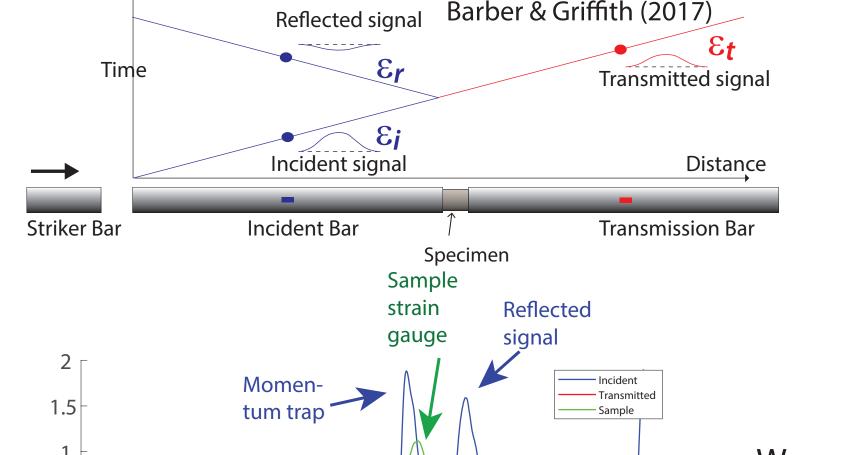


## ABSTRACT

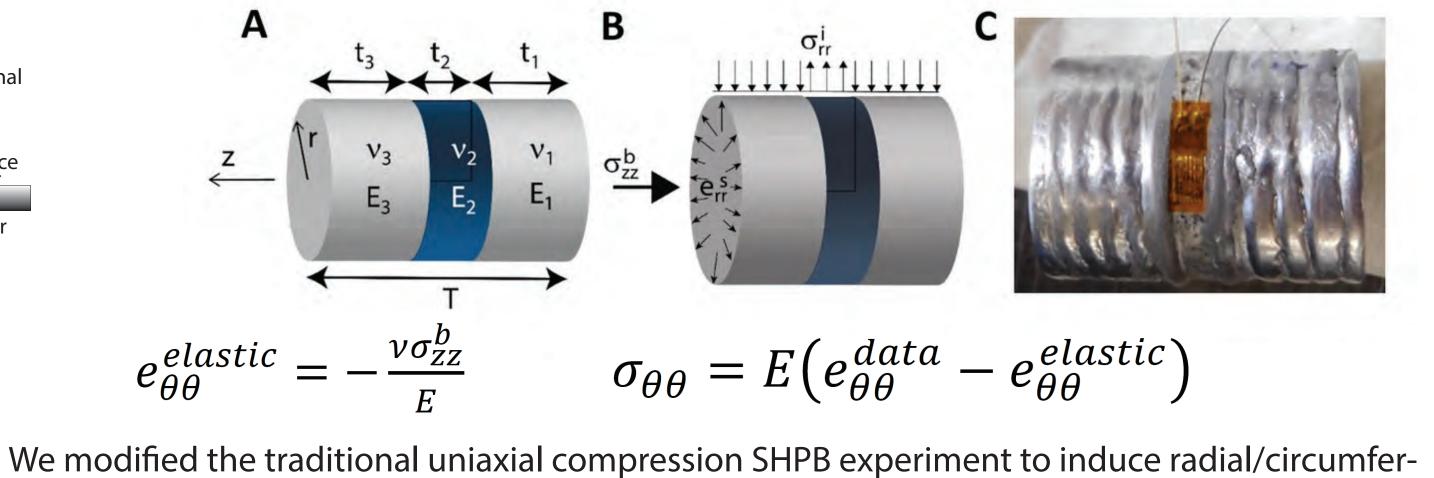
The presence of pulverized (highly fragmented, but weakly strained) zones extending 100-200 m from major strike slip faults, including the San Andreas Fault, have been attributed to impulsive compressive stresses associated with propagating earthquake rupture tips. However, theoretical and experimental evidence suggests that such zones may be formed on the transient tensile side of passing ruptures. These pulverized damage zones represent long-lived inelastic off-fault deformation that affect fault dynamics throughout the seismic cycle. We explore the tensile origin of pulverized fault rocks associated with major strike slip faults through a modified Split-Hopkinson Pressure Bar (SHPB) experiment that induces 2D isotropic tension. In the experiments, a sandwich sample configuration is used in which a rock disk is bonded between two cylinders composed of more compliant material such as lead or polycarbonate. Axial shortening during experiments results in radial and circumferential tension in the rock disk due to radial expansion of the compliant end materials. Experiments on both porous granular (sandstone) and crystalline (granite) rocks enable us to evaluate variations in tensile stress based on the rock type. We validate strain and strain rate histories collected on SHPB strain gauges using high speed photography and digital image correlation. Our modified SHPB experiments on Westerly Granite show that at strain rates of 25s<sup>-1</sup> to 170s<sup>-1</sup>, the rock fails by an isotropic pattern of polygonal fractures. Under similar conditions, deformation of Berea Sandstone is accommodated by distributed grain boundary failure and pore space expansion, therefore preventing fragmentation by fracture growth. These results explain asymmetric off-fault damage observed in the field where crystalline rocks 100-200 m from the core of the fault are pulverized, but adjacent porous sedimentary rocks appear to be relatively undeformed.

## **EXPERIMENTS**

The Split Hopkinson Pressure Bar (SHPB) is a reliable high strain rate loading technique used to assess the dynamic strength and constitutive response of rock. A uniaxial compressive wave is generated by striker bar impact with the incident bar and is recorded by strain gauges on the incident and transmission bars. This results in a simple load history described by single compressive sinusoidal loading and unloading cycle.



Modified sample configuration to induce radial/circumferental isotrpic tension



### PROBLEM

Brittle fragmentation is a fundamental process in the near-tip field of propagating earthquakes. This damage creation - leaving wide swaths of pulverized rocks around large strike slip faults *changes elastic moduli*, results in *rapid slip rate fluctuations* and high frequency content in seismic waves.



#### Rempe, Mitchell, et al. JGR (2013)

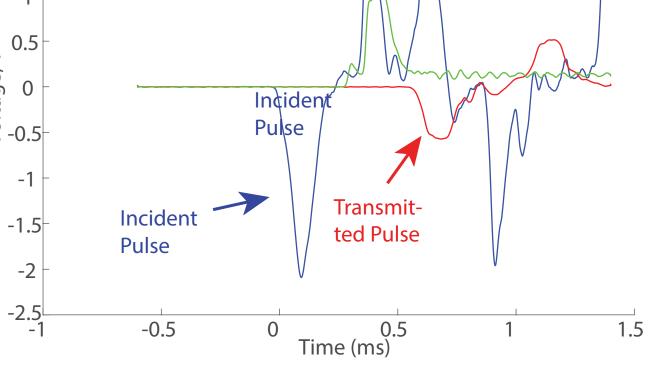
Pulverized damage zones can be 100m thick, but the mechanism for rock pulverization at distances this great from the principal slip zone is unclear:

Using voltage time series from strain gauges, we can recontruct macroscopic axial strain, strain *rate*, and *stress* histories and quantify the *energy* dissipated during sample failure ( $W_d$ ).

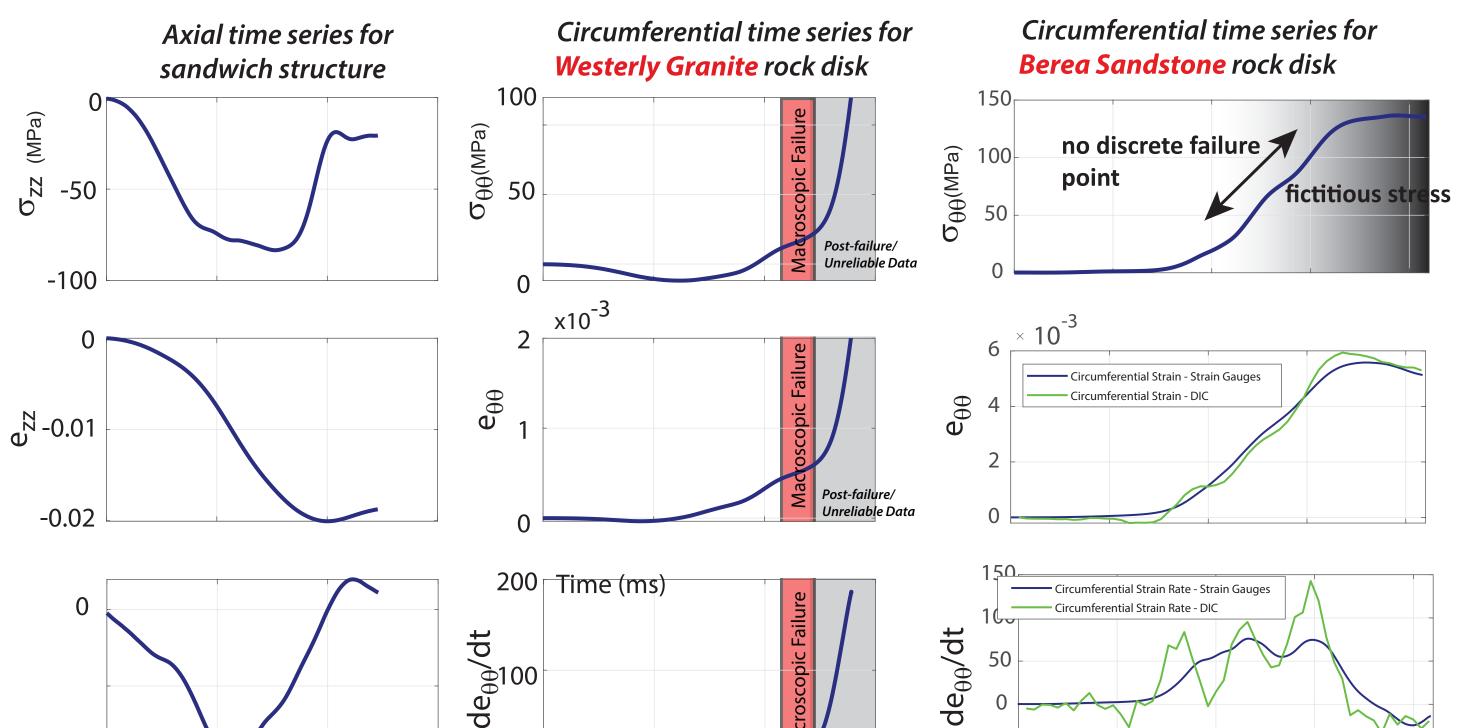
# RESULTS

### Stress, Strain, Strain-Rate Timeseries

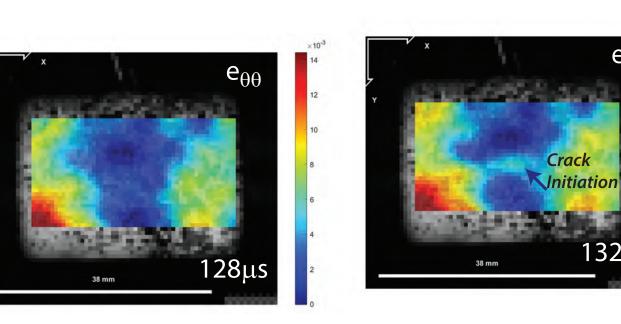
- Example of axial and circumferential strain rate, strain, and stresstimeseries:
- Compressive stress and shortening strain are negative.
- Axial time series represent overall axial values forsandwhich sample; *Circumferential time series repre-*
- sent values for rock disk
- Overall plastic deformation in axial stress plot controlled by yielding of lead
- Brittle failure of rock disk shown by rapid slope increase in circumferential strain rate

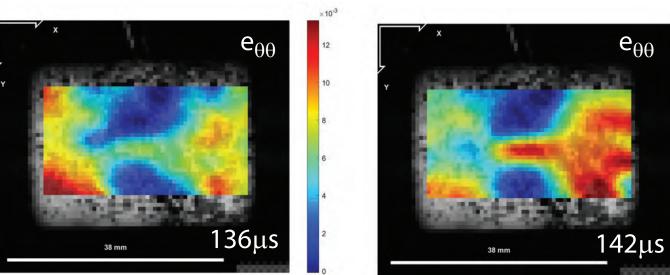


ential tension in a disk-shaped rock specimen. By sandwiching rock disk between two compliant, low compressibility materials (lead cylinders), axial compression causes lead to flow and expand radially outward, pulling the rock specimen apart. This is intended to mimic quasi-isotropic tensile pulses carried by the propagating rupture tip. Tensile failure results in **polygonal** fractures in granite and disaggregation in sandstone that are markedly different than fracture patterns produced in traditional unaixial compression experiments. During the experiments, we can monitor circuferential stress using strain gauge mounted directly on the rock. In all experiments, lead disks are 10-15 mm thick and rock disks are 5-10mm thick.



#### Strain Fields from Digital Image Correlation for Westerly Granite

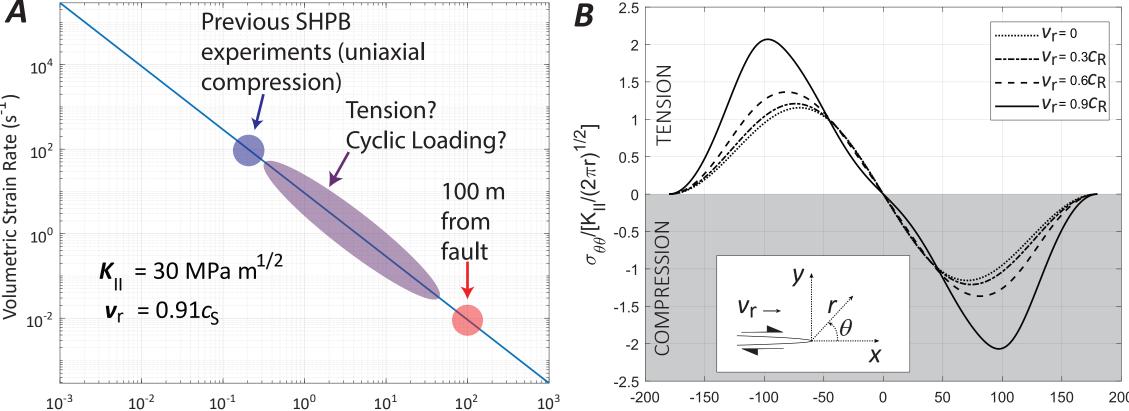




Compressive rock strength increases dramatically with strain rate, and rocks pulverized under compressive loads in the lab require strain rates of ~10<sup>2</sup> s<sup>-1</sup>.

These conditions are only expected within cm's of faults, yet field evidence ■ suggests that rock fragmentation occurs tens of meters from faults.

Strong crystalline rocks seem to be preferentially pulverized compared to weaker sedimentary rocks



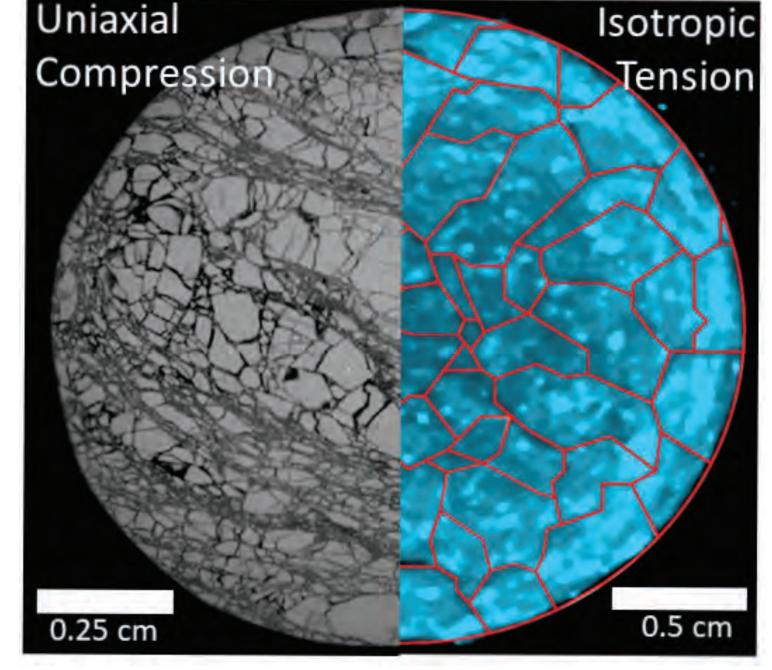
Distance from Fault (m) Peak strain rates as a function of distance from fault (blue line)

 $\theta$  (Degrees Relationship between Mode II near-tip stress field and rupture velocity

# **OBJECTIVES**

We hypothesize that far-field coseismic pulverization occurs via impulsive near-isotropic tensile stresses associated with passing earthquake ruptures, following evidence from numerical models (Xu & Ben Zion, GJI, 2017). We have designed an experimental technique to simulate 2D isotropic tensile stresses in the lab (see Griffith et al., JGR, 2018 for more details), and here we test the following questions:

### **Post-Mortem Rock Structures**



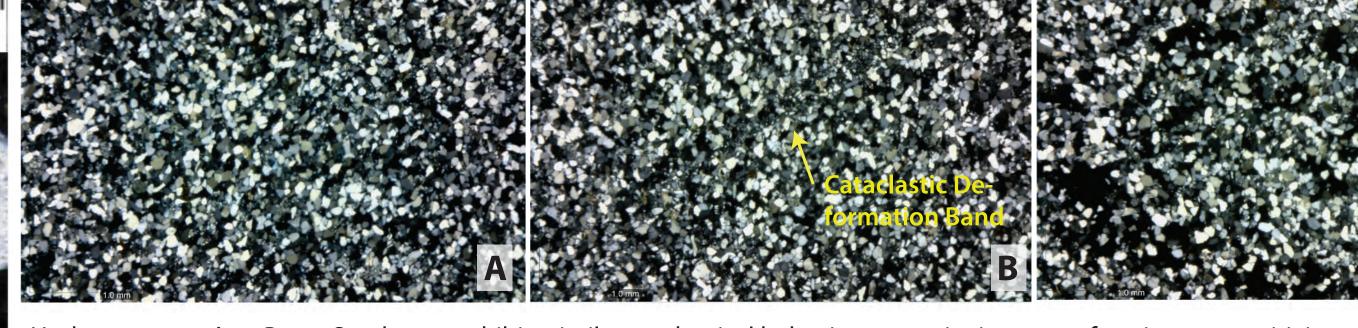
Comparison of experimentally-produced fracture patterns under axial compressive (left) fragmentation of Arkansas Novaculite vs. isotropic tensile fragmentation (right) of Westerly Granite

### Dynamic Tensile Rock Strength

120				11111		1 1 1 1		11111		
120		This Study								
n	⊢ <b></b>	Griffith et al. (2018)								
9		Static Strength, Shmidt & Lutz (1979)								







Under compression, Berea Sandstone exhibits similar mechanical behavior to granite in terms of strain-rate sensitivity to strength but failure under light confinement occurs by formation of cataclastic deformation bands (B). Under isotropic tension, failure occurs by grain boundary failure and disaggregation (C) with no distinct brittle failure strength. This latter deformation mechanism may be impossible to identify in outcrop.

## CONCLUSIONS

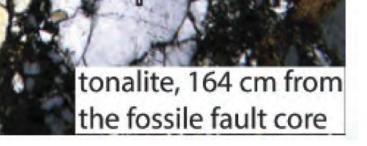
New technique tests dynamic tensile strength with easy sample prep

Technique simulates rapid expansion under shallow confinement, scalable to natural pulverized fault zones

Transition between static and dynamic tensile rock strength btwn 10° & 10<sup>2</sup> s<sup>-1</sup> for Westerly Granite

*No descrete failure occurs in similar experiments for Berea Sandstone* 

*Tensile and Compressive fragmentation produces distinctly different microstructures in both* Fragmented Tonalites from the Clark crystalline and granular rocks Read More House Canyon, western Salton Trough IMPLICATIONS Here: (Whearty, Rockwell & Geary, 2017) Tensile fragmentation can explain heavily fragmented rocks at large distances from faults AGU100 ADVANCING EARTH AND SPACE SCIEN ournal of Geophysical Research: Soli Granular sedimentary rocks appear to simply dissagregate during Range from static (Schmidt & Lutz, 1979) to shock (Cohn & Ahrens, 1992) *impulsive tensile loading<sup>,</sup> which may explain why stronger crystal* line rocks are preferentially fragmented in pulverized zone tensile strength of Westerly Granite ain rates on the order of 10<sup>2</sup> s<sup>-1</sup>, point rs of magnitude smaller. We present evidence from new modified Split-Hopkinson Pressure B occurs in the range 10<sup>0</sup> to 10<sup>2</sup> s<sup>-1</sup> in Asymmetric damage patterns with microstructures diagnostic of our experiments. Berea Sandstone coseismic stress state may imply long term preferred rupture direc may form in a state of isotropic tension. The resulting fragment size is independent of strain rat does not exhibit an abrubt tensile ighness. Our results provide a solution to the strain rate-distance scaling proble tivity experiments and field observations of pulverized rocks and also explain the asymmetric distribut ulverized fault rocks about strike-slip faults



strand of the San Jacinto Fault in Rock



What is the relationship between strain rate and rock strength under tensile loading?

How do crystalline granitoid rocks differ in their mechanical response to impulsive tensile loads compared to granular sedimentary rocks?

How are these mechanical behaviors reflected in rock structure?

How well do the experimental results scale to the natural prototype (pulverized damage zones)?

