

Creep Fronts and Asperity Interactions in Laboratory Earthquake Sequences Illuminate Delayed Earthquake Triggering

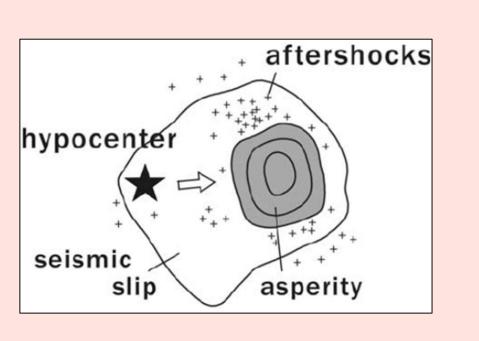
PENNSTATE.

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1. Background and Motivation.

Seismic asperities are fault sections that are stronger and more seismogenic than their surroundings. Their properties may control earthquake size, and interaction between asperities is important for aftershocks, and earthquake triggering.

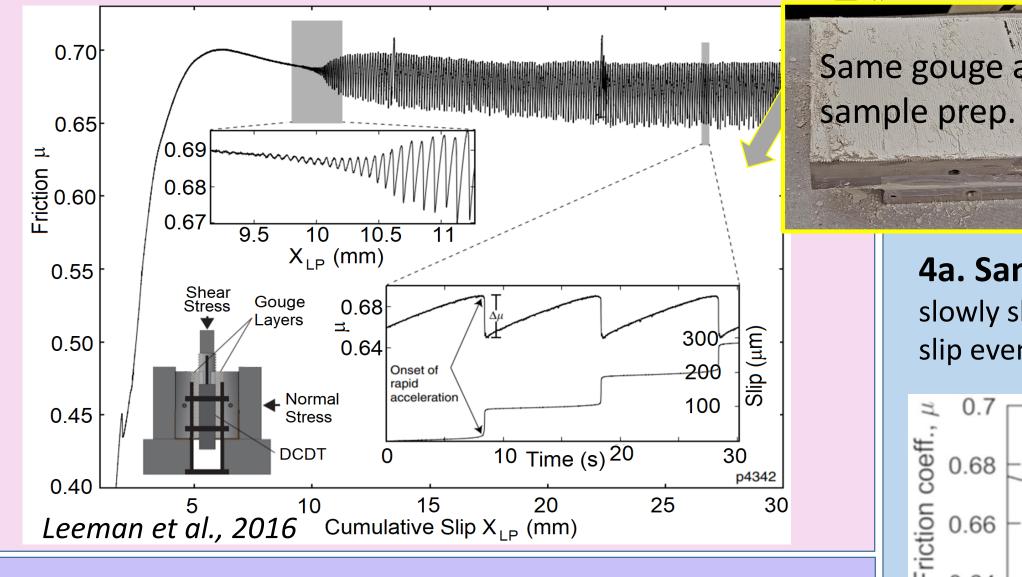


Velocity-strengthening

Yamanaka and Kikuchi, 2004

Chen and Lapusta, 2009

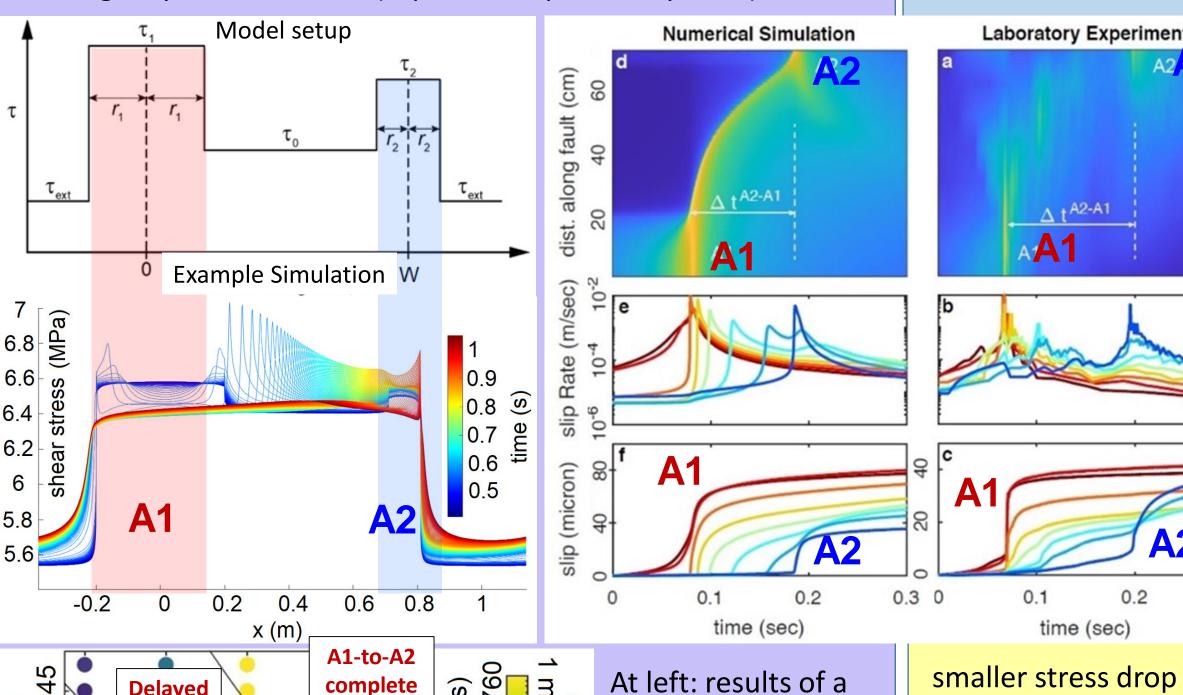
3. Previous quartz gouge research establishes friction **evolution.** With continued shear strain, friction properties transition from velocity strengthening (b-a < 0) to velocity weakening (b-a > 0) while D_c decreases from 15 mm to 2-3 mm (the fault becomes more unstable). Friction evolution is well characterized with previous, smaller experiments shown below (Leeman et al., 2016; Scuderi et al., 2017). Those friction parameters are inputs for corresponding numerical models.



6. Numerical Models: We simulate the 2-asperity interactions with fully dynamic, 2D spectral boundary integral method. With constant friction properties along the sample, the two asperities are regions with locally high shear stress. The size and stress state of the asperities were first tuned to yield spontaneous rupture of A1 followed by delayed rupture of A2. Then, keeping stress levels at A1 and A2 constant, we studied the triggering behavior as a function of the shear stress level (τ_0) between the two asperities and changing friction parameters. We model shear stress asperities instead of normal stress asperities. This only works because they are single-cycle simulations (asperities rupture only once).

6.4

Cumulative slip, x_{LP} (mm)



suite of simulations.

simulation, color-

coded by the A1-to-

A2 triggering time.

With more slip, the

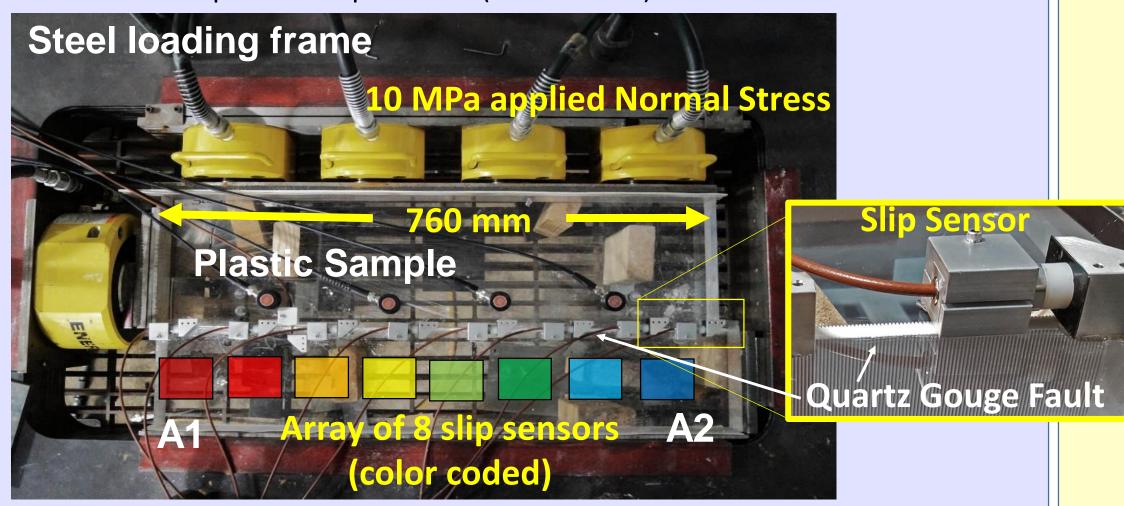
more sensitive to τ_0 .

triggering time is

Each circle is a

2. Laboratory Experiment with a Hybrid Sample.

Our sample has two dominant asperities: A1 (red) and A2 (blue) that result from locally high normal stress at the sample ends. A1 and A2 can rupture independently and interact through post-seismic slip in the form of creep fronts. We track slip with 8 slip sensors (color coded).

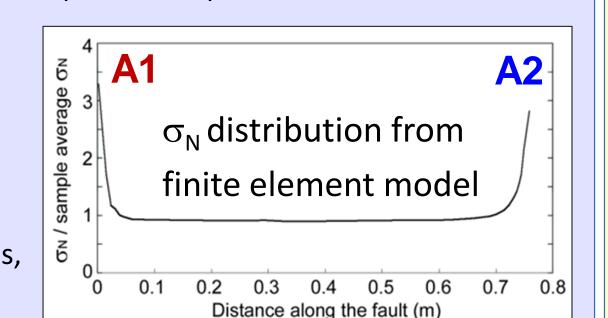


Using a recently-developed technique (Buijze et al., 2020), fault slip occurs within a 2 mm thick layer of powdered quartz (gouge) sandwiched between 760 mmlong plastic forcing blocks. Elastically, the plastic sample behaves like about 10 m

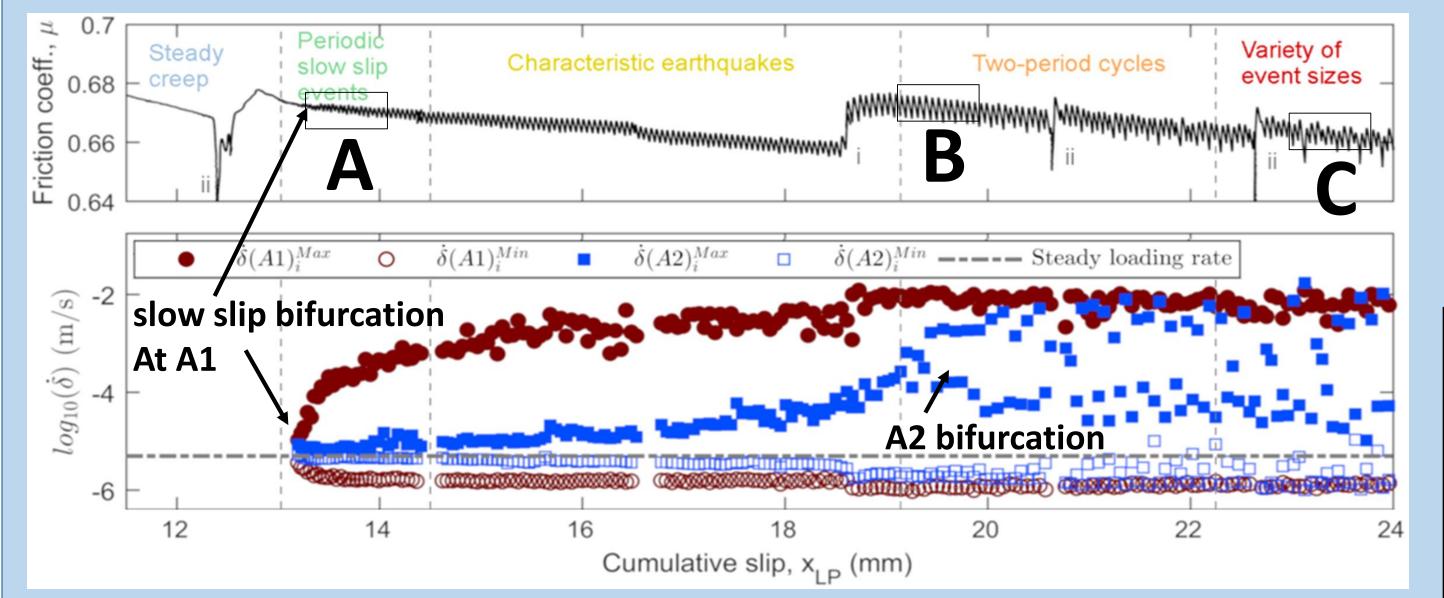
the quartz gouge, and is well characterized by rate- and statedependent friction. We use identical Same gouge and gouge layer

of rock, but friction is controlled by

preparation methods, so friction matches previous studies.



4a. Sample Behavior Evolves From Simple to Complex. We apply constant normal stress and slowly shear the sample at 6 μ m/s. Over many mm of cumulative slip, steady sliding gives way to slow slip events (A), characteristic stick-slip events, two-period cycles (B), and then a variety of event sizes (C).



4b. Creep Fronts Propagate Between the Asperities.

The figures on the left and right show zoom-ins of local slip measurements and corresponding slip rate maps showing rupture of one asperity, propagation of a creep front, and subsequent rupture of another asperity. Typically, A1 ruptures first (left), but sometimes it's reversed (right).

7. Conclusions: The simulations show that small changes in initial shear stress τ_0 (50 kPa) cause two orders of magnitude variation in the average triggering speed of the creep fronts. Fronts propagate faster at higher τ_0 , consistent with other recent studies (Garagash, 2021). Using strain gages, we find a similar stress dependence in our experiments: 20 kPa variation in initial overstress τ_0 can cause an order of magnitude change in A1-to-A2 triggering velocity.

The sensitivity to τ_0 explains the oscillatory behavior of the A2 bifurcation observed at $x_{IP} = 19$ mm: stronger A2 events reduced stress levels in the region surrounding A2 more significantly so the subsequent creep front propagates slowly and triggers weaker A2 events. Weaker A2 events have

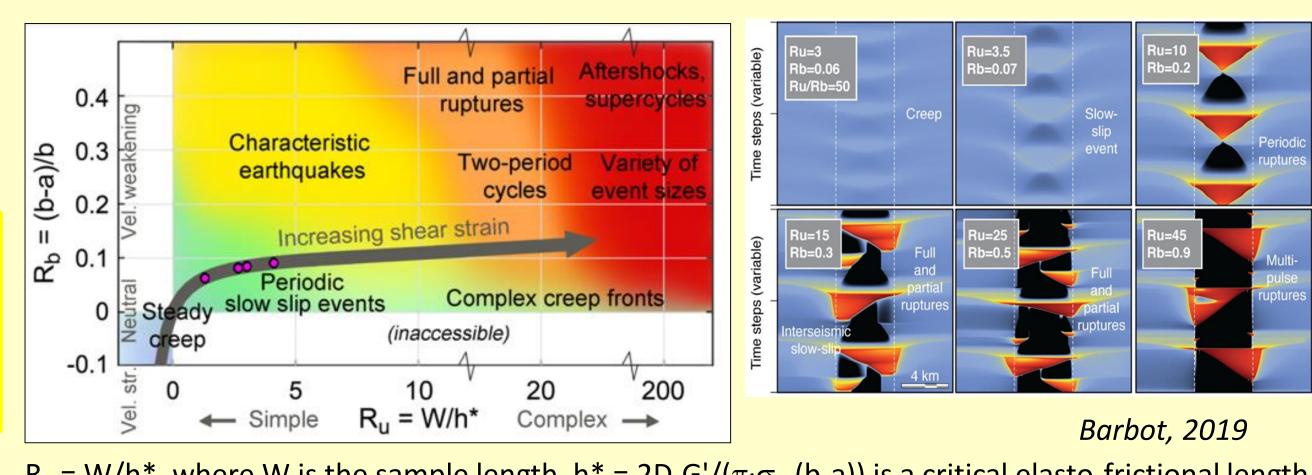
smaller stress drop and thus do not reduce stress levels as much, which primes the fault for faster subsequent creep fronts and more rapid triggering of stronger events.

With evolving friction (increased R_{II} and increased creep propagation distance L/L_b), creep front behavior becomes less sensitive to the size of the events that initiate them (hypocentral forcing) and increasingly sensitive to small variations in initial overstress stress levels (Garagash 2021). This is the likely reason for the increasingly complex behavior observed late in the experiment, despite highly periodic behavior early on.

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5. Mapping sample behavior in R_u – R_b space.

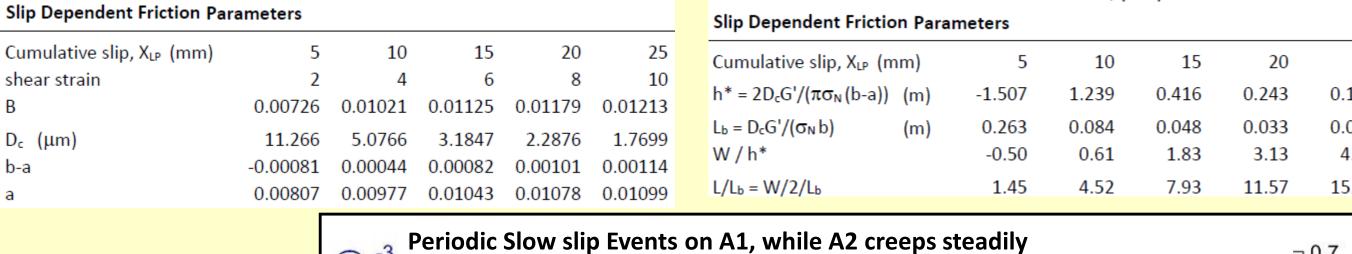
We map the previously established evolution of the friction parameters in R_{II}-R_b parameter space (gray arrow below). We then compare our sample behavior to that of the homogeneous model of Barbot (2019), shown below. Similar results were obtained by Cattania (2019).

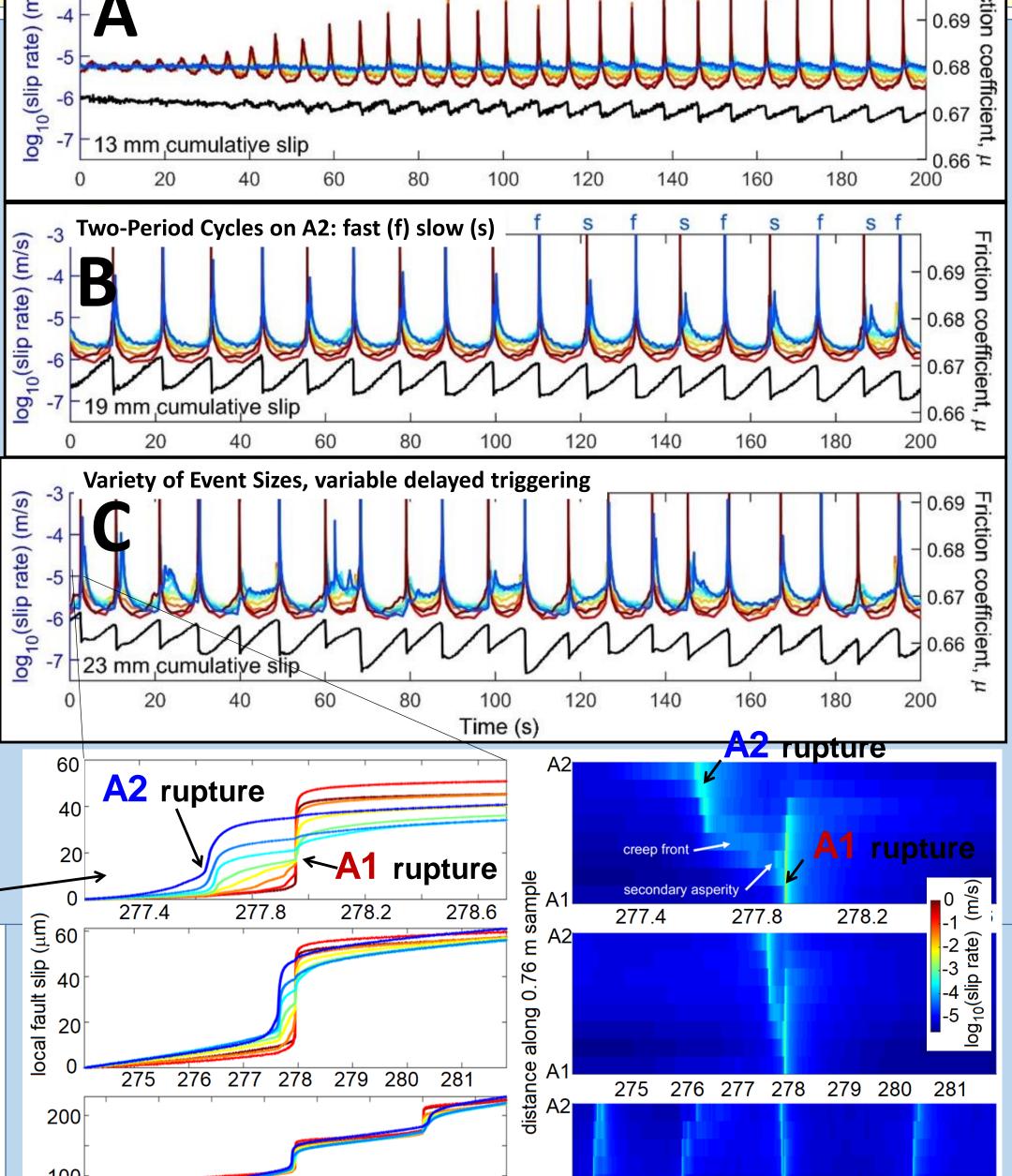


 $R_u = W/h^*$, where W is the sample length, $h^* = 2D_cG'/(\pi \cdot \sigma_{N})$ is a critical elasto-frictional length $R_b = (b-a)/b$, which describes the intensity of frictional weakening behavior. Slip independent Parameters

Our sample with two asperities produces a qualitatively similar behavioral progression to the homogeneous numerical simulations (Barbot, 2019; Cattania, 2019), but at lower R₁₁ levels.

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ters					
	G' = G/(1-ν)		1.69	GPa	
		Poisson's ratio, ν		0.35	
		shear modulus, G		1.1	GPa
		normal stress, $\sigma_{\scriptscriptstyle N}$		10	MPa
		Juinpie iei	19:11, **	0	





time (s)

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time (s)

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