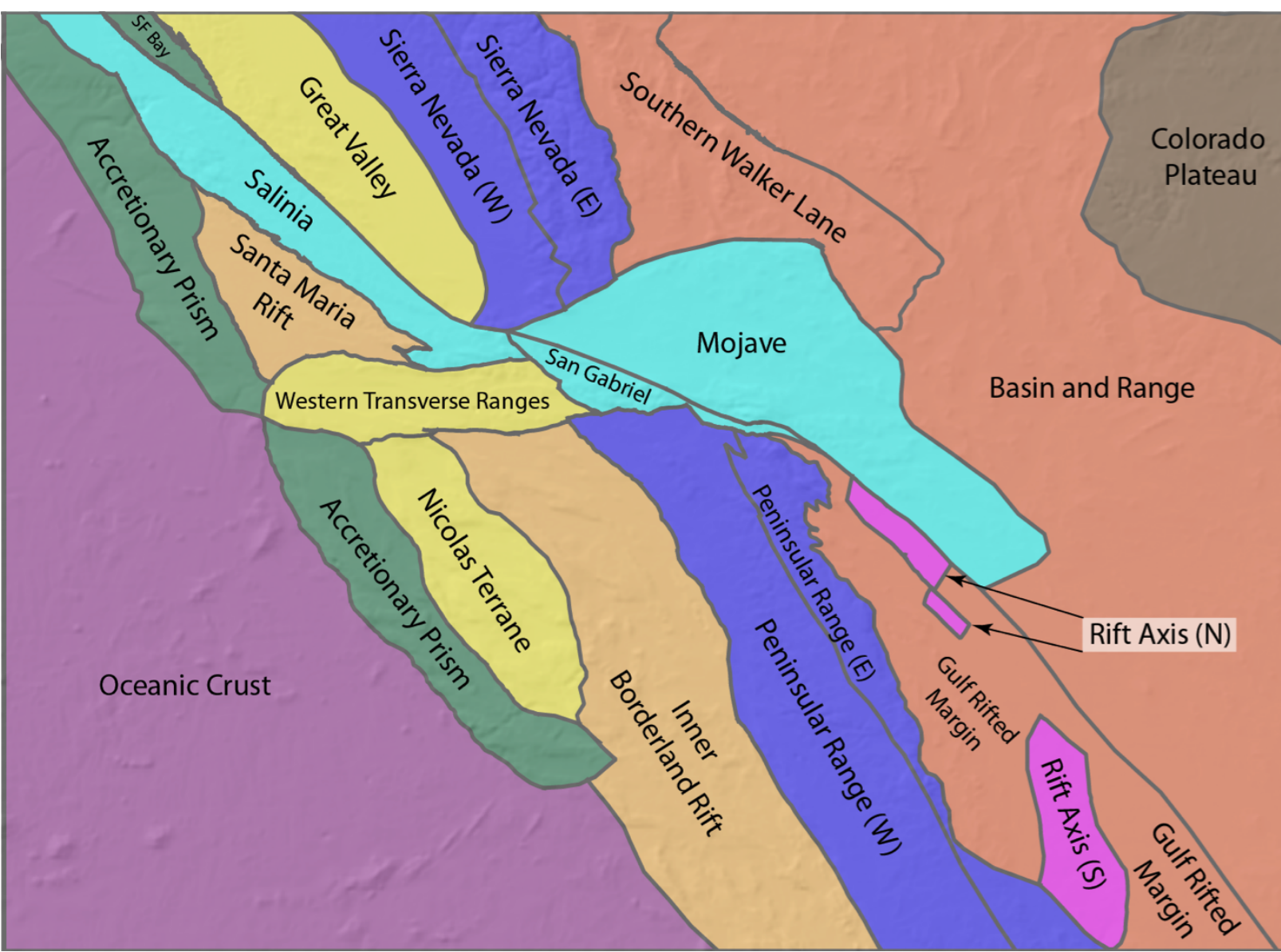
The CRM will include aggregate flow laws and guidance for each GF rock type, for low- and high-strain rocks. Some may require numerical solutions. Equations and parameter guidance for water fugacity (fw) will also be provided.

References:
Huet, B., P. Yamato, and B. Grasemann (2014). The Minimized Power Geometric model: An analytical mixing model for calculating polyphase rock viscosities consistent with experimental data. J. Geophys. Res. Solid Earth, 119, 3897–3924. doi:10.1002/2013JB010453.
Shinevar, W. J., Behn, M. D., Hirth, G. and Jagodzki, O. (2018). Inferring crustal viscosity from seismic velocity: Application to the lower crust of Southern California. EPSL, 494, 83–91.

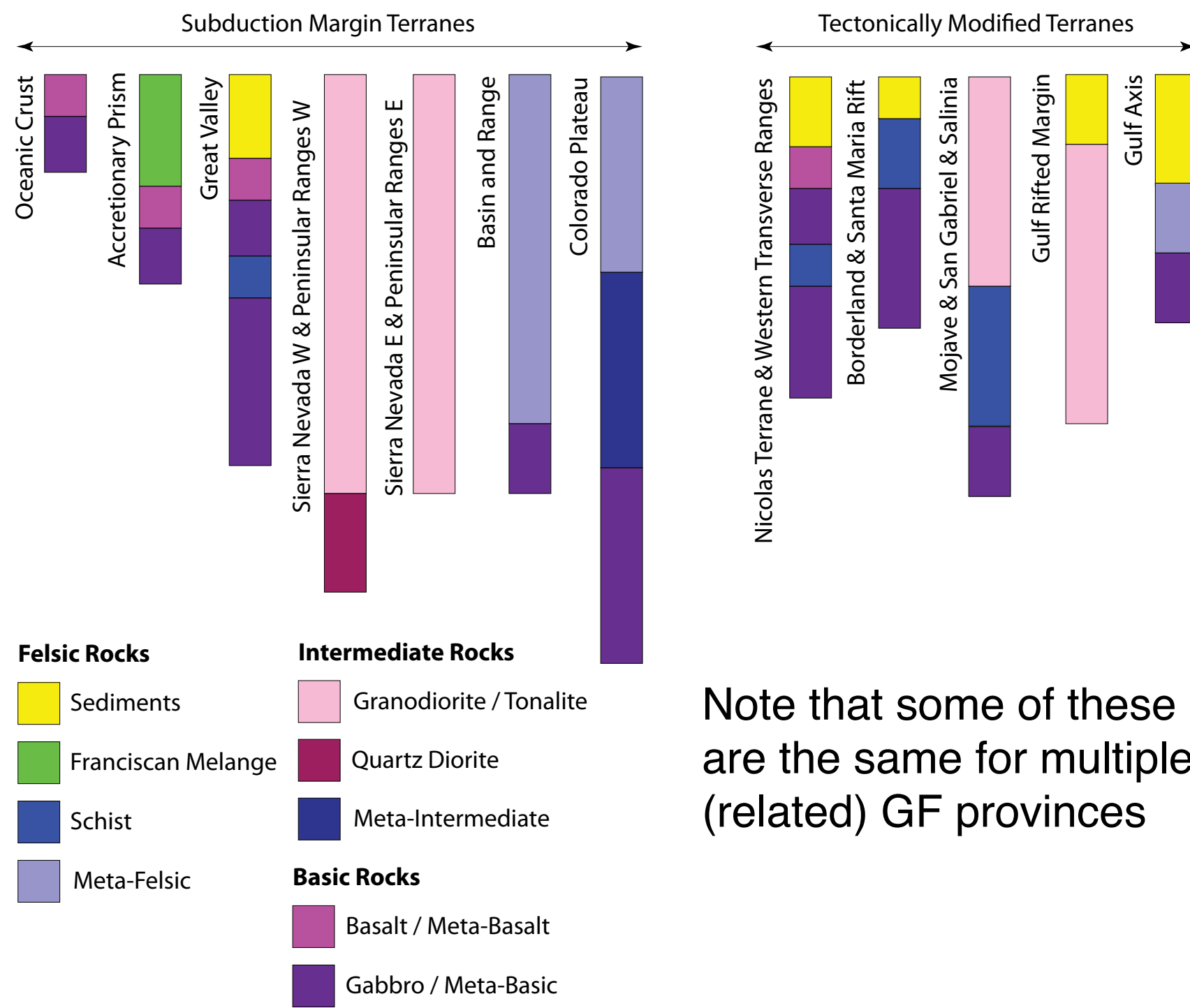
Geologic Framework Model (GFM) and query tool

Preliminary Geologic Framework Model (2018)

Geologic Framework Province Boundaries and Names



Lithology versus Depth: Preliminary GFM

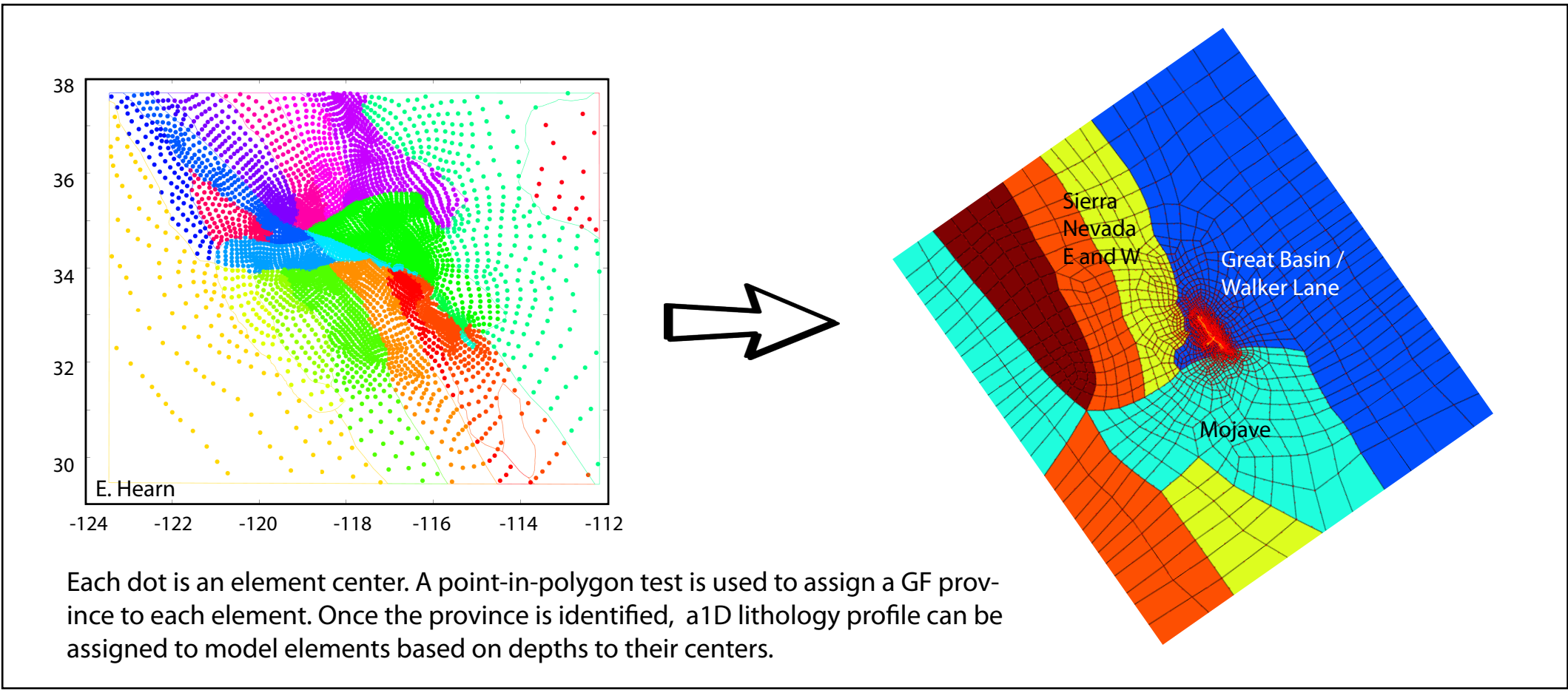


Note that some of these are the same for multiple (related) GF provinces

Examples of GF lithology descriptions for rocks in each province
Proportions of minerals are required to calculate whole-rock flow laws

CRM Crustal Columns										
Slab Format										
	DEPTH			Mineral %						
Domain	Start	End	Rock Type	Quartz	Feldspar	Mica	Pyroxene	Amphibole	Olivine	TOTAL
Sierra Nevada East	0	30	Granodiorite	25	60	10	0	5	0	100
Sierra Nevada East	30	35	Quartz Diorite	15	60	10	0	15	0	100
Sierra Nevada West	0	30	Granodiorite	20	60	10	0	10	0	100
Sierra Nevada West	30	40	Quartz Diorite	15	60	10	0	15	0	100

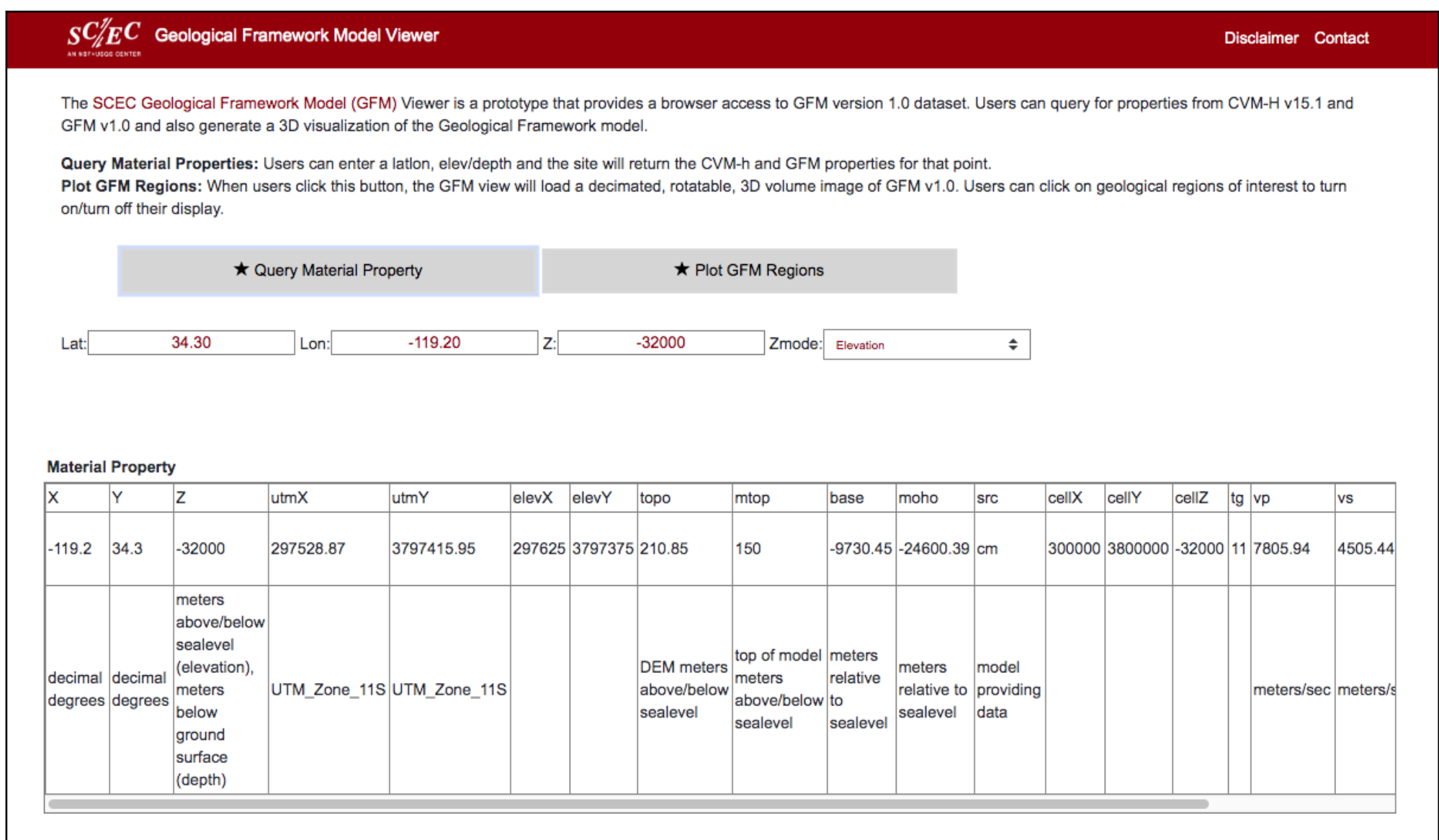
Gridded version of the GFM compatible with the SCEC UVCN



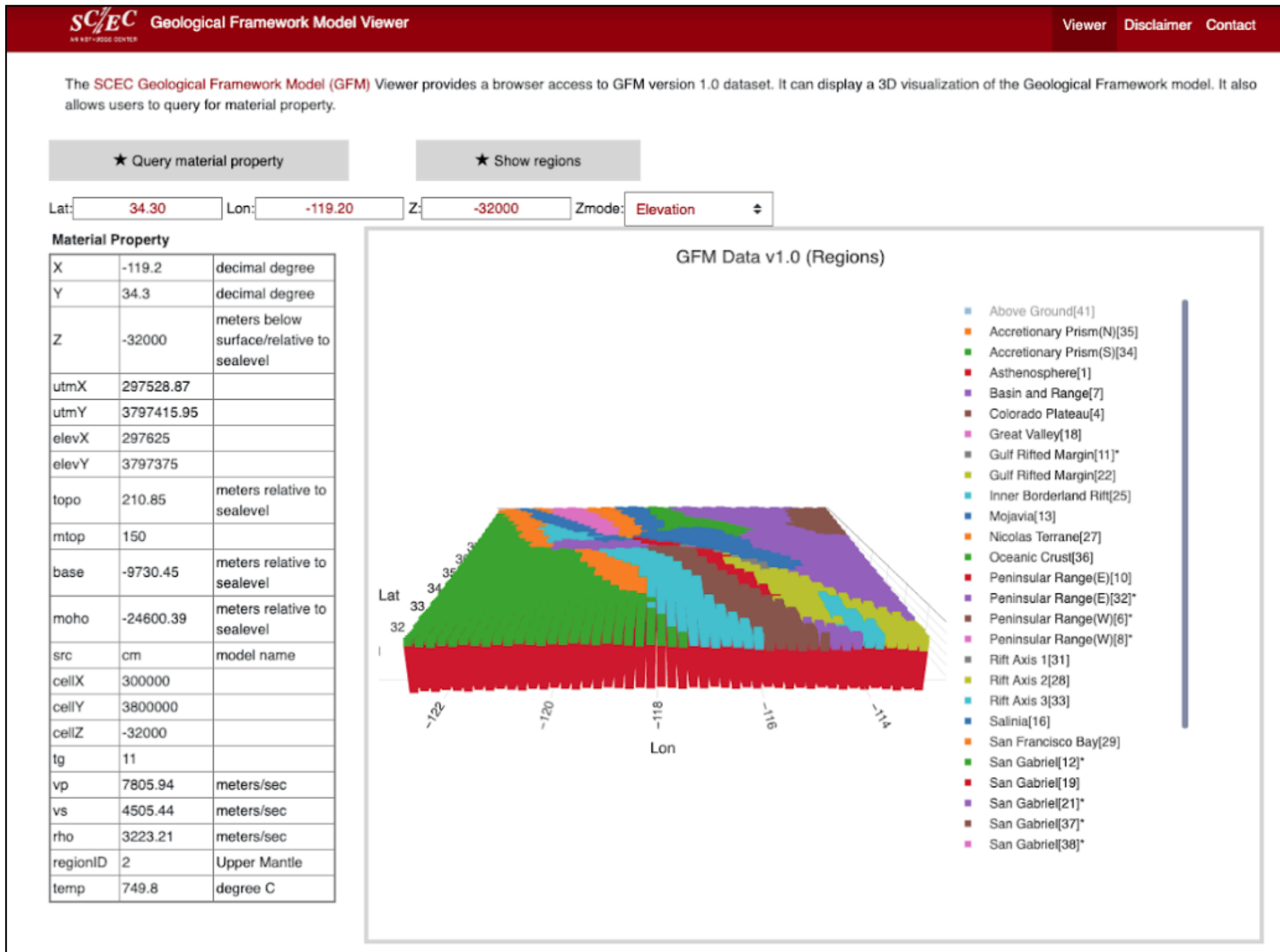
It is possible to use the GF model as is to populate a finite element model (for example) but the UVCN or similar will be needed as the GF moves beyond 1D lithology profiles within terranes with vertical boundaries. Hence, the gridded “LEGO” version (right) and the query tool (below).

A GFM view and query tool is in development

Screen capture 1: querying the GFM



Screen capture 2: displaying the GFM (gridded version)



GFM view/query tool is being developed by Mei-Hui Su and Phil Maechling, with input from John Shaw, Andreas Plesch and Mike Oskin