

USC

MEASURING AND MONITORING NEAR-SURFACE SHEAR WAVE VELOCITY AND ATTENUATION AT GARNER VALLEY



中国地质大学
CHINA UNIVERSITY OF GEOSCIENCES

Lei Qin^{1,2}, Jamison H. Steidl³, Kenneth S. Hudson⁴, Timothy Lamere⁵, Hongrui Qiu^{6,7}, Yehuda Ben-Zion⁸

¹Department of Earth Sciences, University of Southern California, Los Angeles, CA, US

²Hubei Subsurface Multi-Scale Imaging Key Laboratory, Institute of Geophysics and Geomatics, China University of Geosciences, Wuhan, China

³Earth Science Institute, University of California, Santa Barbara, CA, US

⁴Civil and Environmental Engineering, University of California, Los Angeles, CA, US

⁵Manufacturing Design Engineer, Apple Inc.

⁶Department of Earth, Environmental and Planetary Sciences, Rice University, Houston, TX, USA

⁷Earth, Atmospheric and Planetary Sciences, Massachusetts Institute of Technology, Cambridge, MA, USA

⁸Southern California Earthquake Center, University of Southern California, Los Angeles, CA, USA

1. Key points

- **Average in the top 6 m: Impulse response functions (IRFs) at 6-22 m; Vs ~200 m/s, with ~10% seasonal variation**
- **At 5 m depth: Cross-hole array experiment; Vs ~ 210-227 m/s, with ~6% seasonal variation; frequency-dependent Q values are less than 10.**
- **Vs decreases with the precipitation (i.e. water table) and temperature with a phase delay of ~64 days.**
- **S-wave velocity variations decrease with depth**
- **Possible structure variation induced by thermoelastic strains in the top ~1 m**

2. Data - Garner Valley Downhole Array (GVDA)

Figure 1. Event and Array information

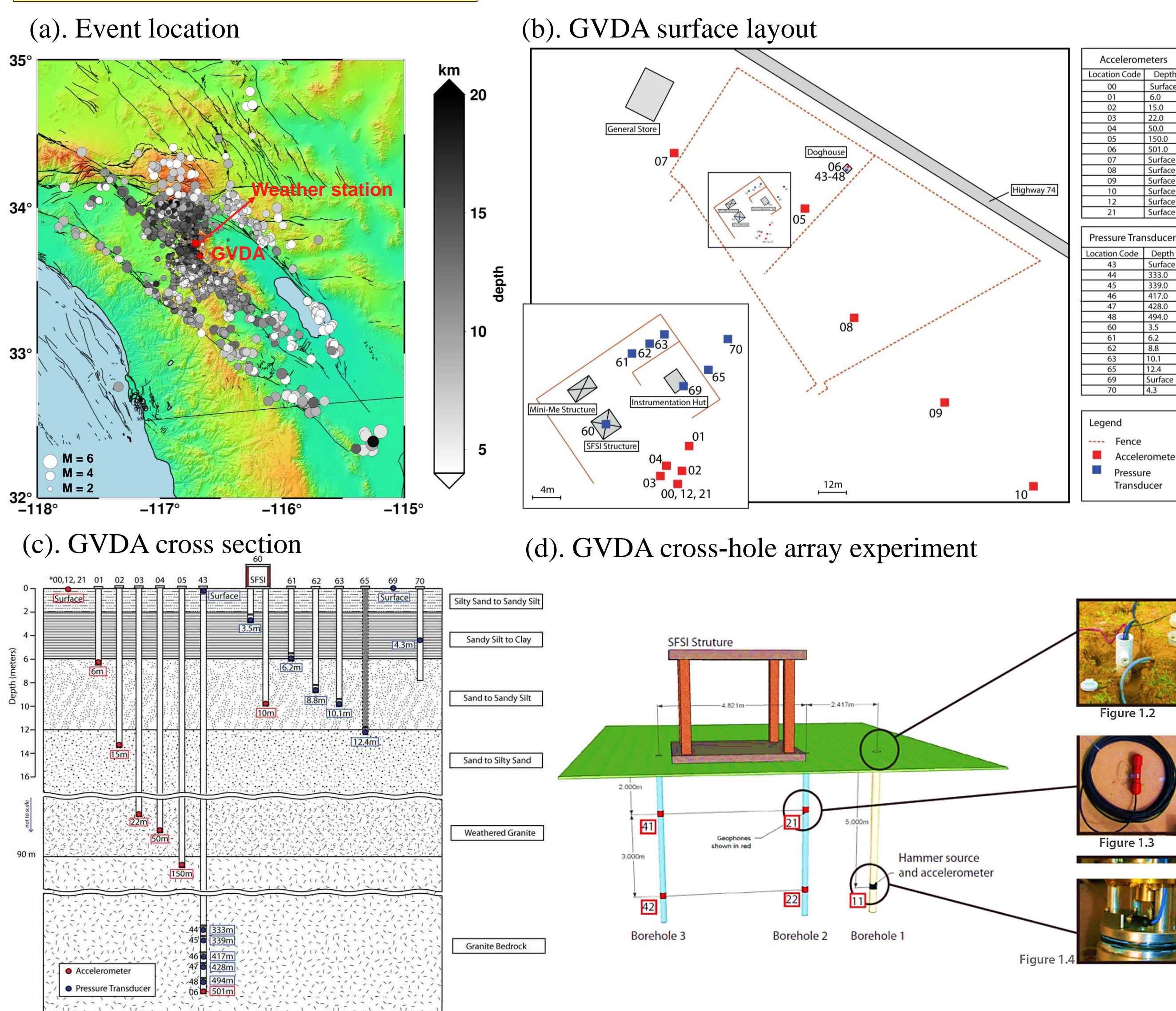


Figure 1. (a). Events (circles) with high-quality (SNR>2) S-wave arrivals at GVDA (red triangle) from 2005 to 2021. The red square shows the location of the Idyllwild weather station that is used in this study. (b) Surface layout of the GVDA, with accelerometers (red squares) and pore pressure transducers (blue squares). (c) Conceptual cross-section placement of the GVDA, showing seismic stations (red dots) at 0-501 m and pore pressure sensors (blue dots) at 0-12.4 m. (d). Cross-hole array experiment at GVDA, with a solenoid powered, dual directional hammer located at 5 m in depth and four geophones, two located at 5 m (22, 42) and the other two (21, 41) directly above them at 2 m in depth. An accelerometer (11) is also mounted to the hammer strike plate. The sampling rate of the geophones is 2000 samples per second. (b)(c)(d) are available from <http://nees.usc.edu/facilities/GVDA>.

Figure 2. Cross-hole array experiment data

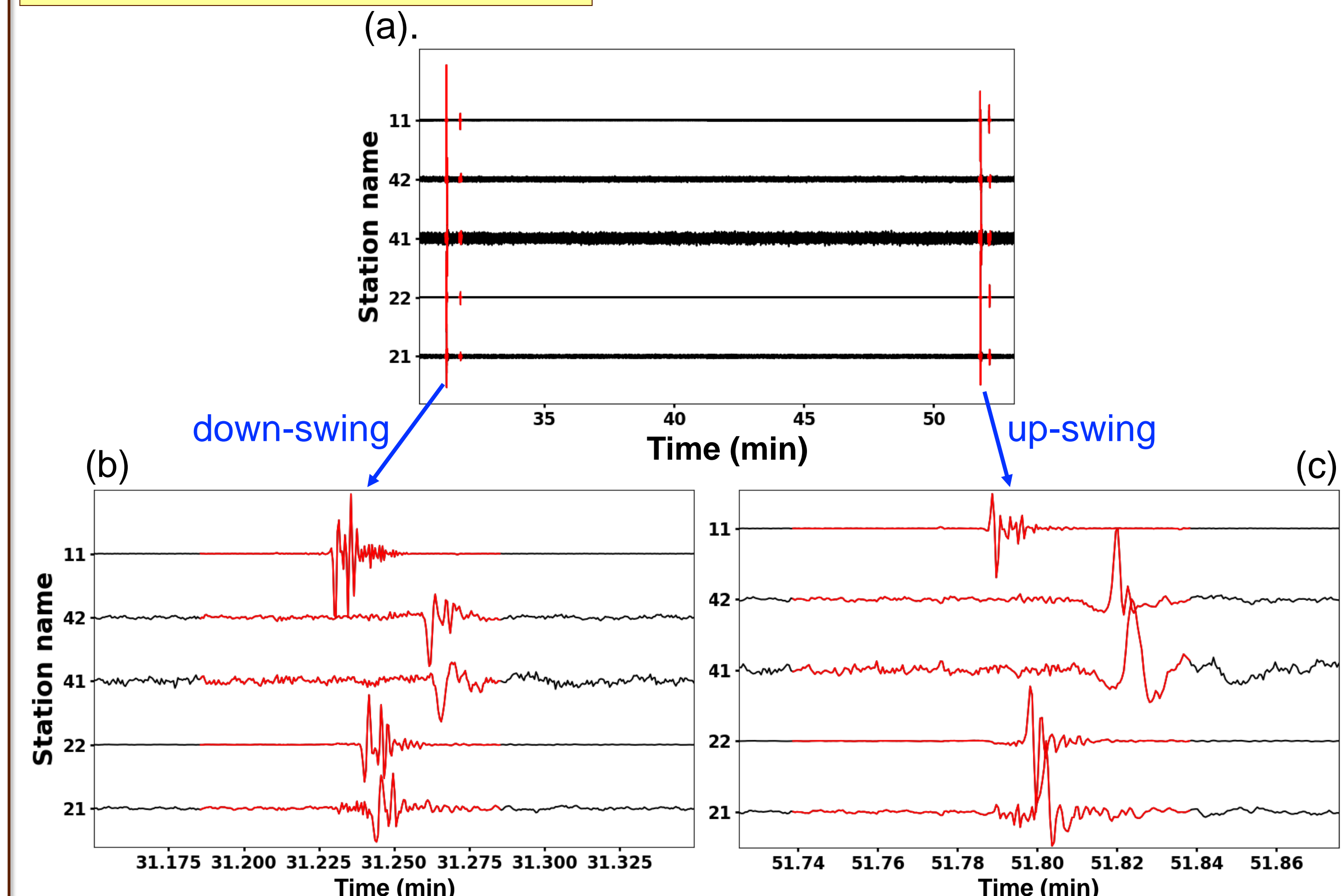


Figure 2. Example waveforms from one cross-hole array experiment, with down- (b) and up- swings (c) of the hammer. Waveforms in red (6 sec centered at the location of the maximum amplitude in station 11) are later analyzed for the S-wave delay time and frequency-dependent Q values between stations. For illustration purpose, the amplitude information is not preserved in the figure.

3. Impulse response functions (IRFs) analysis

Figure 3. IRF examples

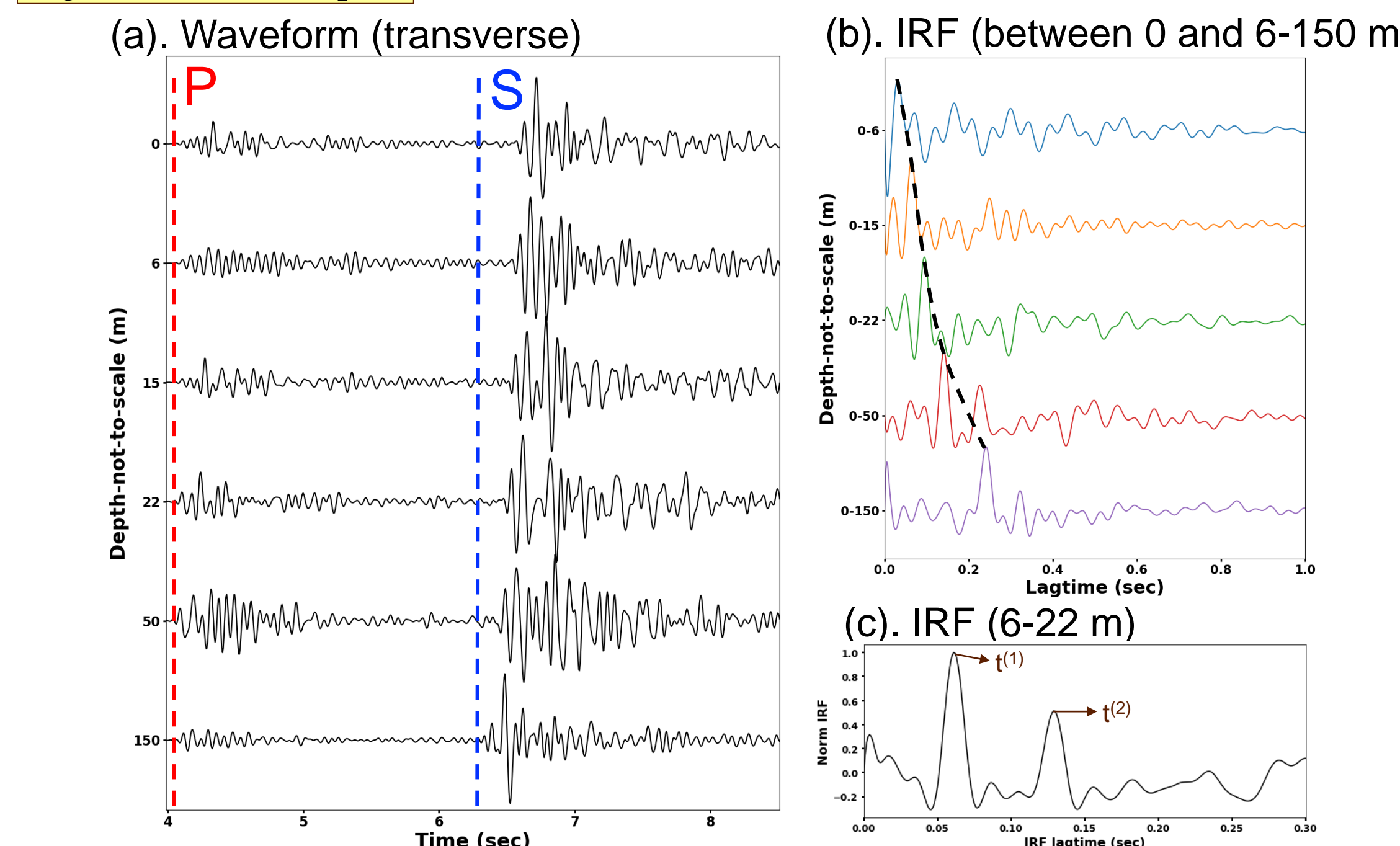


Figure 3. (a) Transverse-component waveforms. (b) IRFs between the surface station and different borehole stations, calculated in the time window 1 sec before and 5 sec after the S-wave arrival. Black dashed line connects the primary peak representing the travel time from surface to borehole stations. (c) IRF between the 6 m (01) and 22 m (03) borehole stations. The primary peak at $t^{(1)}$ (~0.065 sec) represents SH-wave travel time from 22 m to 6 m, and the secondary peak at $t^{(2)}$ (~0.125 sec) shows the travel time from 22 m to the 6 m borehole which is reflected at the free surface.

Figure 5. dt/t vs. precipitation

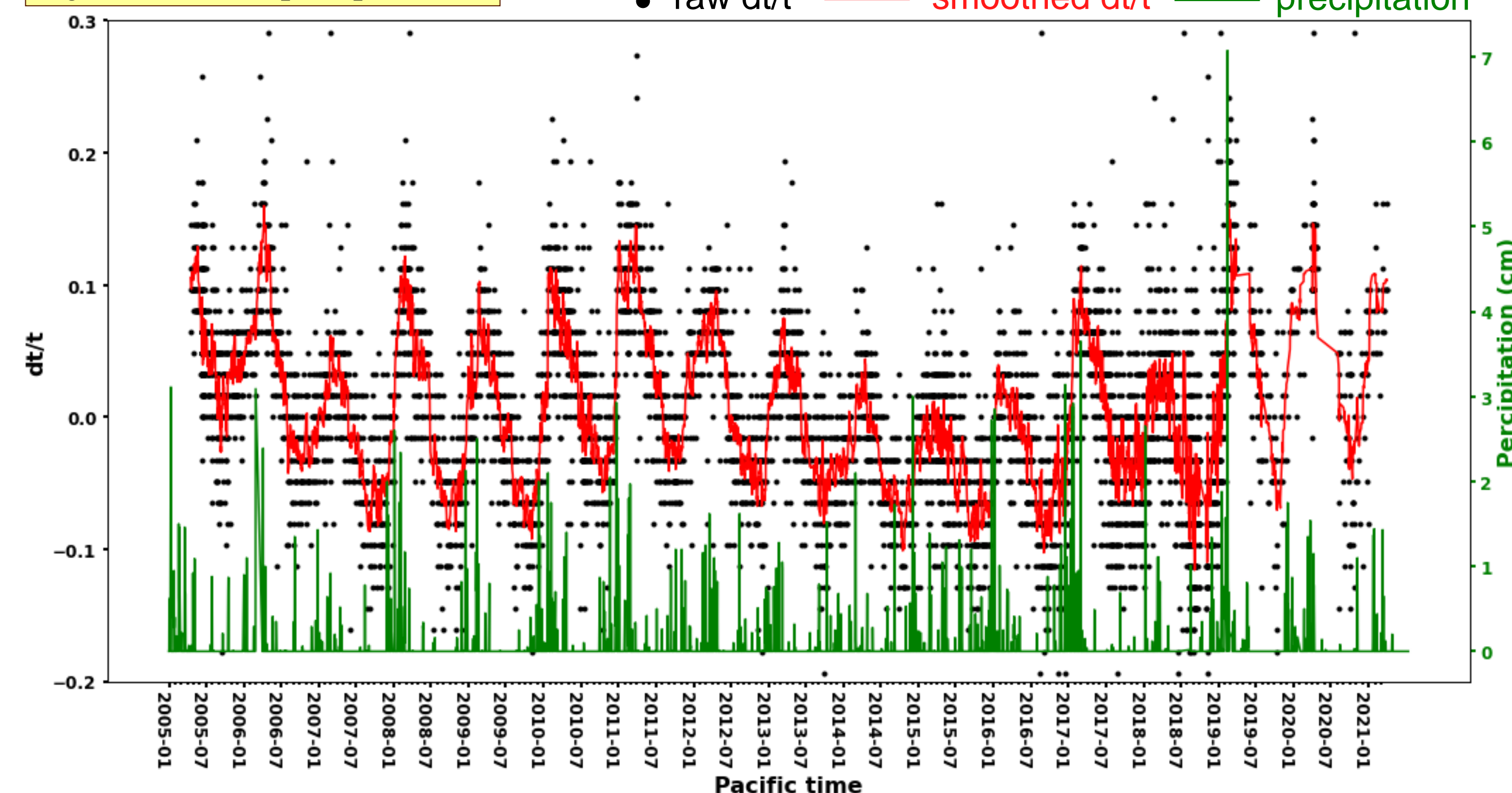


Figure 5. Comparison between the travel time variation (dt/t; black dots and red curve) and the precipitation data (green curve) from the Idyllwild weather station. Here $dt/t = (t - t_0)/t_0$, where t_0 is the median value of all the t measurements (black dots). The red curve is the 11-point smoothed dt/t.

4. Cross-hole array experiment

Figure 7. Q value estimation

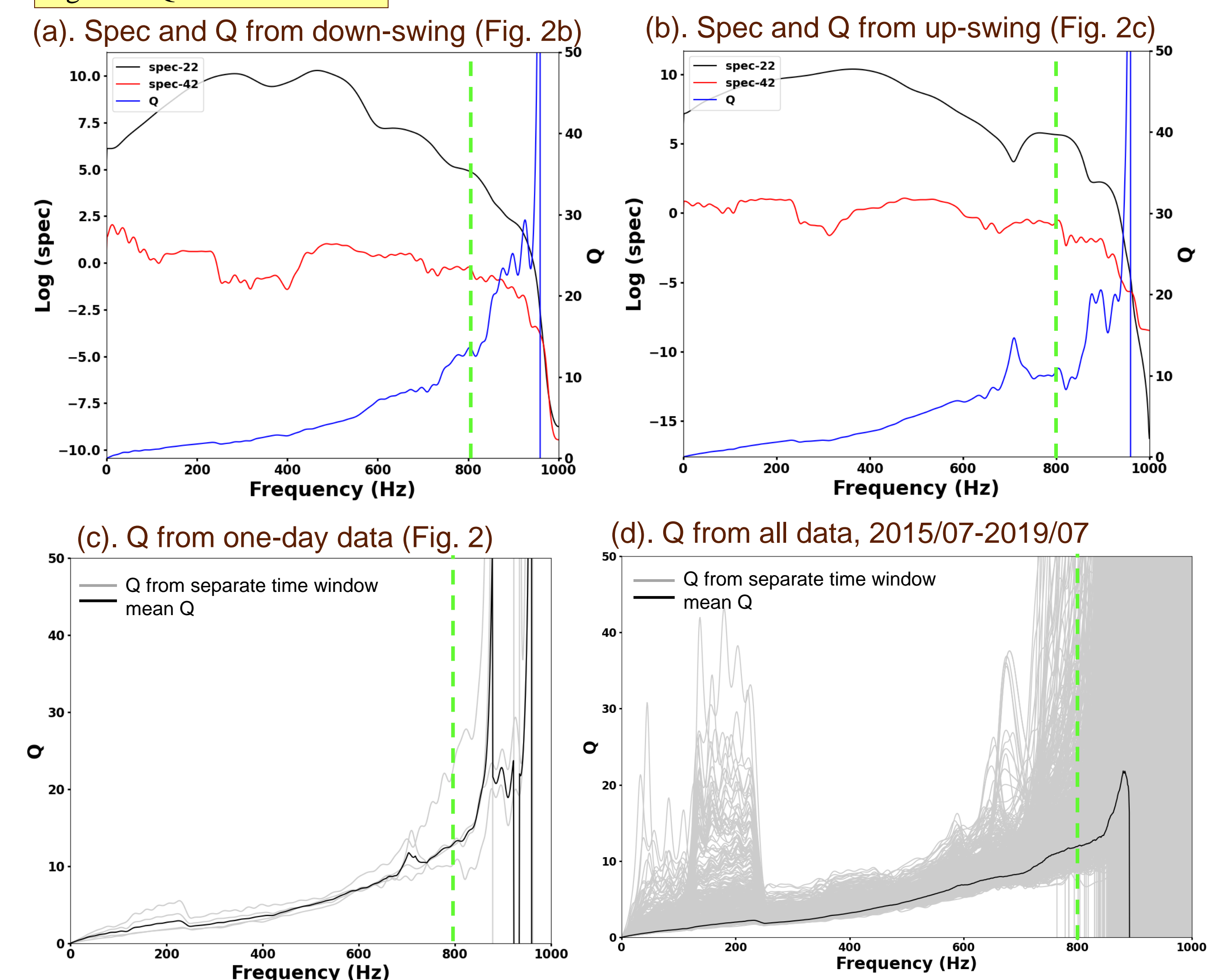


Figure 7. (a) Power spectral density of down-swing data (Fig. 2b) at station 22 (black curve) and 42 (red curve), and the corresponding frequency-dependent Q values (blue curve). The results are reliable up to 800 Hz (green dashed line). (b) same as (a), but for up-swing data (Fig. 2c). (c) Frequency-dependent Q values (gray curves) from the four time windows in one-day data (in red in Fig. 2a), and the corresponding median value (black curve). (d) All Q values from cross-hole array experiment data (gray curve) and the corresponding median values (black curve).

Figure 4. IRF results 1-30 Hz IRF; 01 (6 m)-03 (22 m); transverse component

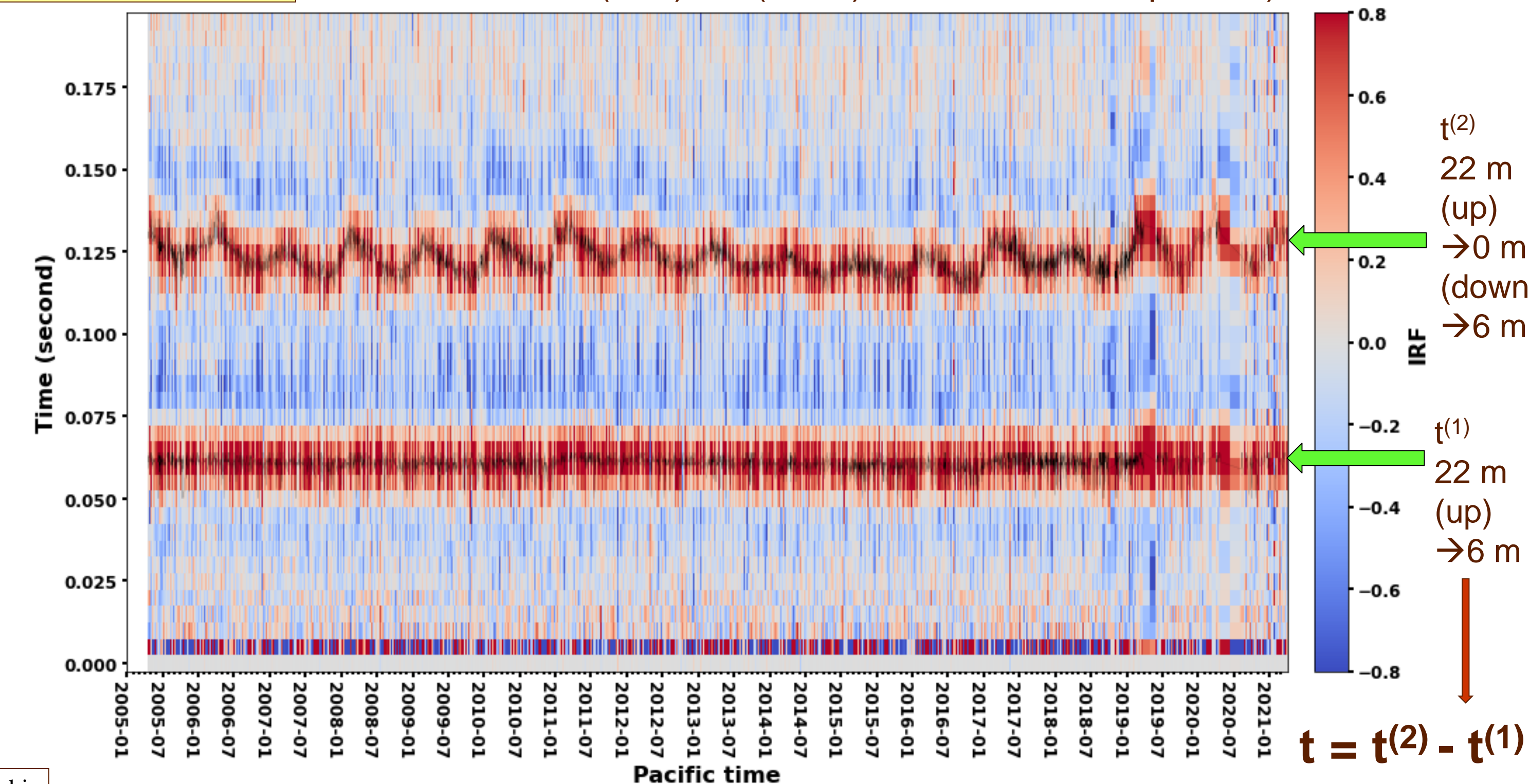


Figure 4. SH-wave IRFs between the 6 m (01) and 22 m (03) borehole stations (similar to Fig. 3c). Waveforms are band pass filtered at 1-30 Hz and each IRF is normalized by the maximum amplitude. Horizontal axis shows the event occurrence time and vertical axis is the IRF lagtime. The change of $t = t^{(2)} - t^{(1)}$ represents the temporal change of the material in the top 6 m.

Figure 6. dt/t vs. temperature

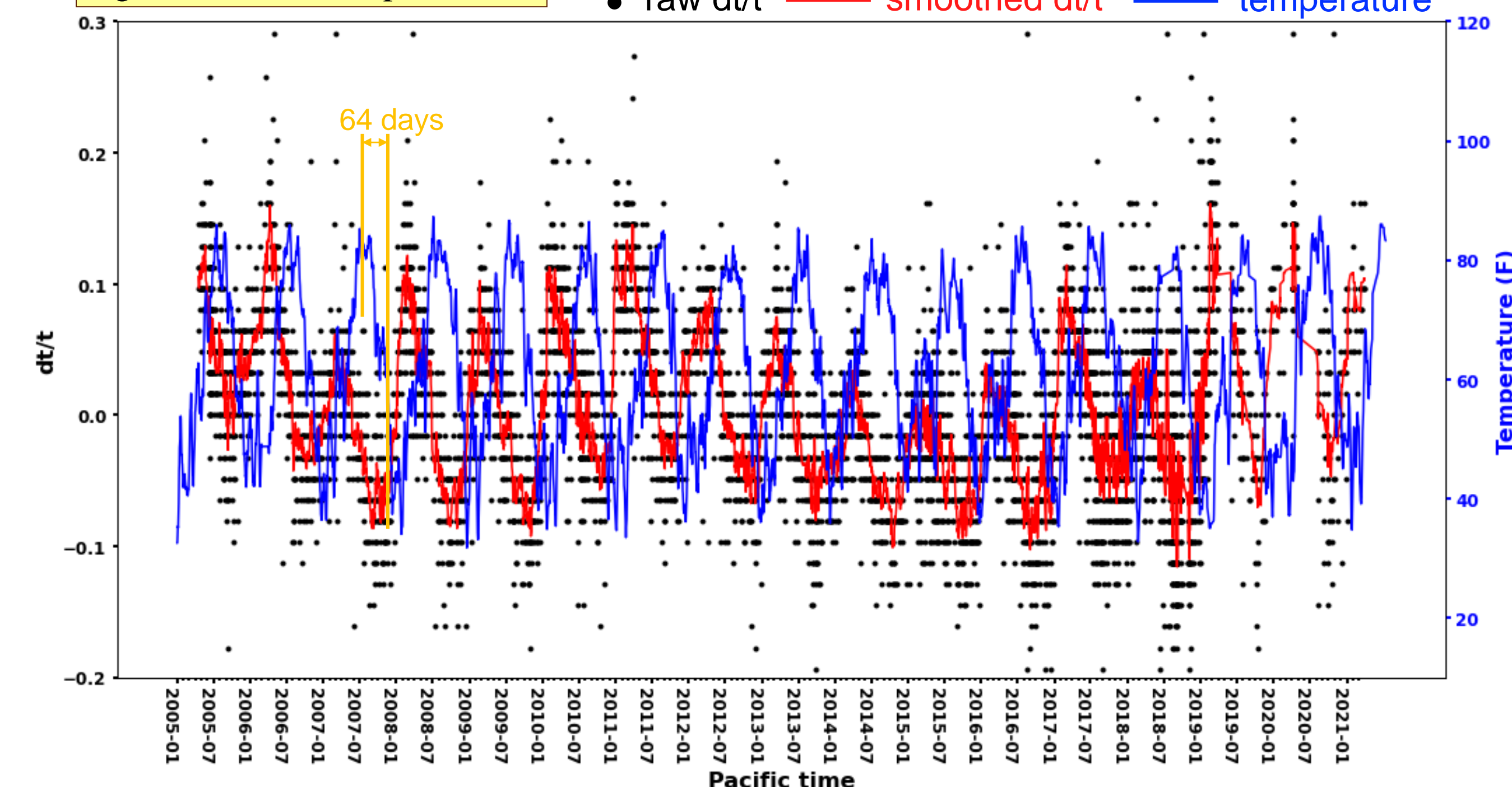


Figure 6. Comparison between the travel time variation (dt/t) and the temperature data (blue curve) from the Idyllwild weather station. Here the black dots and red curve are the same as in Fig. 5. The apparent phase delay of -dt/t (i.e. dv/v) curve relative to the temperature is ~64 days.

Figure 8. S-wave travel times between stations 42 and 22 obtained via cross-correlation

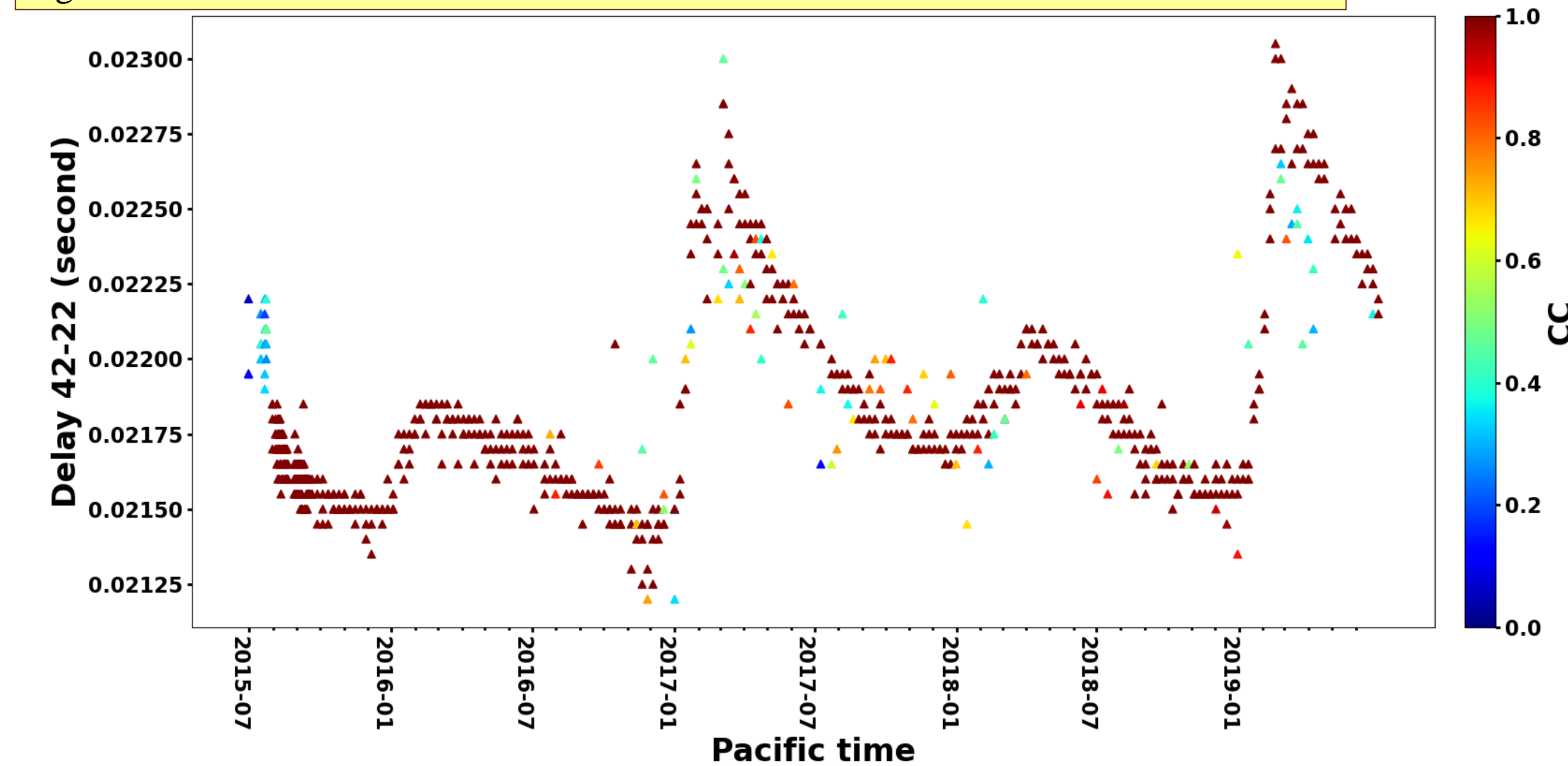


Figure 9. Comparison of dt/t in the top 6 m (red curve) and at the 5 m depth (black triangles).

