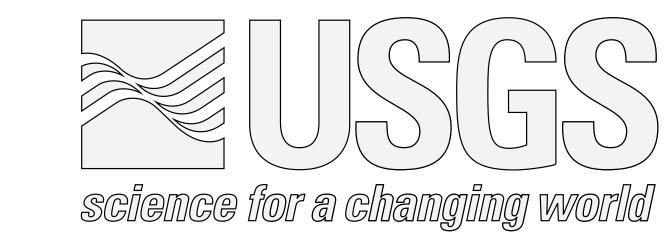
An automated method for building non-planar 3D fault models from earthquake hypocenters

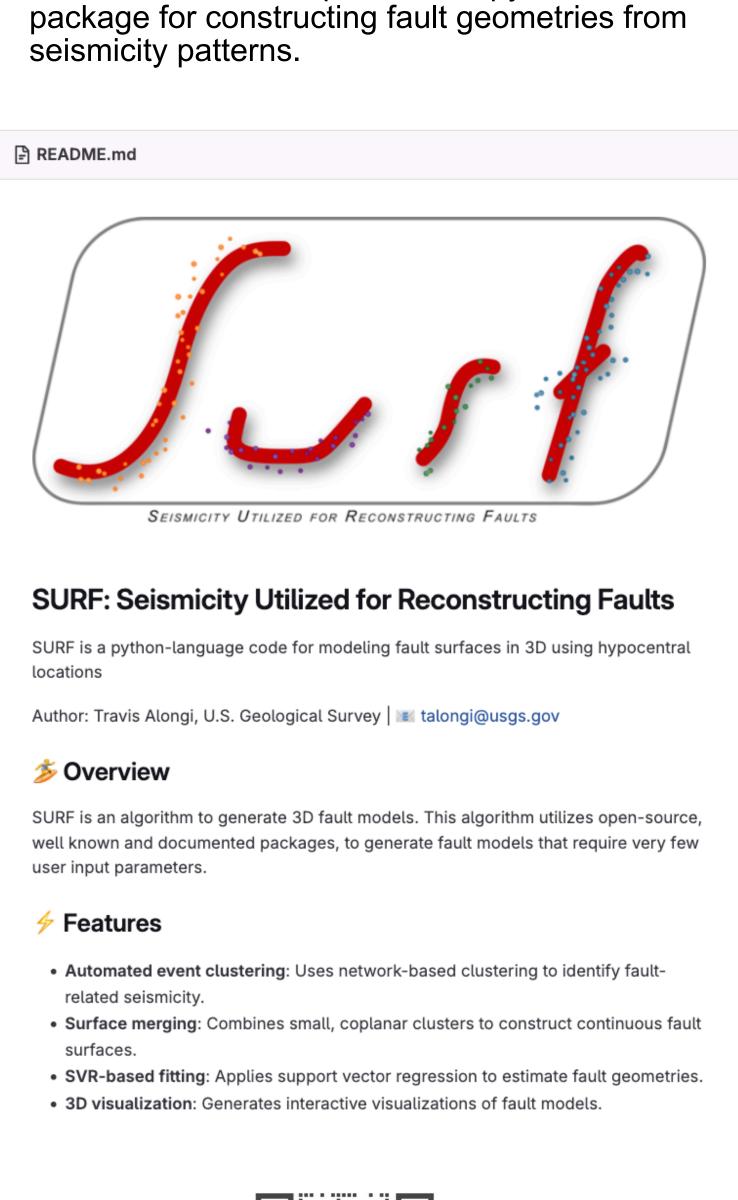
Travis Alongi, Robert Skoumal David Shelly, Alexandra Hatem

Geologic Hazards Science Center, Golden, CO



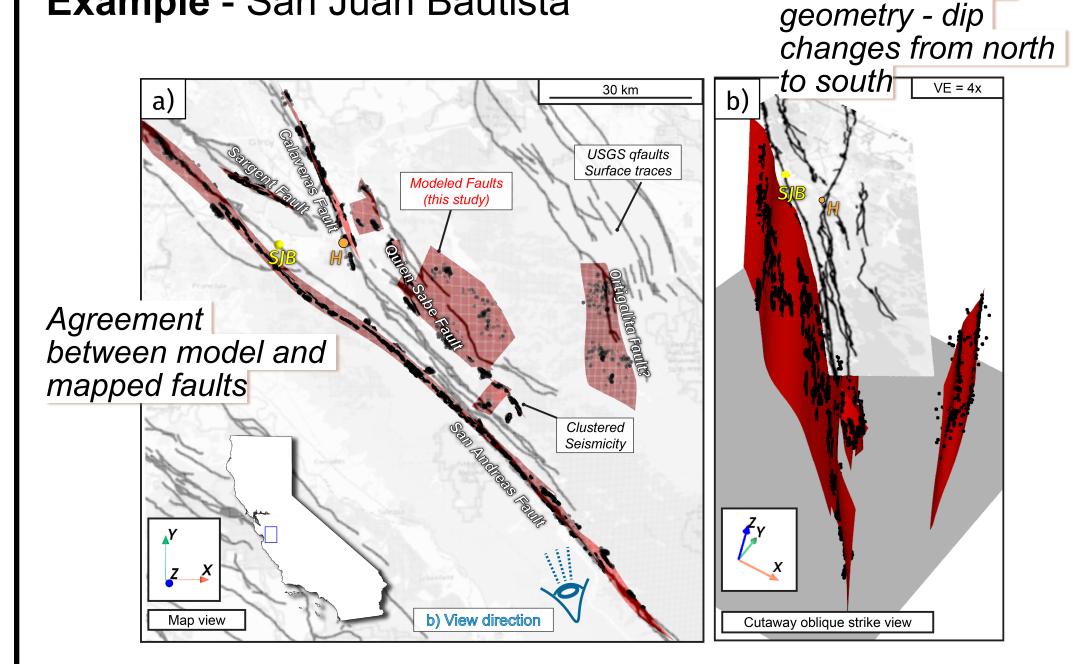


Now available! an open source python

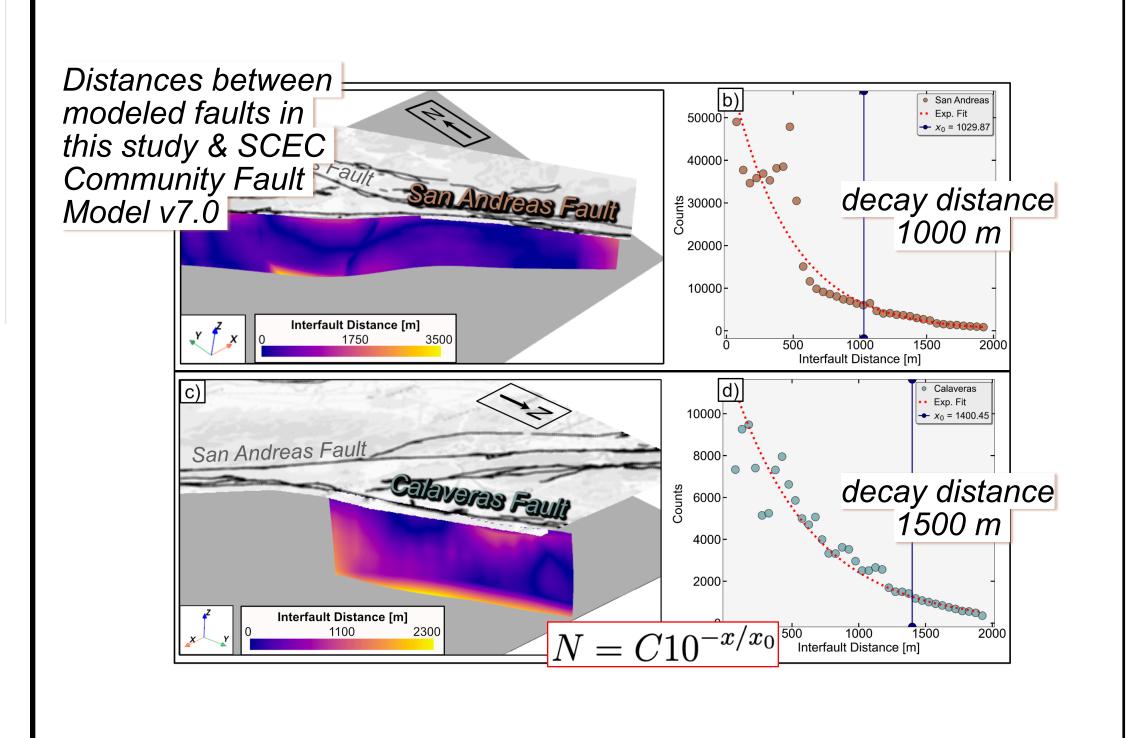


RESULTS

Example - San Juan Bautista



How do results compare to existing fault models?



Non-planar fault

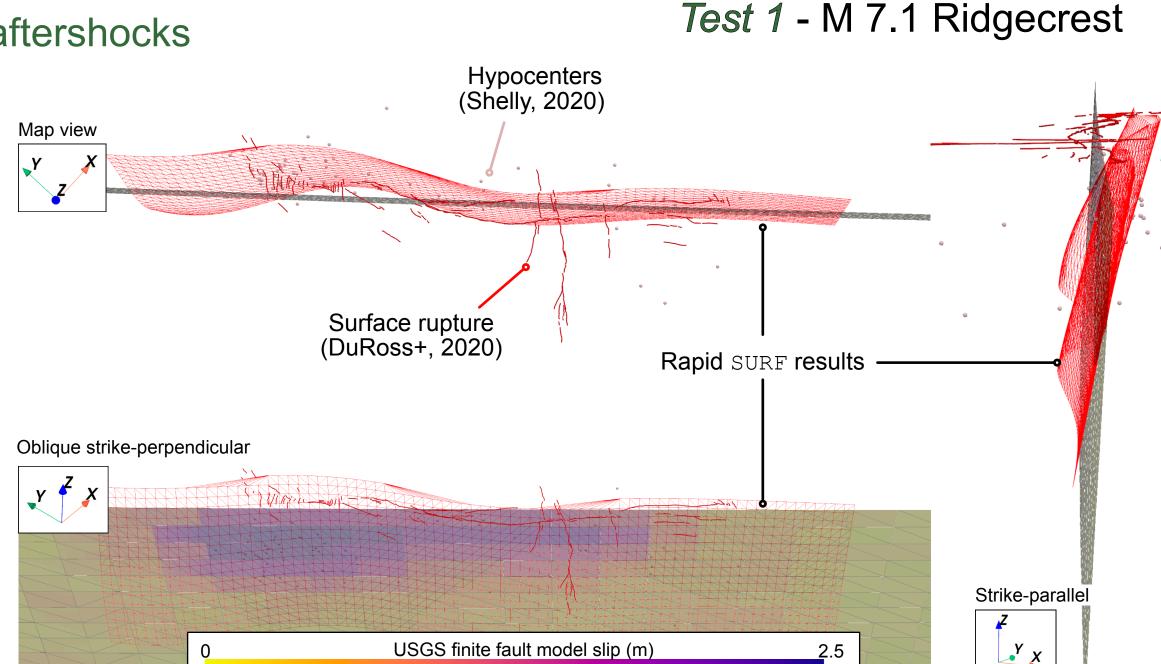
Hazards application

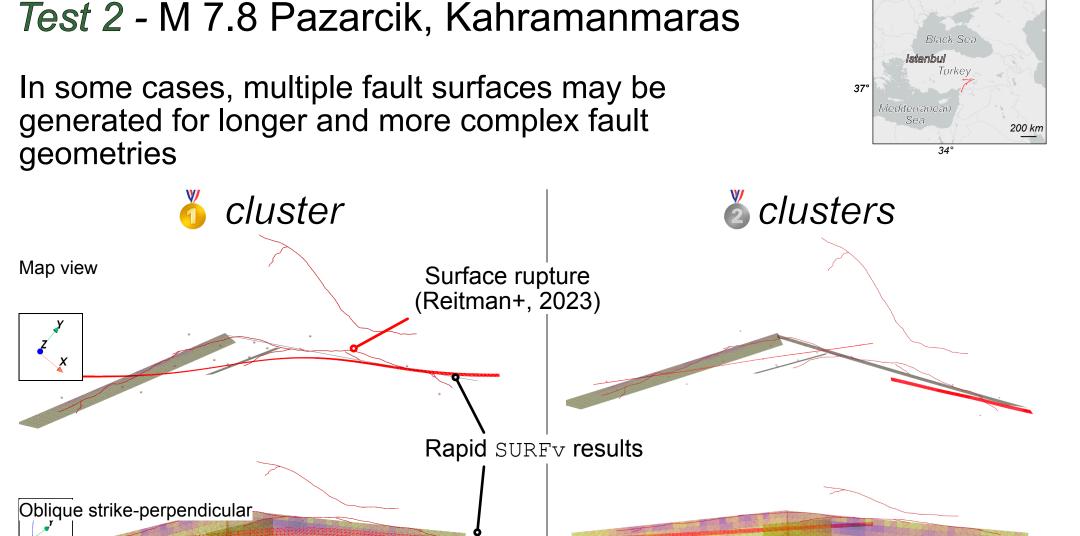
Rapid fault geometry calculation with early aftershocks

We can speed up the finite fault inversion process using early aftershocks to help define the fault geometry and build the mesh with

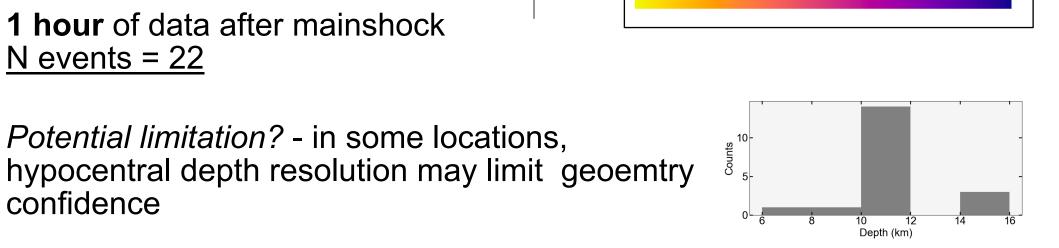
15 minutes of data after the mainshock N events = 66

Note- agreement between the USGS finite fault result & the rapid SURF results





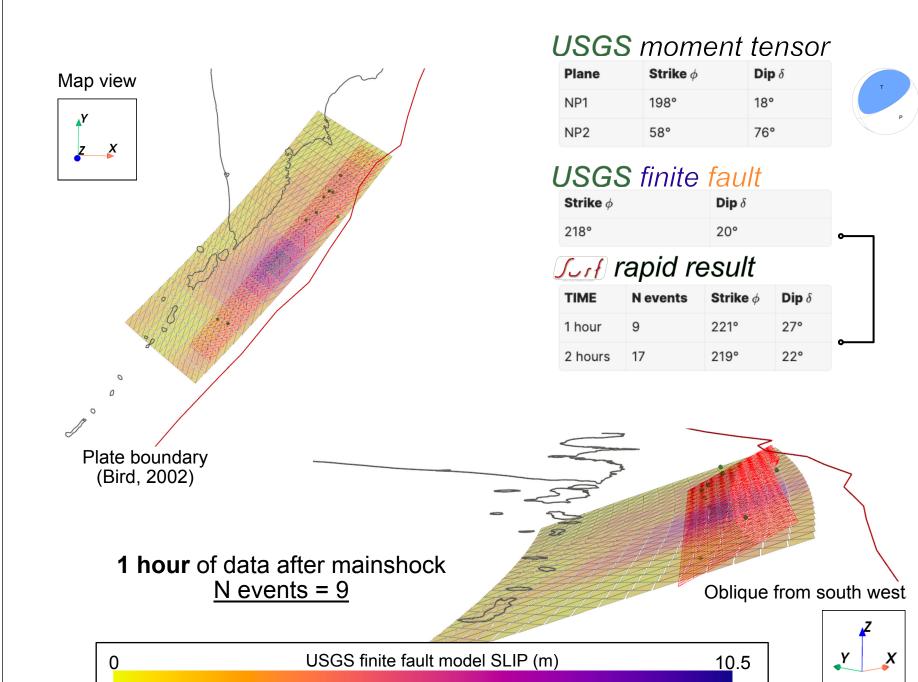
1 hour of data after mainshock N = 22



USGS finite fault model SLIP (m) 11.5



Works well for dipping subduction zone events. Note- comparable geometry in 1 to 2 hours (~ 15 events)

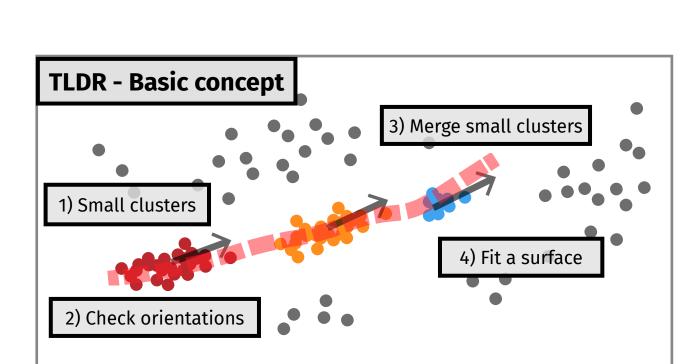


HOW IT WORKS

Try it out!

https://code.usgs.gov/esc/surf

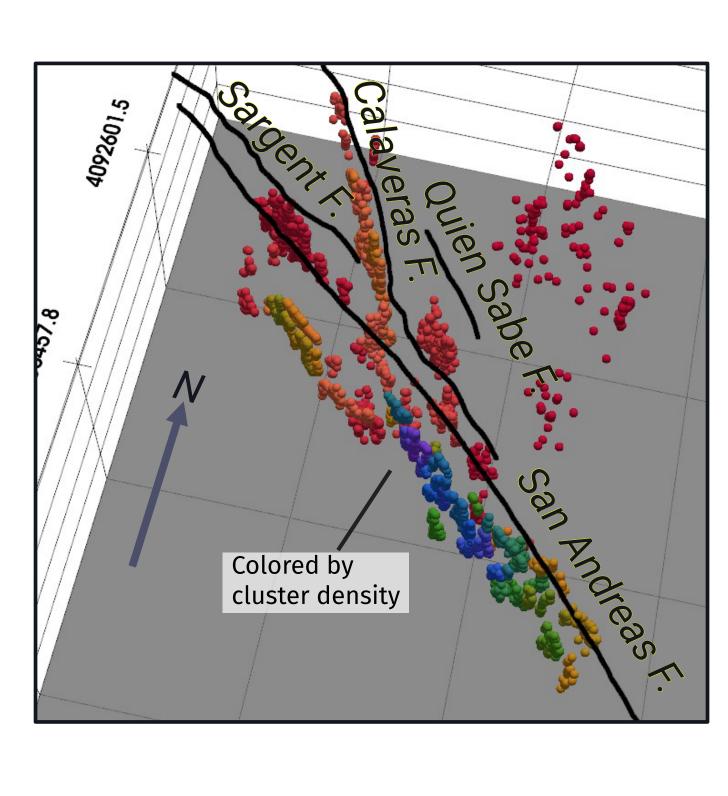
(Alongi et al., 2025 - SRL)



1. clustering

Hierarchical Density Based Spatial Cluster Applications with Noise HDBSCAN

Requires only one parameter- the minimum number of events to be considered a cluster We set this to 40 events allowing orientations to be determined and remain small



2. a) measure coplanarity

a)

Earthquake cluster merging process.

1. Fit clusters with a plane using RANdom SAmpling Consensus to determine its normal vector, n.

2. Next, the *k*, nearest neighbor clusters are identified and the direction vector is calculated.

3. normal vectors and direction vectors are used to construct the measurement matrix M, that quantifies coplanarity of neighboring clusters

Cluster Merging

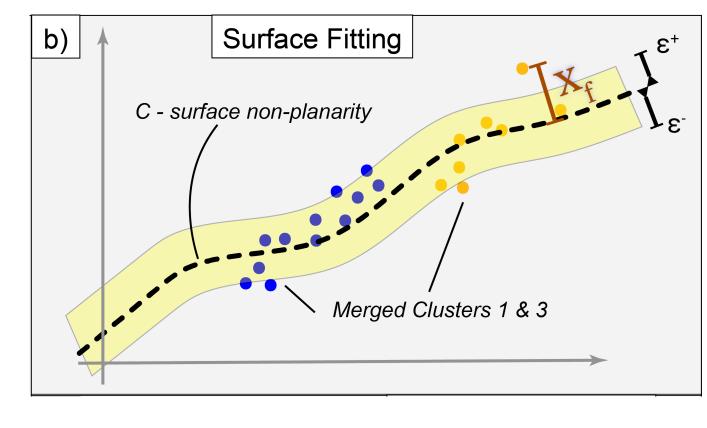
3. b) surface fitting

A cartoon illustration to help build intuition about the support vector regression (SVR) surface fitting parameters, \tilde{C} and ϵ .

confidence

C describes how non-planar the modeled fault surface is, with higher values indicating greater non-planarity.

ε describes the length scale near the surface where points do not impact the fit.

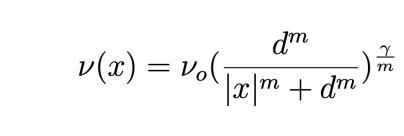


 $ec{N}=ec{n_i}\cdotec{n_{jk}}$ $\vec{D} = \vec{n_i} \cdot \vec{r_{jk}}$ $\mathbf{M} = |\vec{N}| \cdot (1 - |\vec{D}|)$

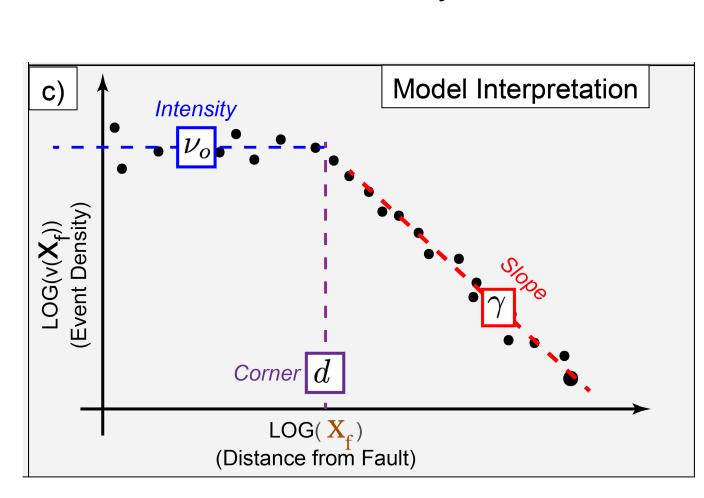
4. c) quality indicators

To assess the reliability of the algorithm's outputs, we measure the spatial distribution of earthquake density around the modeled fault surface.

The relationship between earthquake density and distance is an inverse power law relationship. We fit the following equation,

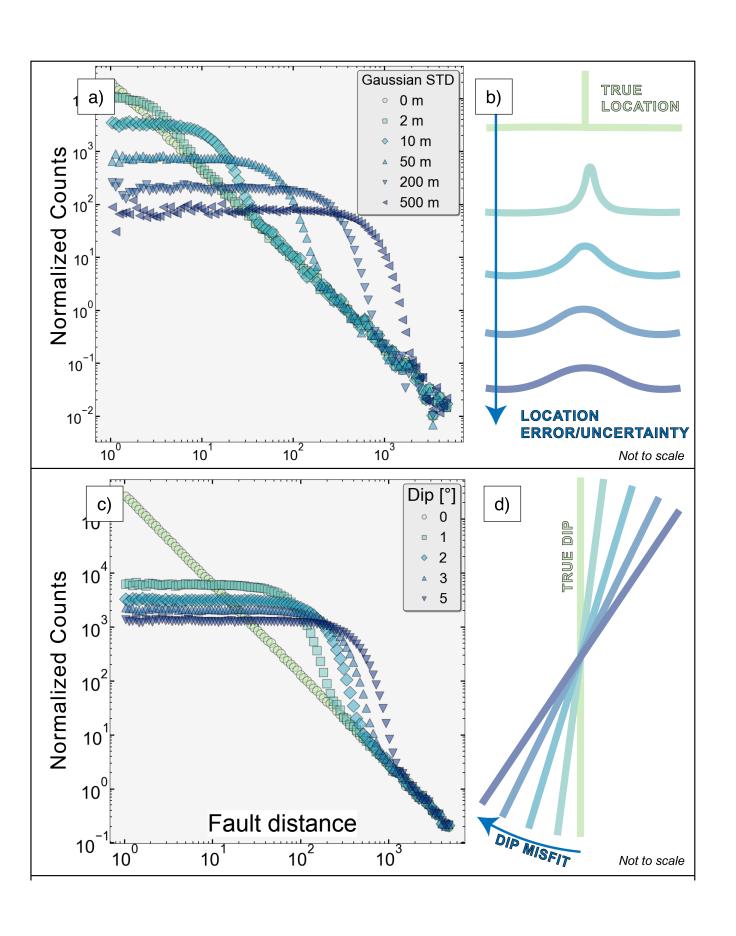


Schematically



5. Impact of misfit and unmodeled complexity on a power law distribution of earthquakes

The points (green circles) are prescribed an idealized distribution around the fault with lateral distances being power law distributed from a line.



Note, examples feature a similar result of moving the corner of the power law further from the fault with increasing unaccounted for fault complexity or fit errors.

This information is preliminary and is subject to revision. It is being provided to meet the need for timely best science. The information is provided on the condition that neither the U.S. Geological Survey nor the U.S. Government shall be held liable for any damages resulting from the authorized or unauthorized use of the information