

# Toward a SCEC Community Rheology Model: Technical Activity Group Kickoff and Workshop

CRM ingredients

Geologic Framework

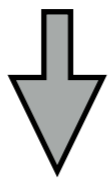
Michael Oskin, UC Davis

Geotherm (CTM)

Wayne Thatcher, USGS

Flow laws and  
parameters

Greg Hirth, Brown  
Whitney Behr, UT Austin



Preliminary  
Mojave CRM

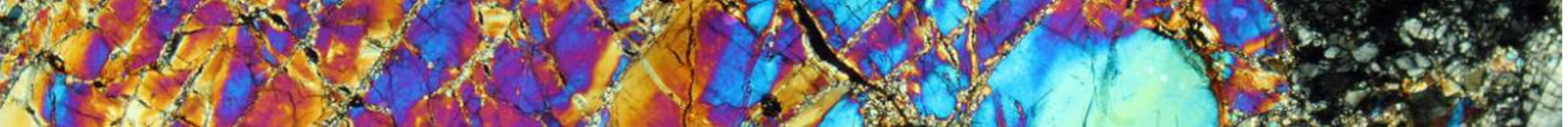
Liz Hearn, Capstone  
Geophysics



## SATURDAY, SEPTEMBER 9, 2017 Morning: CRM description and accomplishments to date, plus speakers and discussions

09:00 - 09:15	Welcome, Motivation, Workshop Outline and Goals	<i>Elizabeth Hearn</i>
	<u>Geologic Framework and Community Thermal Model (CTM)</u>	<i>Michael Oskin, Wayne Thatcher</i>
09:15 - 09:30	Overview of the geologic framework, with emphasis on constraining the lithospheric architecture of the Mojave region	<i>Michael Oskin</i>
09:30 - 09:45	<del>Tectonic modification of the Mojave Desert region during laramide shallow-angle subduction</del>	<del><i>Alan Chapman</i></del>
09:45 - 10:00	Applying the depth distribution of seismicity to probe the rheology of the seismogenic crust in southern California	<i>Egill Hauksson</i>
10:00 - 10:15	Discussion: How should geologic and tectonic inferences of lithosphere composition guide the CRM?	<i>All</i>
10:15 - 10:30	Preliminary CTM, with an emphasis on the Mojave region	<i>Wayne Thatcher</i>
10:30 - 10:45	Discussion: How important are thermal versus compositional differences?	<i>All</i>
10:45 - 11:00	<i>Break</i>	
	<u>Inferring Rheology, Preliminary Mojave CRM</u>	<i>Greg Hirth, Whitney Behr</i>
11:00 - 11:15	Overview of rheologies for Mojave rock volumes and shear zones, guidance on differential stress, volatile content and other parameters for these flow laws	<i>Whitney Behr</i>
11:15 - 11:30	Inversion of seismic velocity for rheology	<i>William Shinevar</i>
11:30 - 11:45	Discussion: Geologic vs. seismic inference of rheology	<i>All</i>
11:45 - 12:15	Preliminary Mojave CRM effective viscosities and their consistency (or not) with Hector Mine Earthquake postseismic deformation	<i>Elizabeth Hearn</i>
12:15 - 12:30	Discussion: Assessing the CRM with deformation data and models	<i>All</i>
12:30 - 13:30	<i>Lunch</i>	

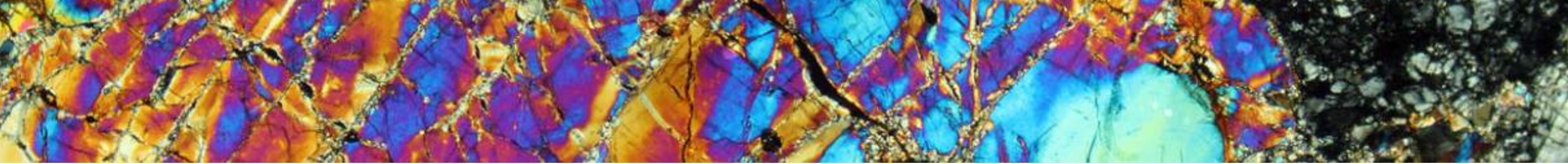




## **SATURDAY, SEPTEMBER 9, 2017 Afternoon: Workshop participants help us plan the future of the CRM**

- 13:30 - 14:30 Disciplinary Breakout Session I *All*  
What can each scientific discipline contribute to the CRM over the next year and by the end of SCEC5, for the Mojave and southern California?
- 14:30 - 14:45 Breakout Session I Summaries *All*
- 14:45 - 15:00 *Break*
- 15:00 - 16:00 Interdisciplinary Breakout Session II *All*  
Identify important scientific questions than can be addressed with a CRM (and other existing resources) by the end of SCEC5 - particularly, problems that relate to the five SCEC5 science questions
- 15:00 - 16:15 Breakout Session II Summaries *All*
- 16:15 - 17:00 Next steps for 2018 and beyond: Our TAG *Elizabeth Hearn, Michael Oskin, Greg Hirth, Whitney Behr, Wayne Thatcher*  
Other GF provinces, shear zones, delivery and sharing, future workshops and closing comments
- 17:00 *Adjourn*

**The bouncer shows up at 5 to kick us out so this room can be prepped for the evening poster session. No lollygagging!**



## SCEC Community Rheology Model: The Mojave

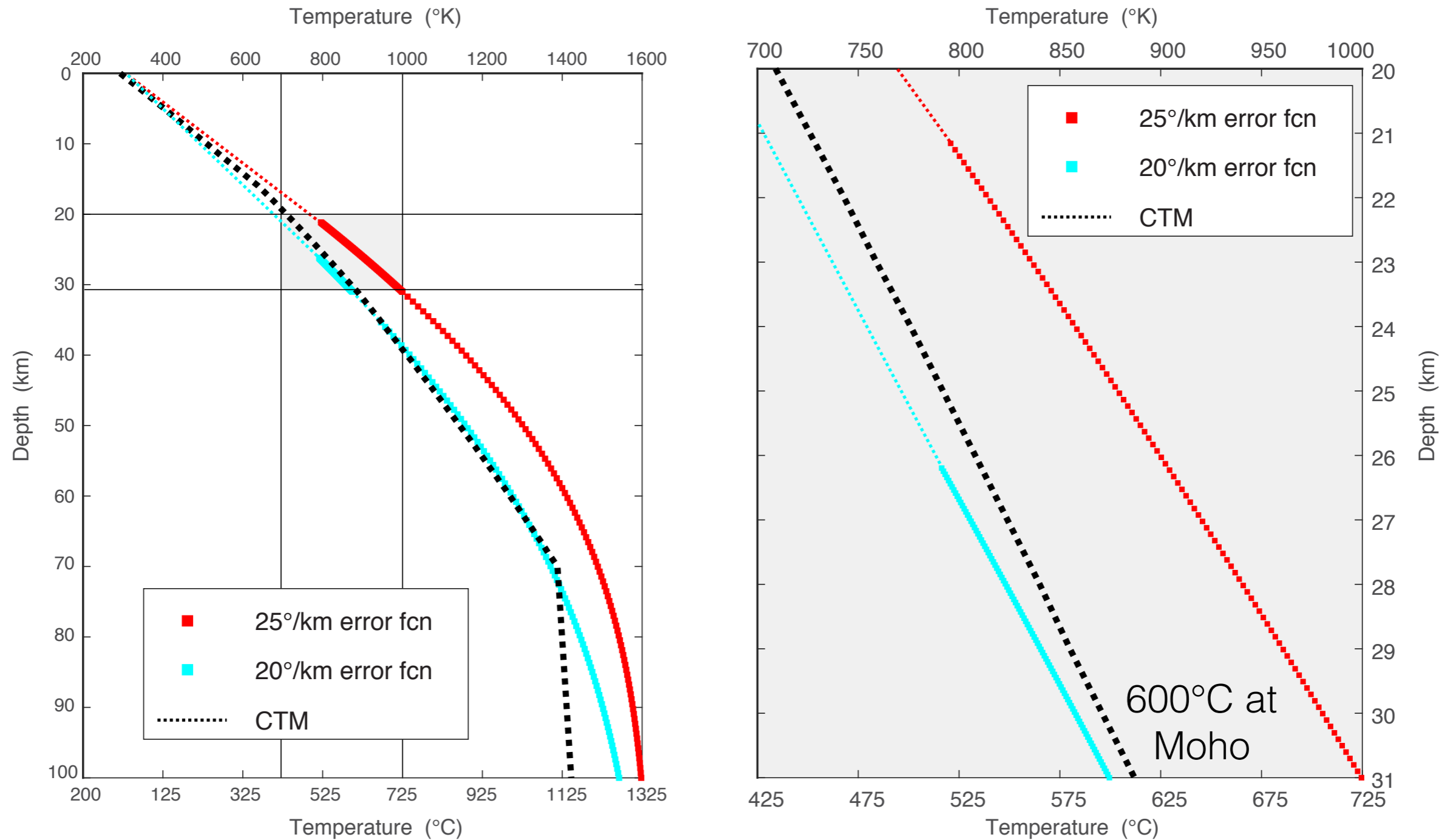
Geotherm (CTM)

Geologic Framework

Flow laws and  
parameters



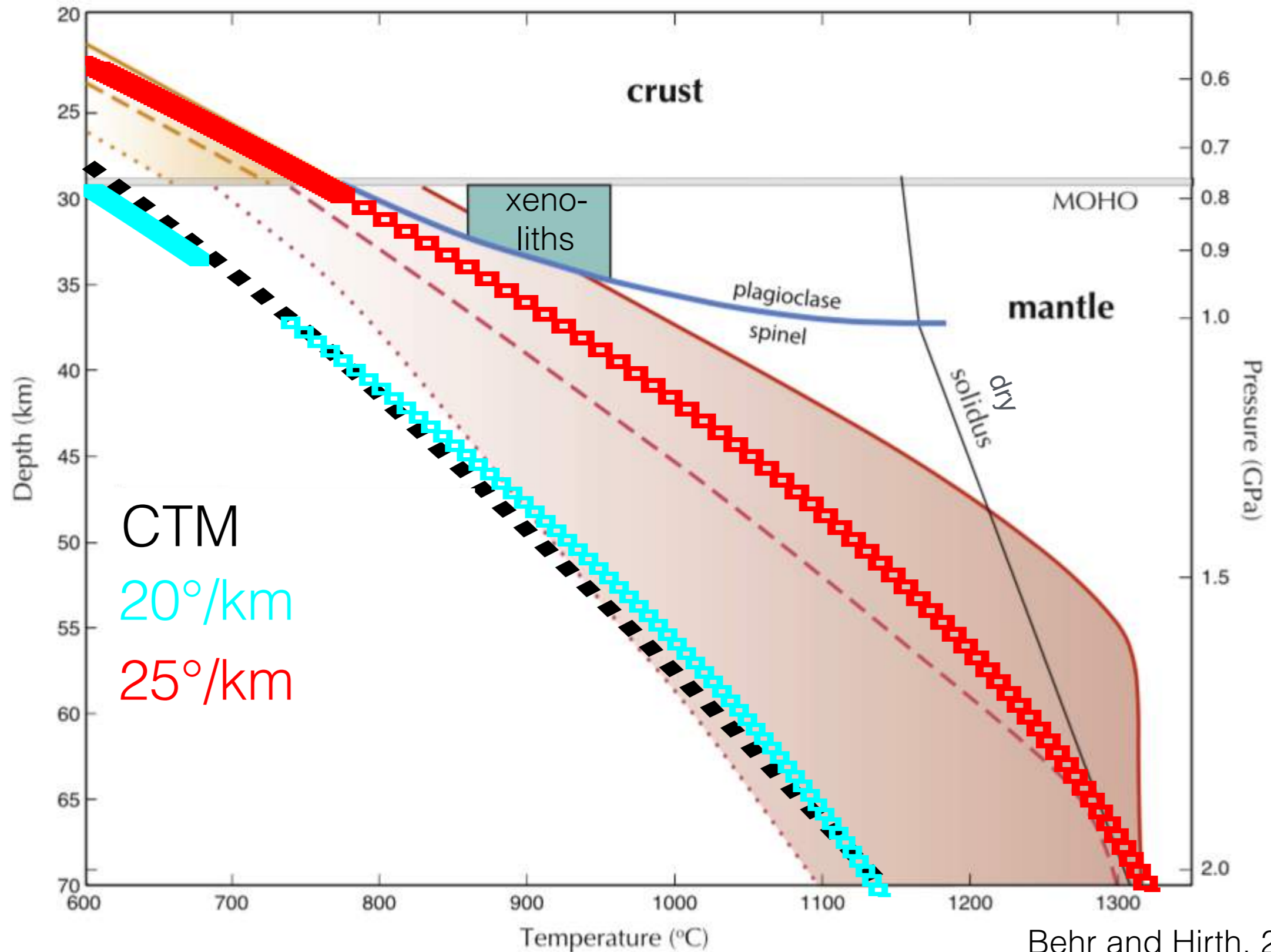
# Mojave CTM: Temperature vs. depth



Mojave CTM has cool lower lithosphere because high heat production is modeled in the felsic crust



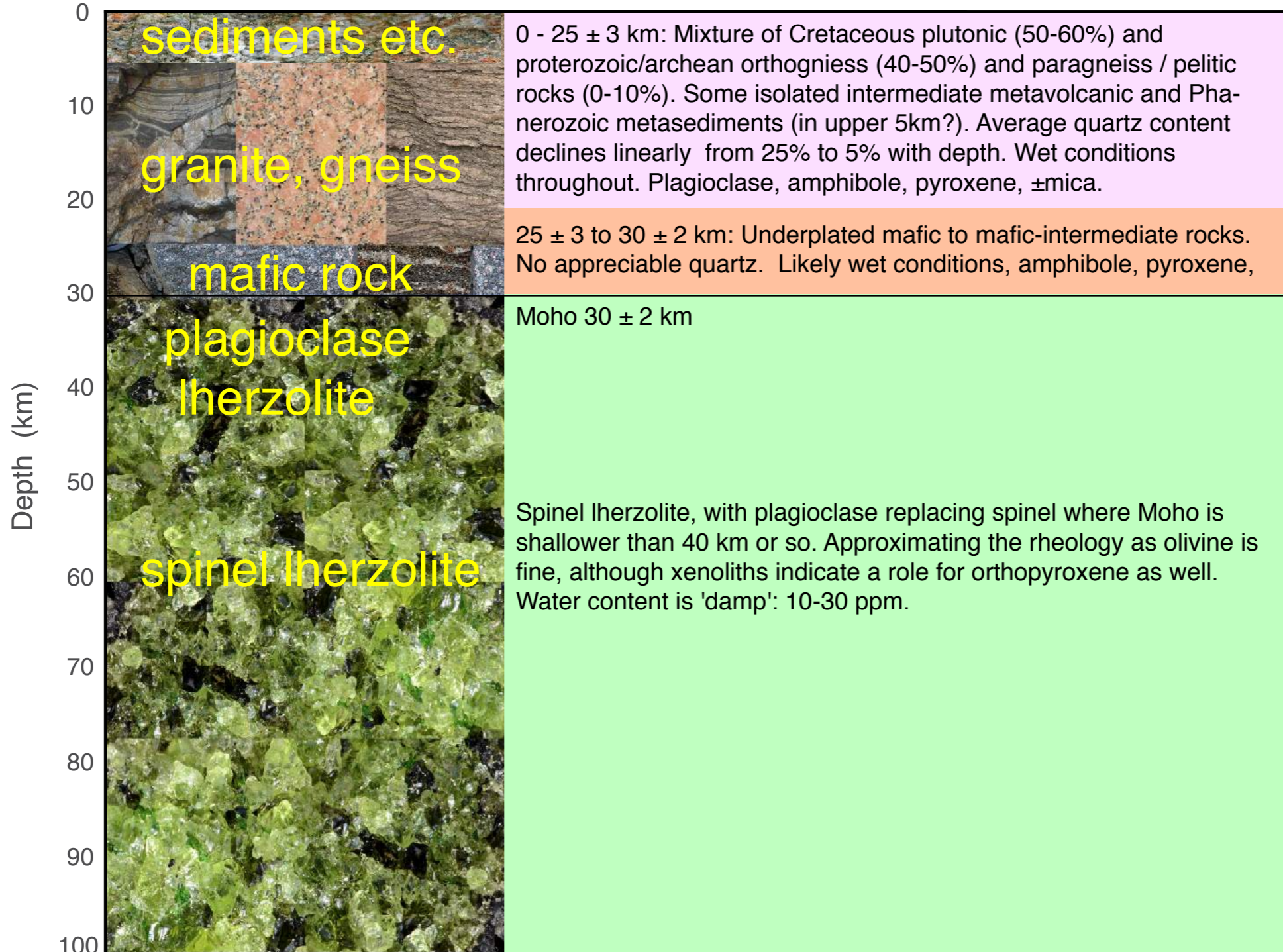
# Mojave CTM: Comparison with other studies



# Mojave Geologic Framework

simplified column

geologist's description



# Dislocation Creep: Power-law flow equation

$$\dot{\epsilon} = A\sigma^n e^{\frac{PV-Q}{RT}} f_{H_2O}^r$$

$$\eta_e = \frac{\sigma}{\dot{\epsilon}}$$

- T from Community Thermal Model
- P from density\*g\*depth
- assume stress or strain rate
- other parameters from flow laws

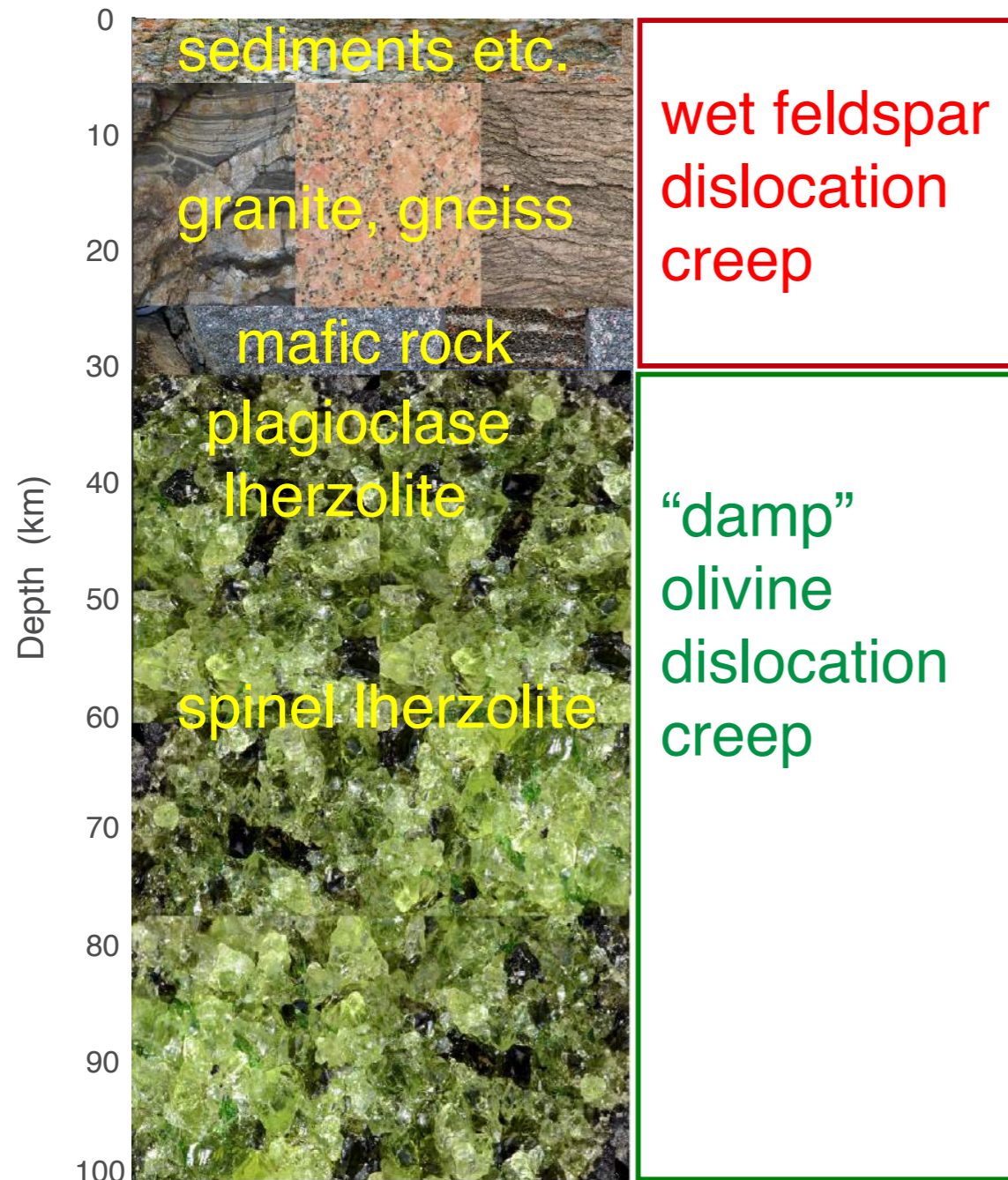


- which fugacity table or equation do I use?
- some versions have COH or H/Si (or ppm?), not water fugacity
- do they mean engineering strain rate? differential stress? torsion vs. uniaxial.
- assume what for stress or strain rate?
- some versions include another fugacity term and/or a grain size term
- some versions include a term for melt.
- what if it's actually diffusion creep?
- what about transient rheology?



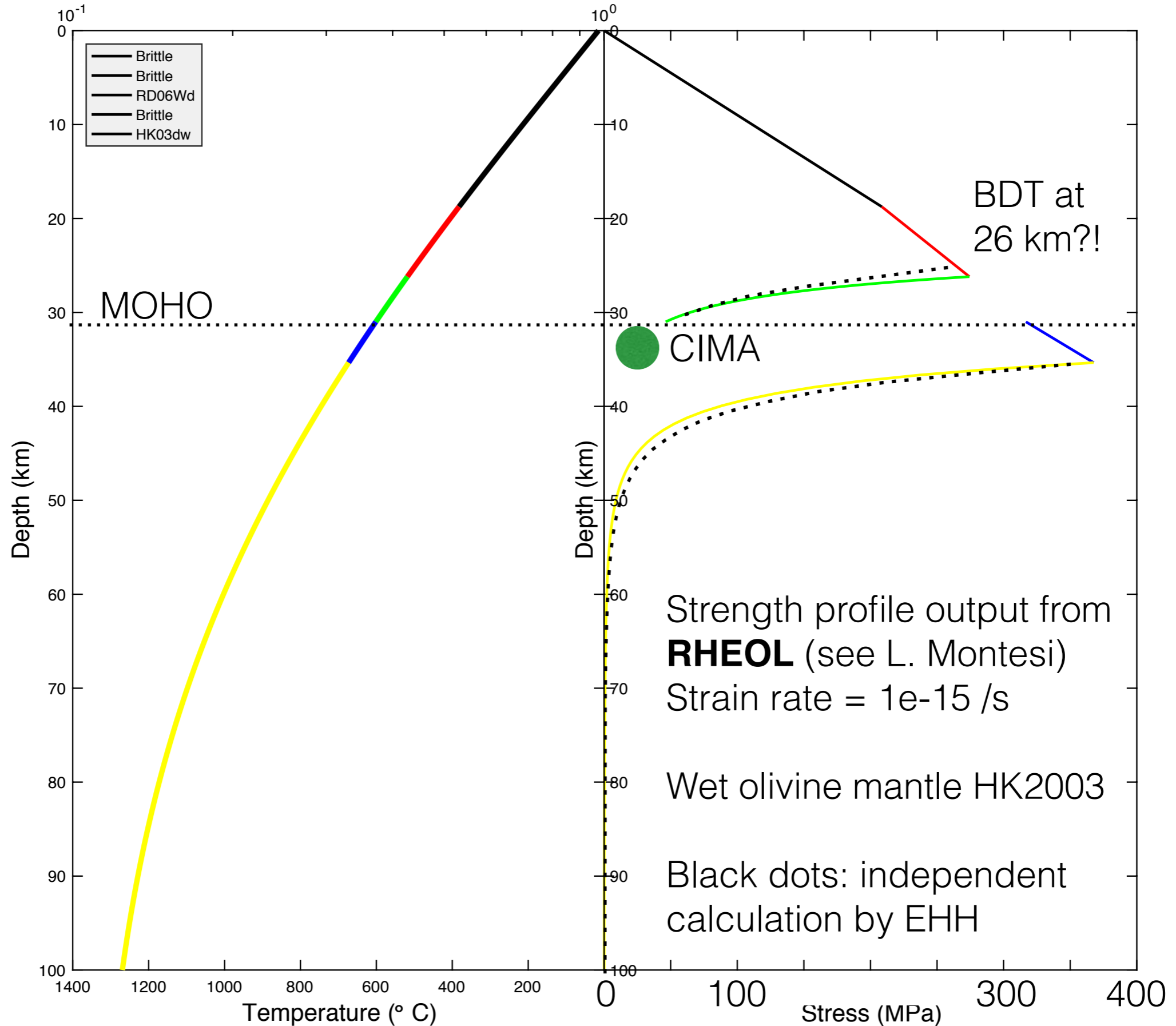
# Power-law flow equation guidance (Mojave)

$$\dot{\epsilon} = A\sigma^n e^{\frac{PV-Q}{RT}} f_{H_2O}^r \quad \eta_e = \frac{\sigma}{\dot{\epsilon}}$$

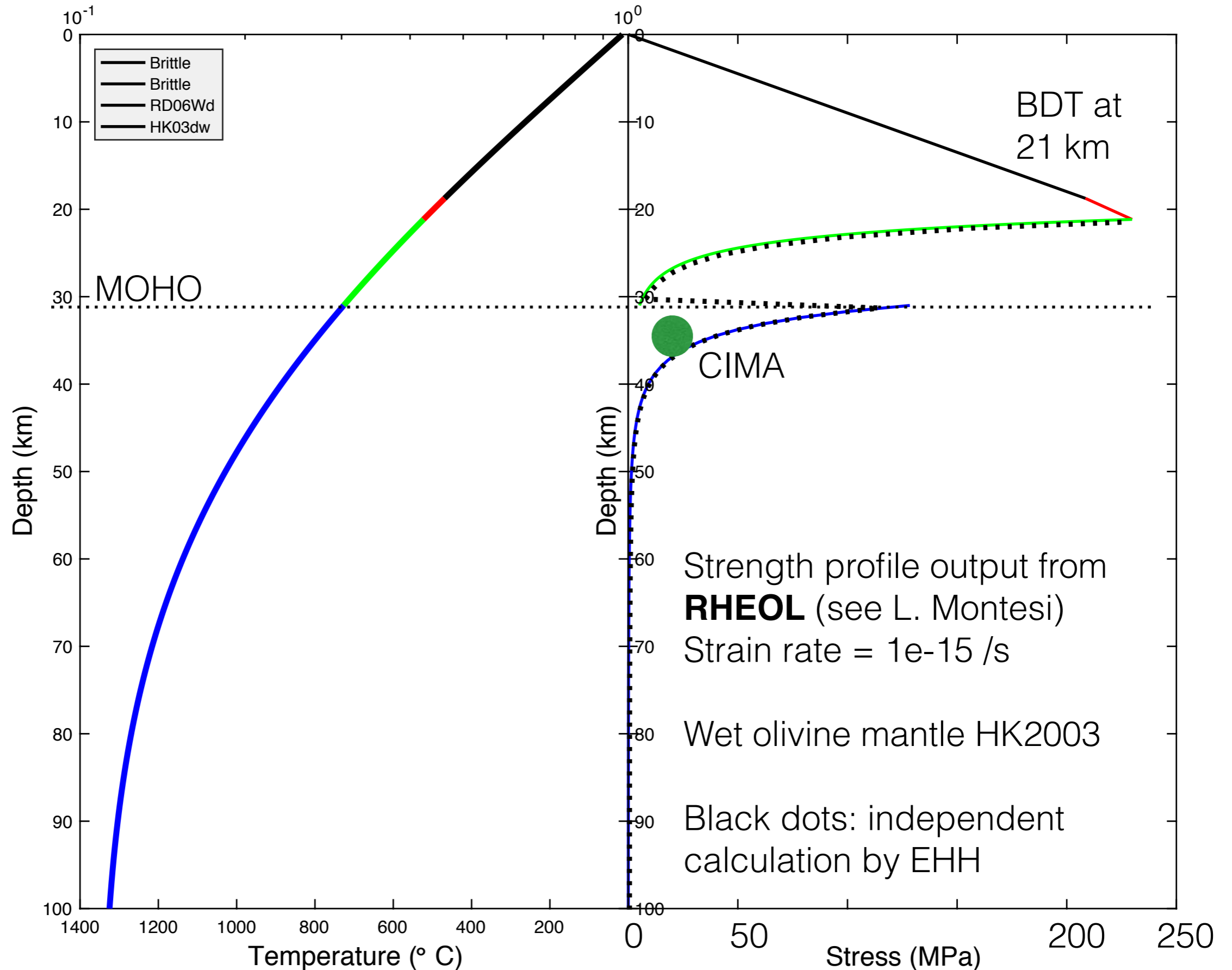


- parameters A, n, r, V and Q from [Rybacki et al. \(2006\)](#) for wet anorthite
- water fugacity consistent with saturation: depends on P and T (Pitzer and Sterner, 1994)
- parameters A, n, r, V and Q from [Hirth and Kohlstedt \(2003\)](#) for wet olivine, *constant COH equation*
- assume COH = 300 (i.e. 300 H atoms per  $10^6$  Si atoms).
- COH substitutes for fugacity in equation

# Mojave strength profile assuming cool geotherm



# Mojave strength profile assuming hot geotherm



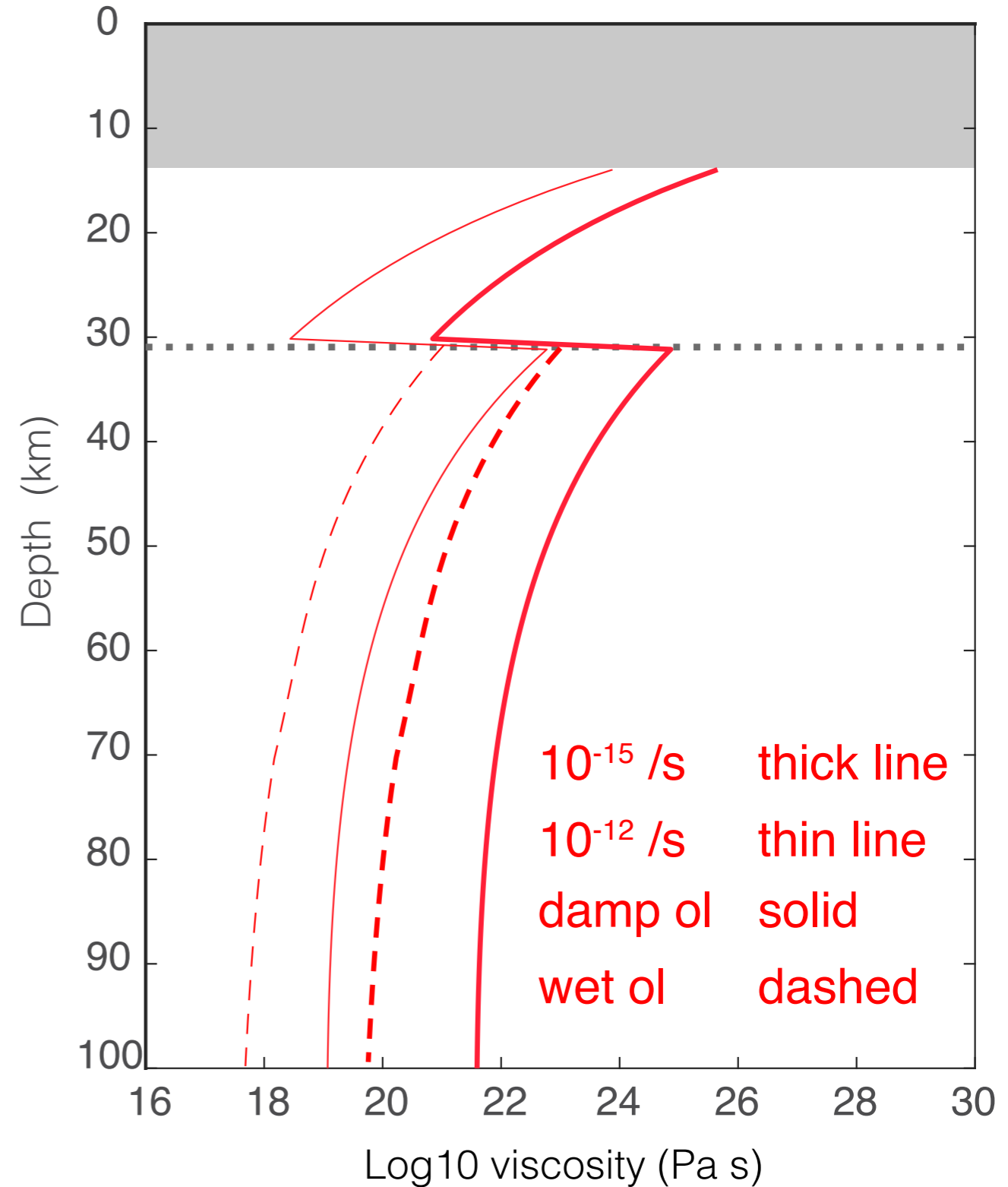
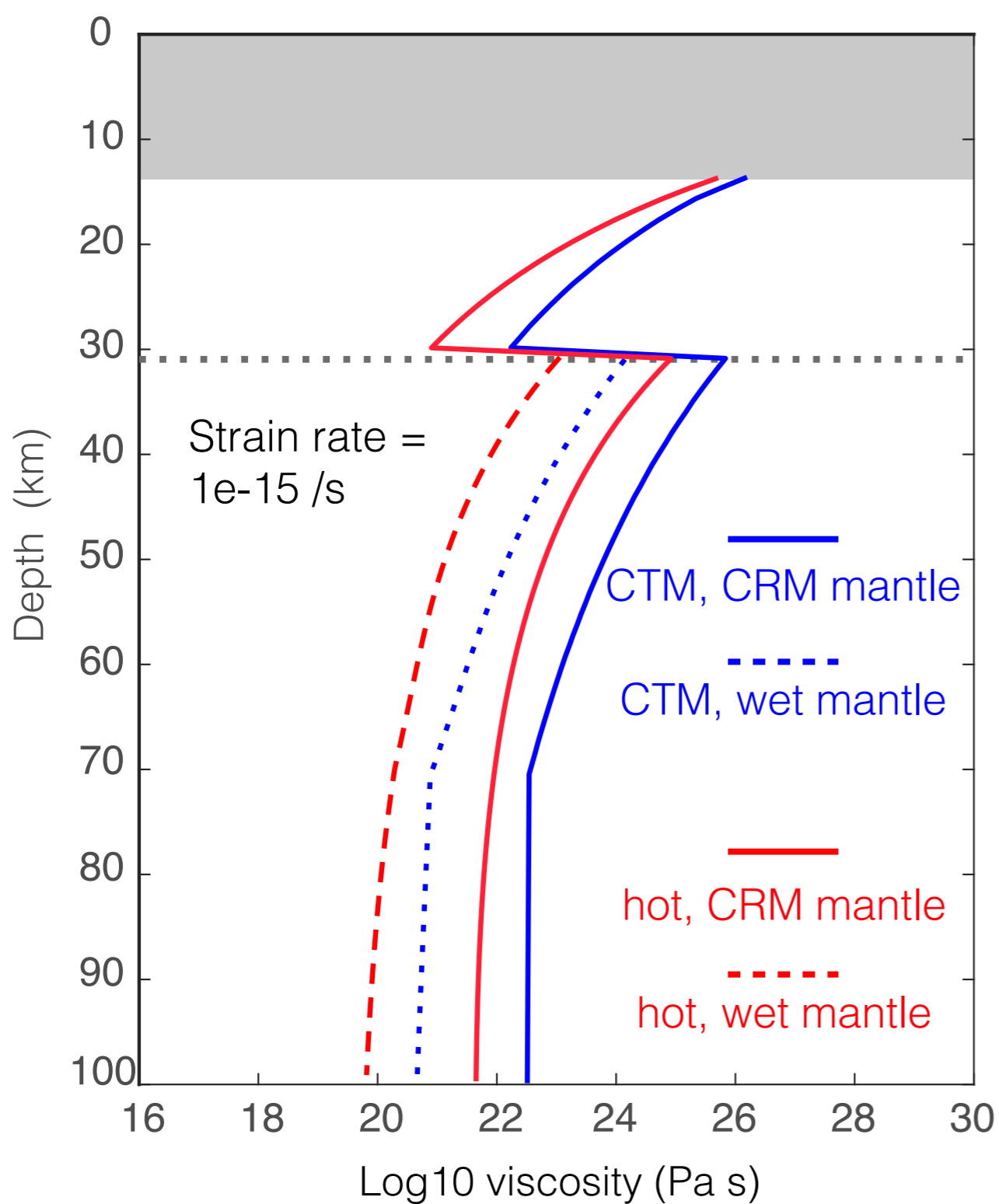


To faithfully represent the CRM I needed to

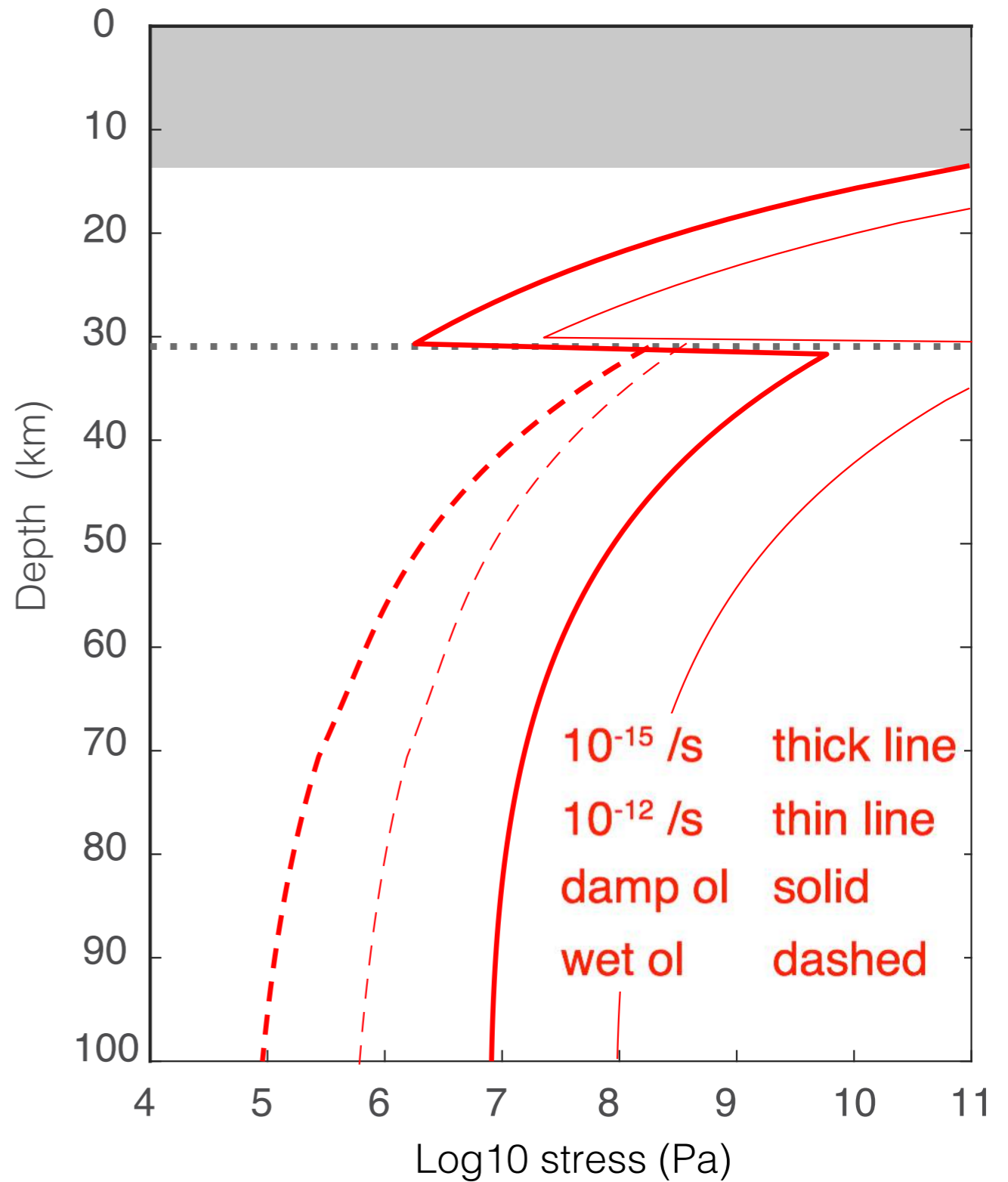
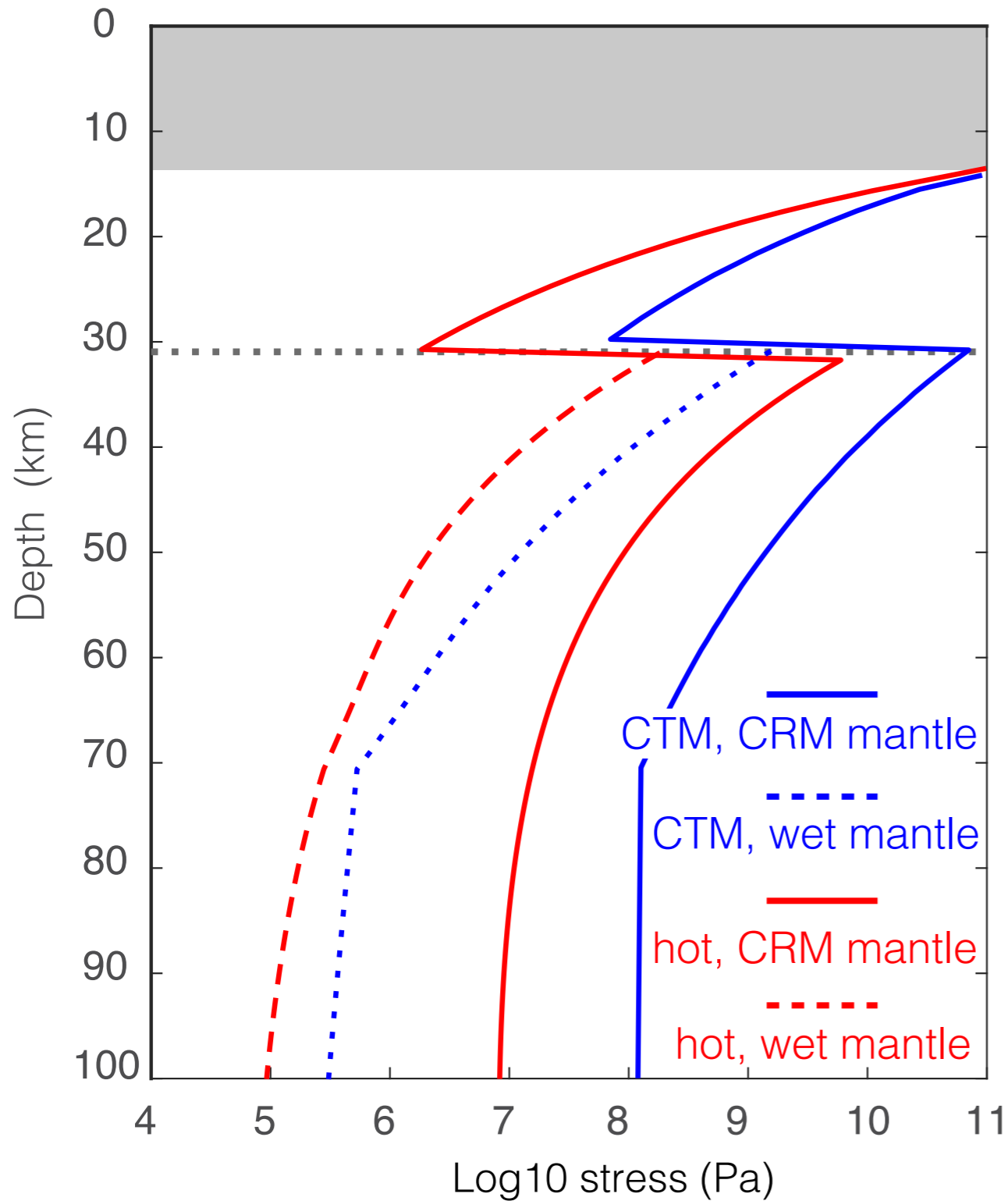
- (1) read in the CTM
- (2) use the “damp” olivine flow law

I calculated effective viscosity and differential stress using a MATLAB code (benchmarked to RHEO calculations)

Higher temperatures, higher water content and higher strain rates: lower effective viscosities

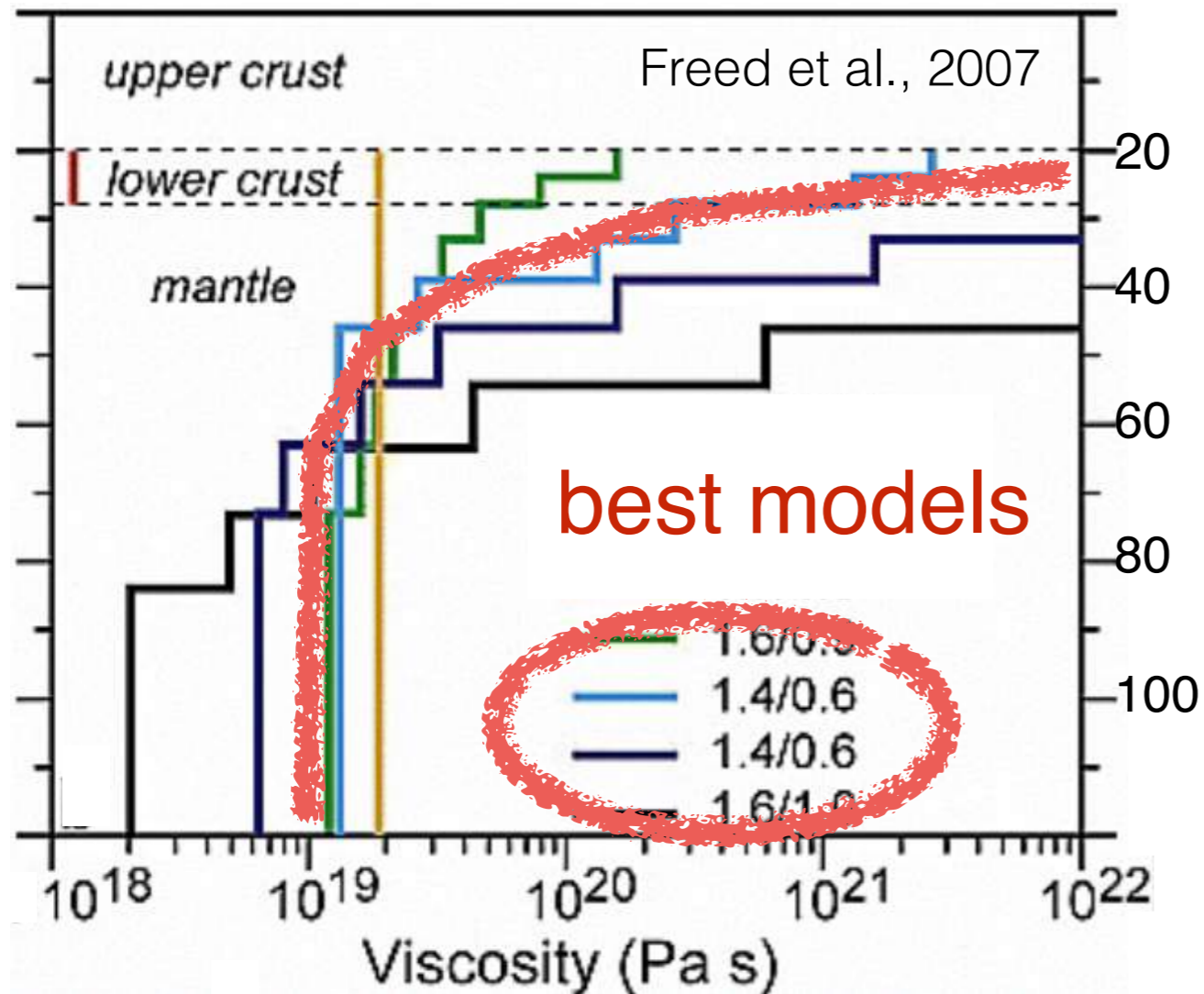


Lower temperatures, lower water content or higher strain rate: higher differential stress



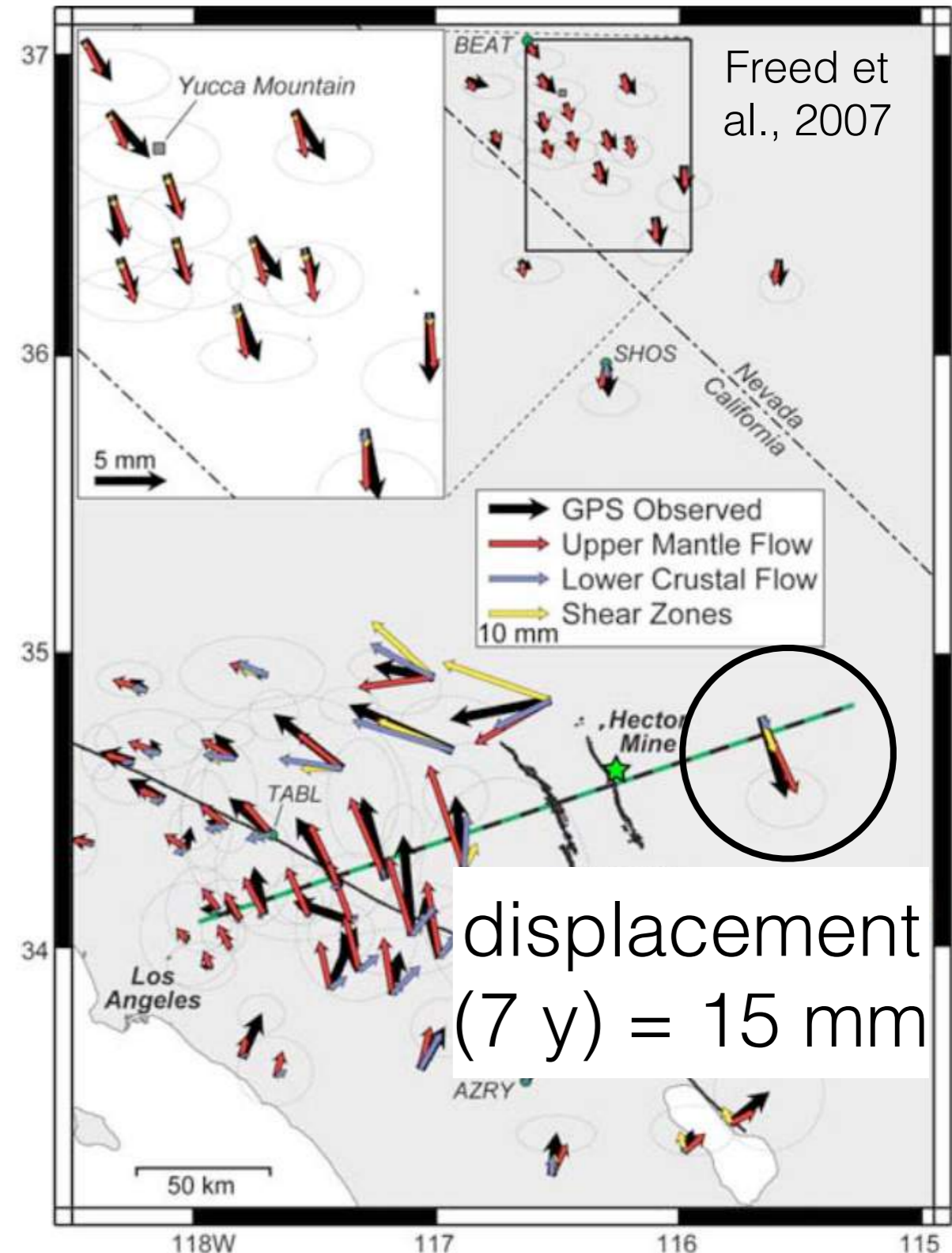


# Hector Mine Earthquake Postseismic Deformation: Inferred viscosity structure, excluding near-field GPS data

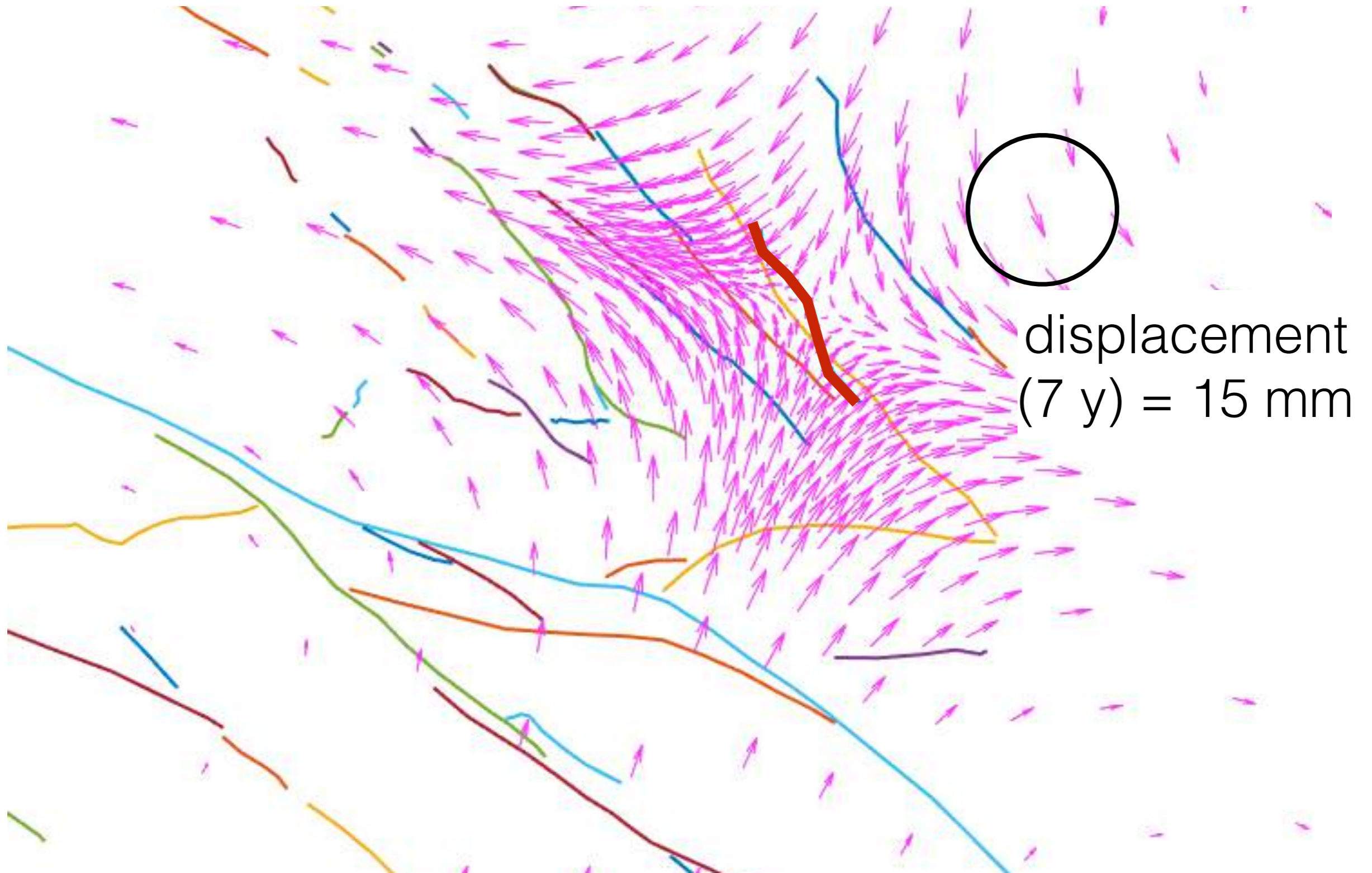


“Creme brûlée” model

Similar results for the Mojave region from post-seismic deformation models, e.g. Pollitz et al., 2000; Pollitz and Thatcher, 2010



# Modeled Hector Mine post-seismic displacements (preferred viscosity structure of Freed et al. 2007)



# CTM geotherm, CRM rheologies

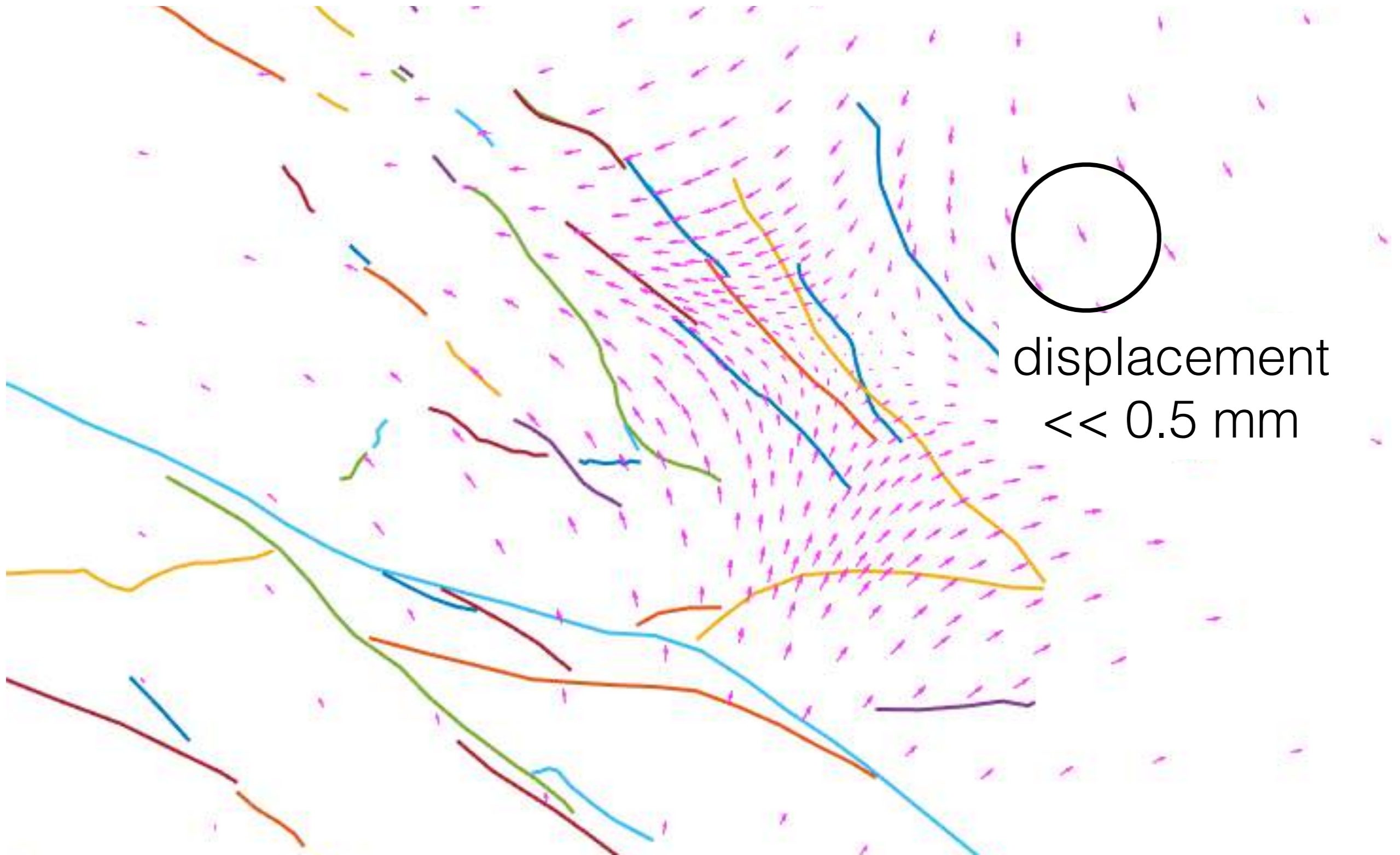
assumes pre-earthquake strain rate of  $10^{-15}/s$

earthquake changes stresses

<b>Depth (km)</b>	<b>eff viscosity_o Pa s</b>	<b>differential stress_o MPa</b>	<b>n</b>
<b>0-15</b>	elastic	large	
<b>15-20</b>	1.5E+25	14,000	3
<b>20-25</b>	1.1E+24	1100	3
<b>25-31</b>	1.5E+23	150	3
<b>31-40</b>	2.5E+25	25000	3.5
<b>40-55</b>	2.2E+24	2200	3.5
<b>55-80</b>	1.3E+23	130	3.5
<b>80-120</b>	1.0E+23	100	3.5
<b>120-200</b>	1E+21	1	3.5



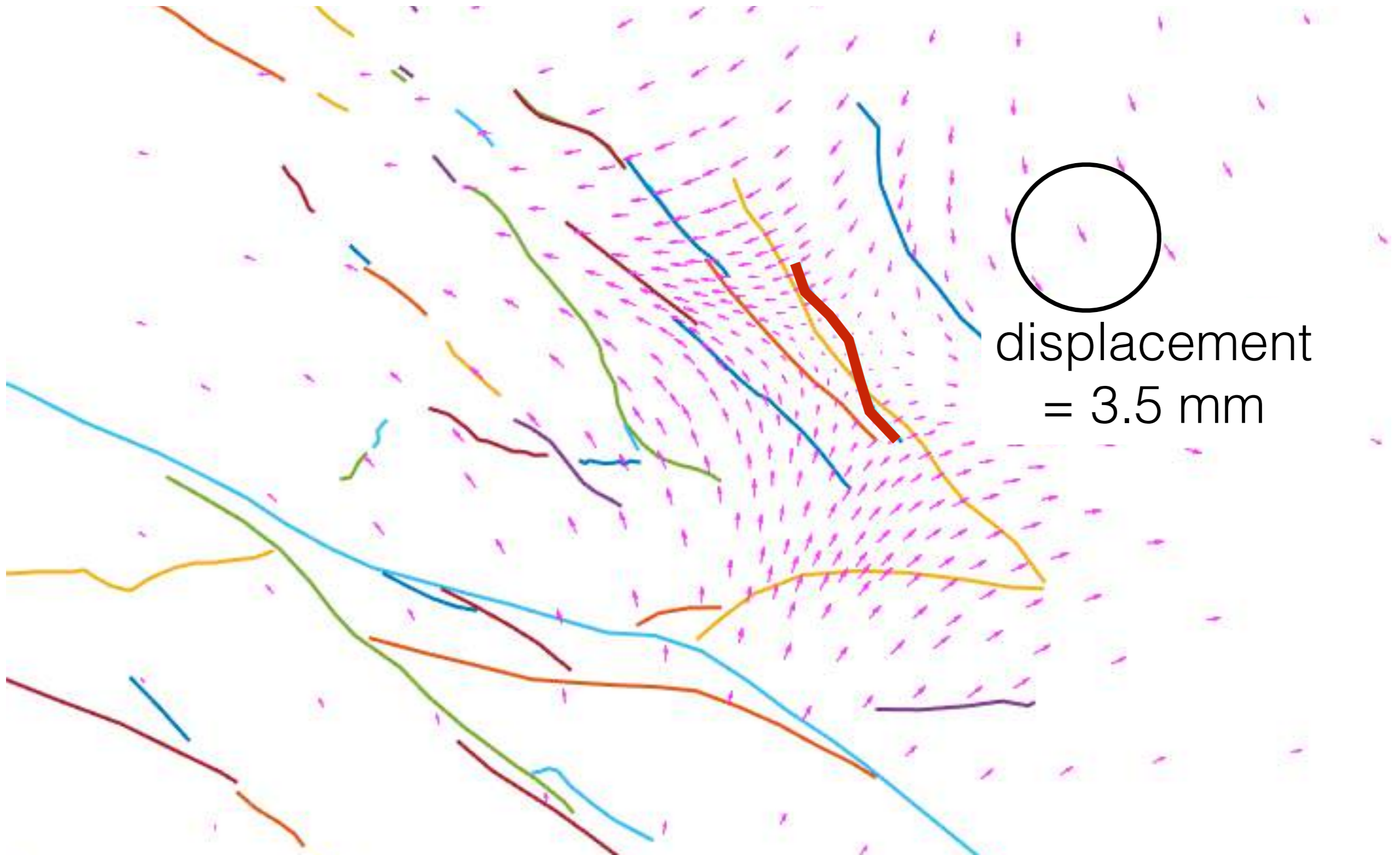
# Modeled Hector Mine post-seismic displacements (Mojave CTM and CRM)



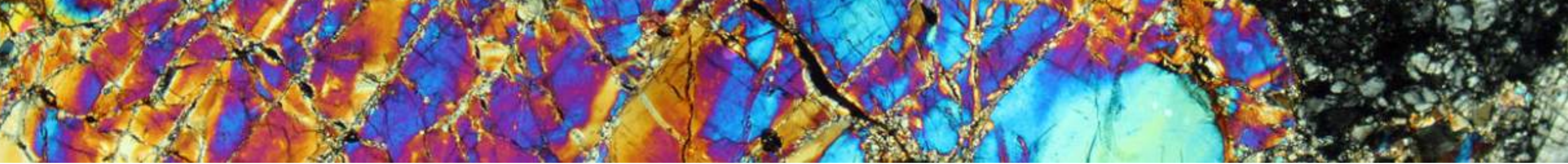
# Modeled power-law flow parameters: CRM but using hot geothermal and wet (not damp) olivine

<b>Depth (km)</b>	<b>eff viscosity_o Pa s</b>	<b>differential stress_o MPa</b>	<b>n</b>
<b>0-15</b>	elastic	large	
<b>15-20</b>	6E+24	6000	3
<b>20-25</b>	1.7E+23	170	3
<b>25-31</b>	2.2E+22	22	3
<b>31-40</b>	3.4E+22	34	3.5
<b>40-55</b>	2.3E+21	2.3	3.5
<b>55-80</b>	2.5E+20	0.25	3.5
<b>80-120</b>	7.3E+19	0.07	3.5
<b>120-200</b>	1.0E+19	0.01	3.5

Modeled Hector Mine post-seismic displacements.  
Hot geotherm, wet olivine flow law.







Mojave CRM/CTM **is not** consistent with mantle relaxation needed to explain intermediate- to far-field post seismic deformation

Mojave CRM/CTM **is** consistent with afterslip or shear zone creep (not modeled) within a strong lithosphere

## Thoughts

- add ductile shear zones (separate rheological entity)
- increase background strain rate
- increase temperatures (more like Basin and Range CTM)
- assume higher COH in mantle
- continue testing CRM with numerical models!