Detecting and Reducing the Effects of Postseismic Deformation in Observed GPS Time Series Using a Hybrid Forward/Inverse Method in Southern California

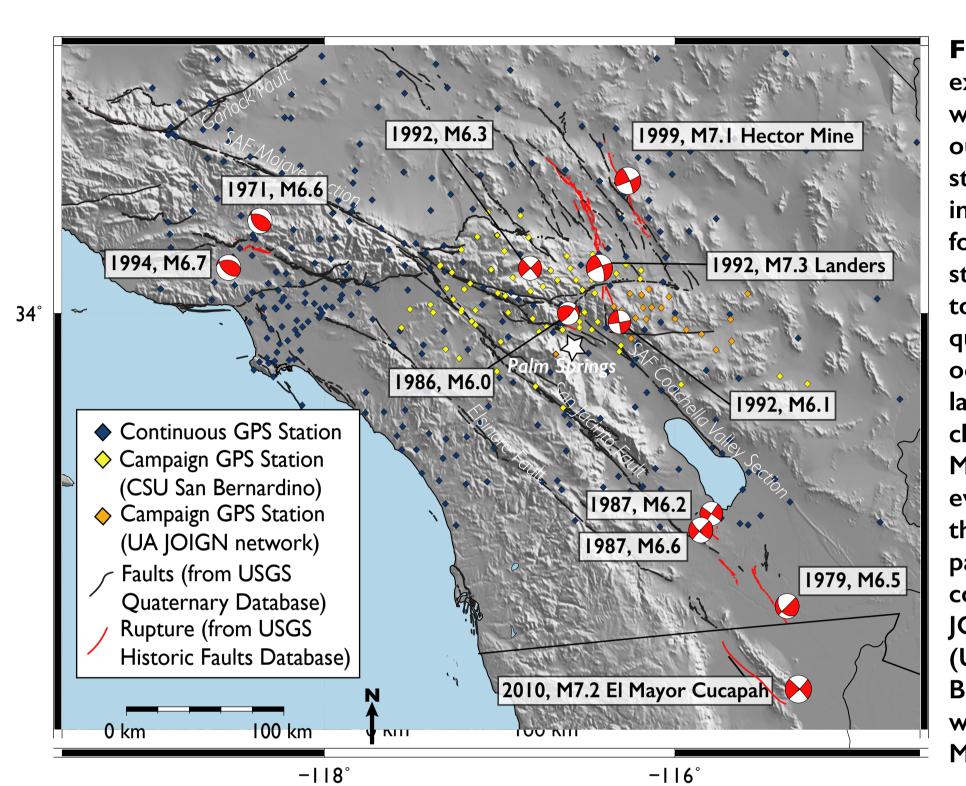
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Motivation

- + Viscoelastic postseismic deformation caused by large magnitude earthquakes can last for years to decades after the main event.
- + High precision GPS instruments in southern California may still be observing postseismic deformation from historical and recent earthquakes in the southwest U.S.
- + This type of long-term deformation can bias both crustal GPS velocities and geodetic-based fault slip rate estimates, but can be difficult to detect in GPS time series.



works of continuous and campaign style GPS stations in southern California used in this study, in addition to all M6.0+ earthquakes that have occurred in the last 50 years (excluding the recent M7.1 Ridgecrest event which is off the map). Campaign GPS data comes from the JOIGN network (UA) and the San Bernardino network run by Sally McGill (CSUSB).

Wang et al. (2006)

Wang et al. (2009)

SRCMOD Database

JSGS Earthquake

Wang et al. (2006)

Method

GOAL: To identify and reduce the effect of ongoing postseismic viscoelastic relaxation on GPS time series

(I) FORWARD MODELING Choose Earth elastic and viscosity structure | Figure 2 | Broermann (2017)

Calculate Green's Functions with PSGRN software

Organize finite fault models for all 217 earthquake events

Calculate model surface displacements for each event at locations of network GPS stations using PSCMP software

Determine which events measurably impact our GPS time series (yr. 2000-2018)

Subtract those events' postseismic displacements from modern time series

(2) INVERSE MODELING

Estimate logarithmic postseismic terms

for recent conspicuous events

from **both sets**, to assess success

of forward model

Estimate velocities from forward model postseismic reduced time series

Estimate velocities from **observed**, | Figure 5 un-reduced time series

Figure 7

Figure 3

Figure 4

| Figure 5 |

Figure 6

TSVIEW / TSFIT

Hector software

Hector software

Method

(I) Forward modeling earthquake displacements

+ PSGRN + Earth Model

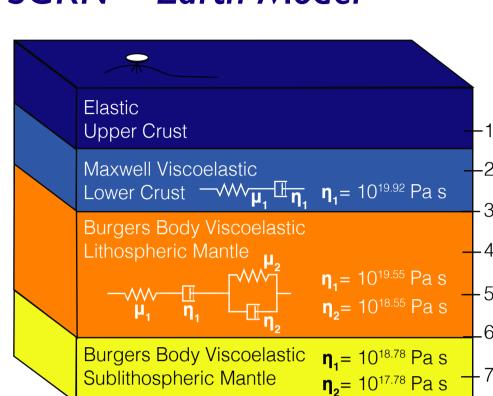


Figure 2. Our chosen viscoelastic Earth structure from This is a layered, laterally homogenous lated in PSGRN -70 km (Wang et al., 2006).

+ 217 ≥M6.0 Earthquakes modeled in PSCMP

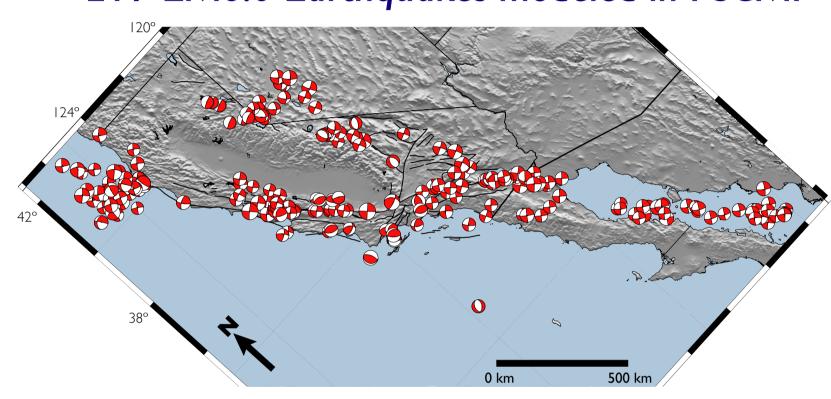


Figure 3. Locations of 216 ≥M6.0 earthquakes since 1800, for which we calculated surface displacements caused by viscoelastic postseismic deformation (one earthquake not shown as it is off map in Sonora, MX). These earthquakes and their rupture characteristics come from the Wang et al. (2009) catalog, and the USGS Earthquake Catalog when not present in the SRCMOD database.

+ 12 Earthquakes with ongoing displacements

Figure 4. Of all 217 earthquakes modeled using PSCMP and our chosen earth structure, 12 events still exhibited measurable ongoing year-cumulative surface displacements ≥0.35 mm in any year 2000-2018 for any station in our southern CA network. This figure illustrates the 0.35 mm contour of cumulative displacements in 2018 for each event, where the areas inside the contour had cumulative displacements >0.35 mm.

Results

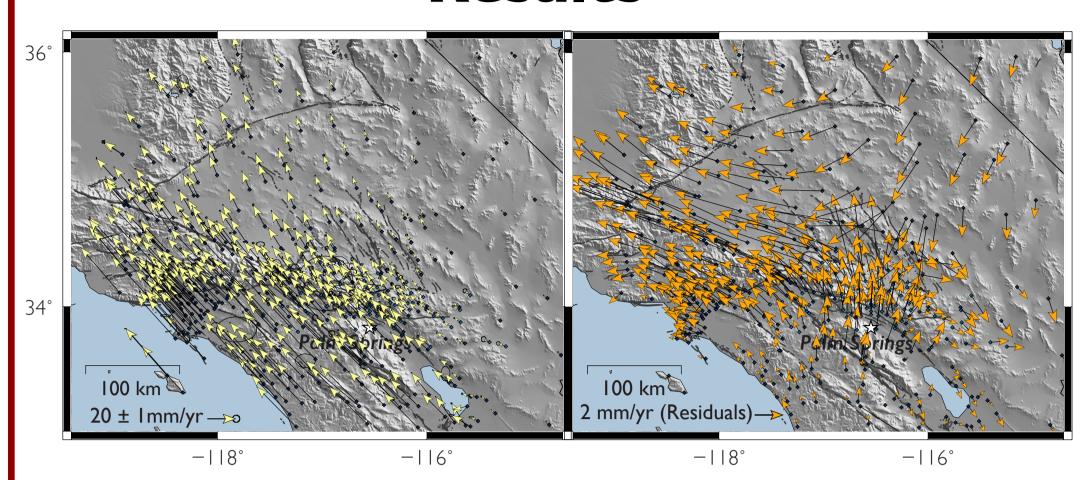


Figure 7. Final postseismic-reduced GPS velocity field in Stable North America Reference Frame (NAM08) (left, yellow) and the residual velocity field (observed velocities - reduced velocities) (right) showing the directions and magnitudes of the removed postseismic biases at each station (note factor of ten difference between scales). Error ellipses are 95% confidence

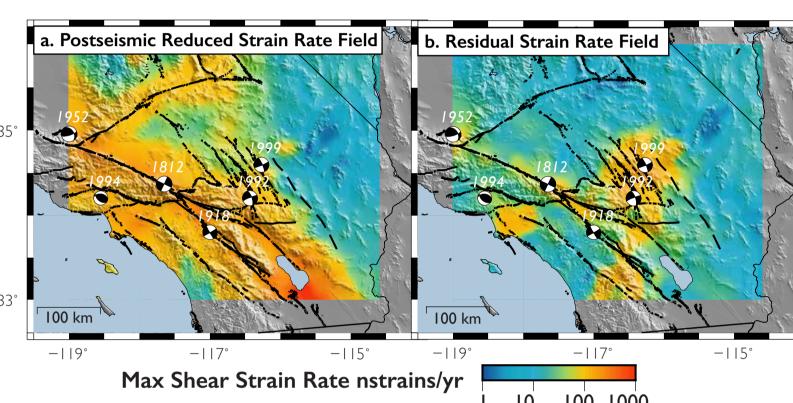


Figure 8. Shear strain rate maps calculated from the (a) reduced velocity field and (b) the residual velocity field. This illustrates the high strain rates in the ECSZ that have been reduced through this analysis.

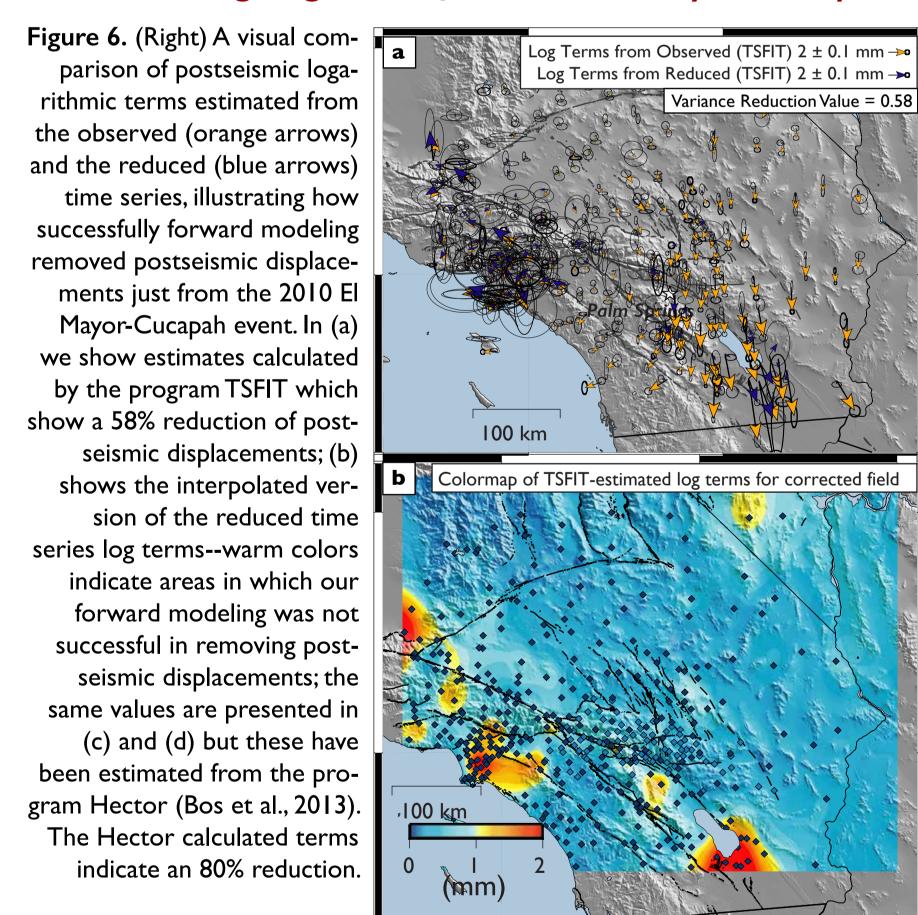
(2) Inverse Modeling: Reduced Time Series & Estimating Logarithmic Terms

+ Reducing Time Series & Estimating Velocities

Time (years) Time (years)

Figure 5. (above) Stacked time series plot of a selection of stations zoomed in to the 2010 El Mayor-Cucapah earthquake (vertical dashed grey line), comparing observed time series (at top of every pair, labeled with "Ob" at right) and forward model postseismic reduced time series (at bottom of every pair, labeled with "R" at right) that have been corrected for the 12 earthquakes shown in Figure 4. Stations BEMT, OPCX, and P511 show examples of successful reduction of postseismic displacements, while station CRRS shows an example of a reduction in which most of the postseismic displacements have been removed, but the coseismic jump has been overcorrected. Station P556 shows an example of the case in which our postseismic reduction have little effect on the 2010 displacements.

+ Estimating Log Terms for 2010 El Mayor-Cucapah



Log Terms from Observed (Hector) 2 ± 0.1 mm →

Log Terms from Reduced (Hector) 2 ± 0.1 mm → Variance Reduction Value = 0.80-118°

Implications for fault slip rates?

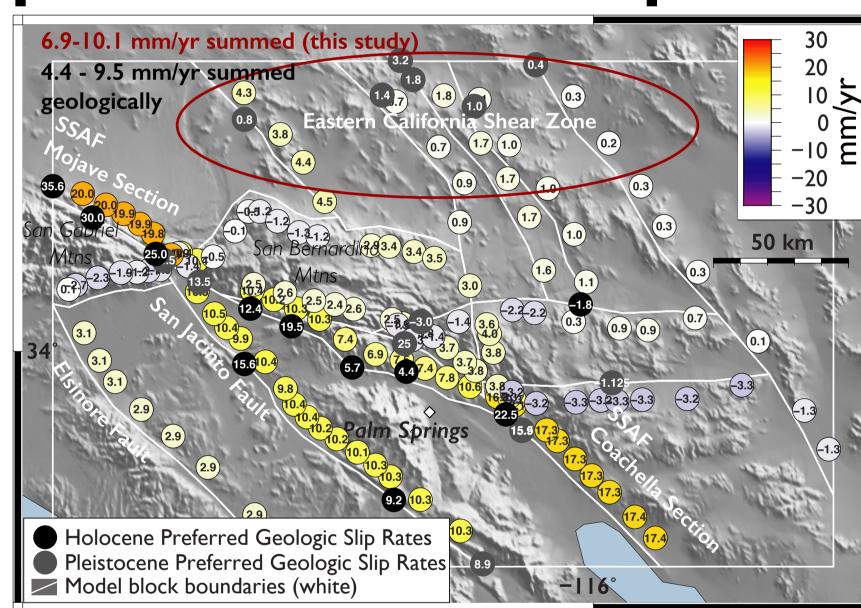


Figure 9. Preliminary slip rate estimates from inverting our postseismic reduced velocity field using the fault block geometry outlined in white using tDEFNODE (McCaffrey, 2005). The most intriguing result is that our geodetic-based fault slip rate values match or nearly match those in the Eastern CA Shear Zone (compare to grey circles with preferred geologic slip rates).

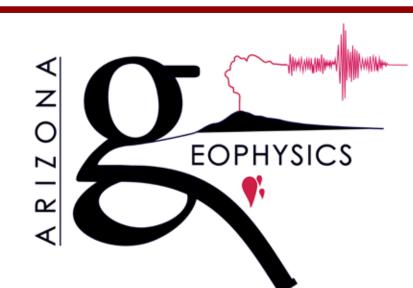
Summary

- + We identify 12 historical and recent earthquakes that have greater than 0.35 mm/yr of summed cumulative displacements in any year 2000-2018, indicating measureable ongoing displacements.
- + We successfully reduce postseismic deformation for the 2010 El Mayor-Cucapah earthquake by up to 80% (58% measured through values from TSFIT and 80% measured throught values from Hector) by removing displacements from these 12 earthquakes from modern time
- + Preliminary block modeling using our postseismic reduced velocity field indicates possible agreement between geologic and geodetic slip rates in the Eastern California Shear Zone

References

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