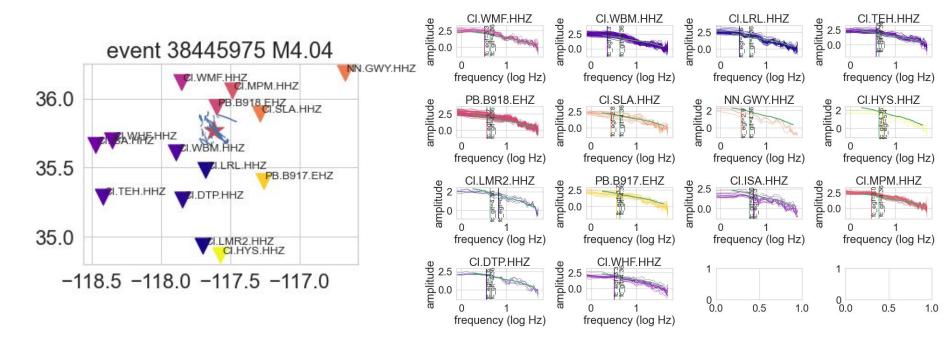
Shanna Chu, comparison of spectral decomposition & eGf spectral ratio methods

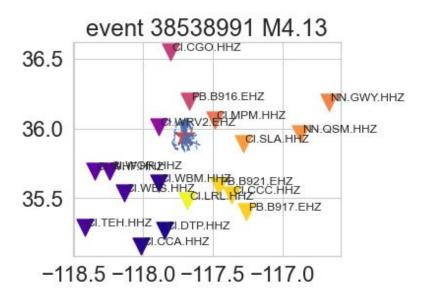
- Data is pre-processed using the same selection criteria: window length, SNR above 3 in 10 frequency bands 1-40 Hz, P-wave
- For one method, I performed spectral decomposition with post-inversion constraint of the event spectra to a spatially-varying correction function derived from pinning stacked events in 0.2 magnitude bins to a Brune model.
- For the second method, I performed spectral ratios using an empirical Green's function. The eGfs are selected to be within 1 source radius of the target events (preliminarily calculated with 2.4 MPa stress drop) and minimum cross-correlation coefficient of 0.5.
- For spectral ratios, deviation from the Brune spectrum can be used to obtain peak-to-peak ratio, a
 proxy for source complexity which has variation across stations. The difference between GIT
 spectra (which uses all stations for an event) also shows a station-by-station difference to the
 spectral ratios.

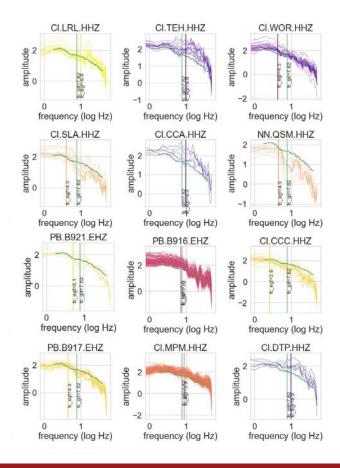
Shanna Chu USGS schu@usgs.gov

 $F_c \sim 4.3$ with GIT and 2.9-5.7 Hz with eGf method. The two methods are closest at PB (borehole) stations.









EXTRA SLIDES

38445975 M4.04 GIT fc=4.36 sd=25.04 MPa 38445975 M4.04 eGf fc=3.03 sd=6.58 MPa 38451079 M4.09 GIT fc=8.21 sd=104.58 MPa 38451079 M4.09 eGf fc=2.44 sd=4.07 MPa 38471103 M3.3 GIT fc=12.84 sd=18.76 MPa 38483215 M3.02 GIT fc=10.77 sd=8.09 MPa 38538991 M4.13 GIT fc=7.82 sd=79.01 MPa 38538991 M4.13 eGf fc=7.51 sd=136.3 MPa 38489543 M2.54 GIT fc=10.61 sd=1.97 MPa

Trey Knudsen and Bill Ellsworth Measuring Source Parameters using Peak Amplitudes

- We measure the seismic moment and corner frequency of earthquakes using narrow band filtered peak displacement amplitudes. Errors are assessed using the bootstrap.
- "Richter"-like frequency dependent attenuation curves are used to correct for geometric spreading and anelastic loss.
- We have also estimated station corrections to the attenuation curves. These may account for path differences, site differences and radiation pattern.
- Results reveal a strong covariance of seismic moment and corner frequency.



Trey Knudsen
Department of Geophysics Stanford University
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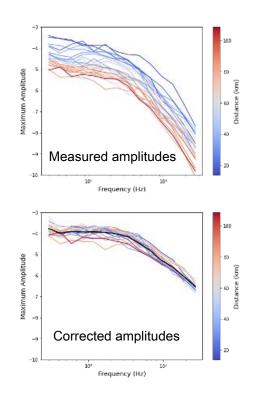


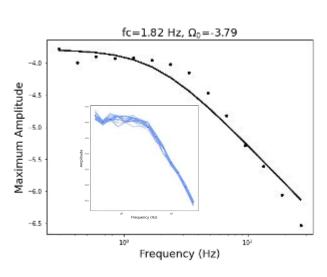
Bill Ellsworth
Department of Geophysics Stanford University
wellsworth@stanford.edu

July 5 00:18 Mcat 4.04 ID 39445975

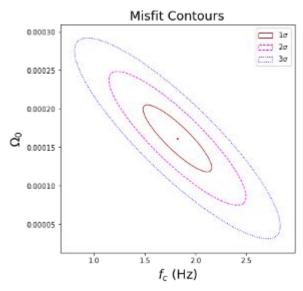
Knudsen and Ellsworth

Mw 4.06, stress drop 9.0 MPa, fc 1.65 Hz





Best-fitting Brune spectrum With bootstrap analysis

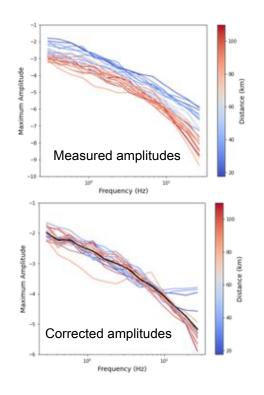


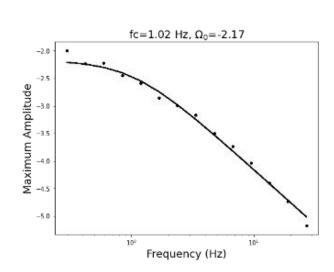
Misfit error analysis
Note covariance between seismic moment
and corner frequency

July 5 11:07 Mcat 5.36 ID 38450263

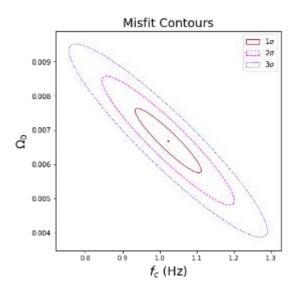
Knudsen and Ellsworth

Mw 5.13, stress drop 65 MPa, fc 1.02 Hz





Best-fitting Brune spectrum

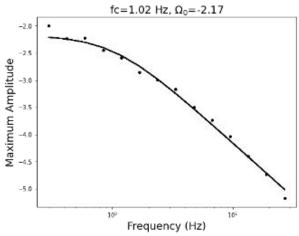


Misfit error analysis

Note covariance between seismic moment
and corner frequency

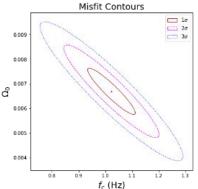
July 5 11:07 Mcat 5.36 ID 38450263

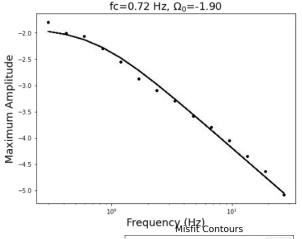
Knudsen and Ellsworth



With station terms

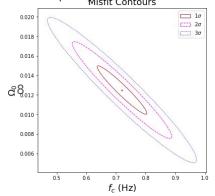
Mw 5.13 fc 1.02 Hz s.d. 65 MPa





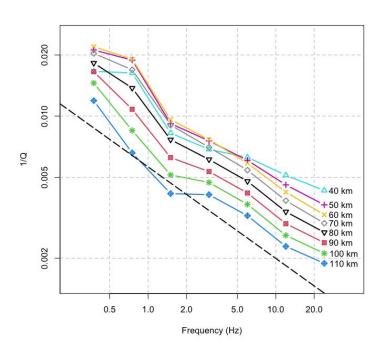
Without station terms

Mw 5.31 fc 0.72 Hz s.d. 42 MPa



Frequency dependence of attenuation

From Al-Ismail, Ellsworth and Beroza (BSSA, 2023)



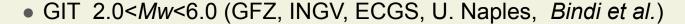
ID	depth ma	g date	time	log w0	fc	MO	Mw	stress drop
		05-Jul-2019 05-Jul-2019				1.67e15 1.97e15		
38471103	7.778 3.30	07-Jul-2019	03:23:26	-4.92	4.62	9.85e13	3.30	10.70
38450263	7.23 5.36	08-Jul-2019 05-Jul-2019	11:07:53	-1.90	0.72	5.17e13 1.03e17	5.31	
		3 11-Jul-2019 08-Jul-2019				1.67e15 1.39e13		
38496551	10.11 2.57	' 09-Jul-2019	05:17:45	-5.93	16.28	9.63e12	2.62	45.78



Multi-Method Comparison

Kevin Mayeda, Dino Bindi, Paola Morasca, Jorge Roman-Nieves, Bill Walter, Doug Dreger, Taka'aki Taira, Chen Ji, Ralph Archuleta

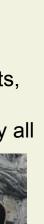




- Energy integration Mw> 4.0(UCSB Chen Ji & Archuleta)
- Finite Fault Inversion 3.3<Mw<5.53 (UCB Dreger & Taira)

For common events from the initial focus events, recent extra events, and selected moderate magnitudes events, we find good correspondence for apparent stress and corner frequency for nearly all range of event sizes.

Kevin Mayeda, AFTAC, kevin.mayeda@gmail.com





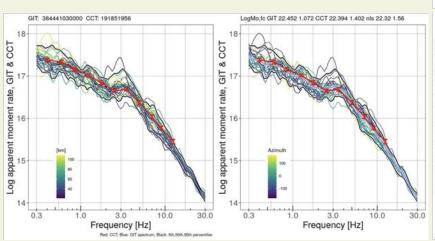


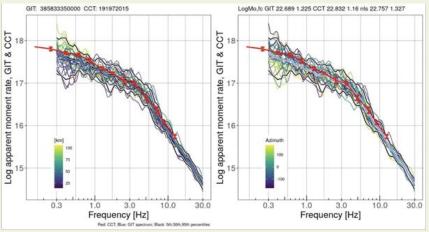


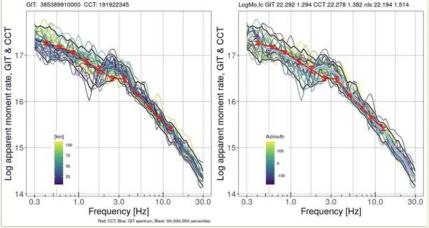


CCT vs GIT

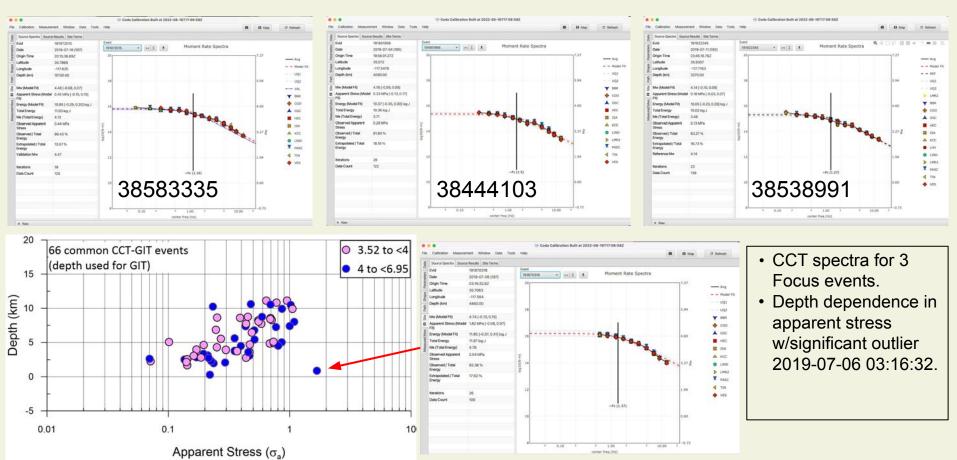
- 3 common 'Focus' events between GIT (*Bindi et al.*, 2021) and CCT spectra are in good agreement.
- 2 events (38444103, 38538991) show a pronounced spectral bump near 3-Hz.
- Correspondence between GIT and CCT also found in central Italy (Morasca et al., 2022)





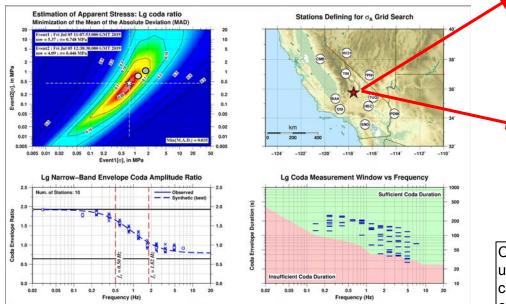


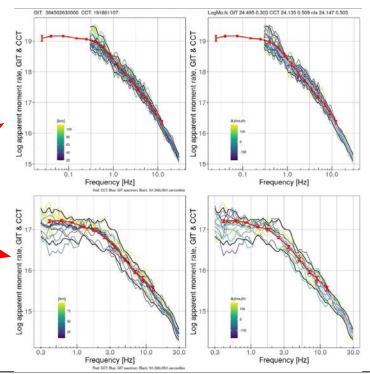
CCT Initial Focus events



GT Source Constraints

CCT uses reference 'ground truth' (GT) spectra for site term corrections using the coda spectral ratio method (*Mayeda et al.*, 2007) avoiding *a priori* scaling assumptions.

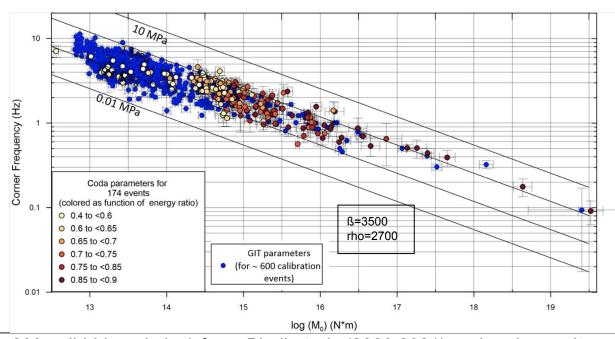




CCT and GIT spectra are in good agreement for an event pair used in determining GT source spectra used in the CCT calibration. These spectra along with independent *Mw*'s for other events provide the station site corrections.

Comparison of GIT vs Coda CT

- Spectra from GIT and CCT methods yield similar radiated energy, Mw and apparent stress.
- For the larger events, we find similar results with those from UCSB (*Ji & Archuleta*, pers. comm., Sept. 2022).



CCT vs GIT: GIT calibration events (~600 solid blue circles) from *Bindi et al.*, (2020;2021) and coda results are shown as white to deep red circles as a function of the ratio of observed to extrapolated energy (174 events) along with error bars derived from NLS inversion. Lines of constant apparent stress are shown as reference and similar to findings from central Italy by *Morasca et al.*, (2022).

Apparent Stress for Larger Events w/UCSB

- We compare results of 25 larger magnitude events also processed by UCSB (*Ji & Archuleta*) using three different methods.
- In general, agreement is good and future study to understand the outliers is planned.

Figure by Chen Ji and Ralph Archuleta:

GR: using the modified Gutenberg-Richter method (*Kanamori, et al.*, 2020). Time domain, no local attenuation correction.

F0: frequency approach. Corrects near surface impedance, k_0, crust attenuation, similar to *Boatwright et al.* (2003). The radiation coefficients are assumed to be 0.6 (*Boore and Boatwright*, 2003). It is on average 20% larger than GR.

F1: Similar to F0, except we use Caltech moment tensor to estimate radiation coefficients and only select the measurements with radiation coefficients larger than 0.25. It is on average 38% larger than GR.

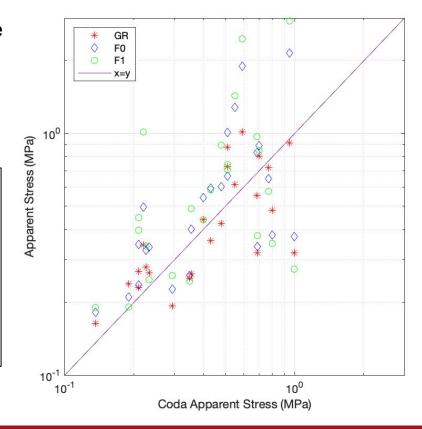


Table of solutions (EXTRA SLIDE)

EventID	Mw	Depth	Sta.	β_0 (m/s)	$\rho_0 (kg/m^3)$	$f_c(Hz)$	$\Delta\sigma_B(\mathrm{MPa})$	σ_a (MPa)	$\sigma_a(CCT^1)$	$f_c(CCT)$	$f_c(\text{Trugman}^2)$
38445975	4 0	2 30	12	3000	2650	1.08	0.50	0.21	0.16	1.54	
38450263			14	3650	2850	0.57	5.18	1.22*	0.89	0.51	0.57
38451079	4.1	7.34	12	3650	2850	1.70	1.53	0.61	0.49	1.78	
38538991	4.0	2.76	11	3000	2650	1.18	0.65	0.15	0.15	1.38	

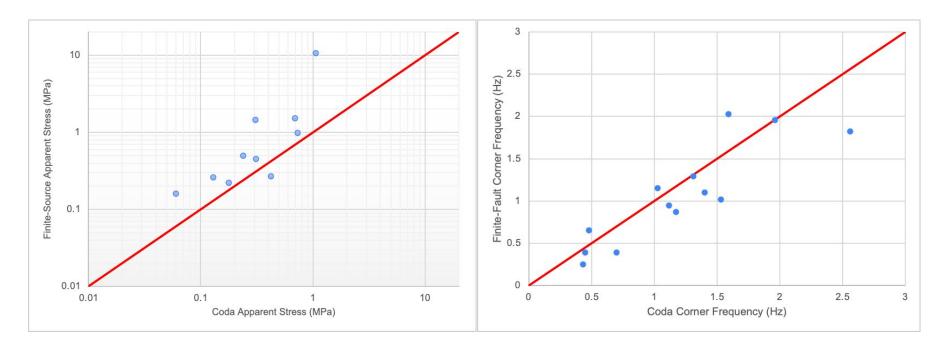
Values are adjusted to consider the rigidity difference

² Trugman (2022, JGR).

8 selected events

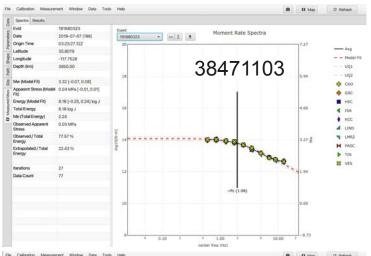
ID	Z	Mag	date-time	CCT fc	CCT Mw	GIT fc	FF fc
38445975	2.3	4.04	'05-Jul-2019 00:18:01'	1.31	4.07	1.02	1.29
38451079	7.3	4.09	'05-Jul-2019 12:38:30'	1.53	4.10	1.54	1.02
38471103	7.7	3.30	'07-Jul-2019 03:23:26'	1.96	3.32	3.43	
38483215	7.7	3.13	'08-Jul-2019 05:02:10'	3.41	3.02	4.29	
38450263	7.2	5.36	'05-Jul-2019 11:07:53'	0.45	5.37	0.27	0.39
38538991	2.7	4.13	'11-Jul-2019 23:45:18'	1.15	4.13	1.25	0.87
38489543	2.8	2.50	'08-Jul-2019 17:30:03'	N/A			
38496551	10.1	2.57	'09-Jul-2019 05:17:45'	5.31	2.55	10.18	

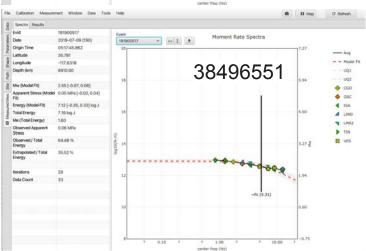
Comparison with Finite Fault Inversion (Dreger)

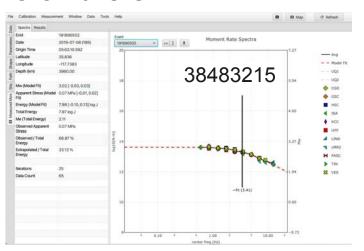


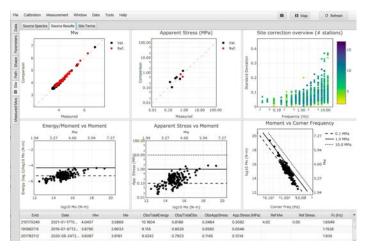
CCT and finite fault estimates of apparent stress and corner frequency for 13 moderate sized events (3.27<Mw<5.53) are in rough agreement. These results are preliminary and outliers will be investigated as will the assumption on *mu*.

CCT Mw 2's and 3's

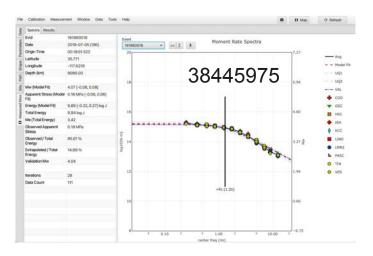


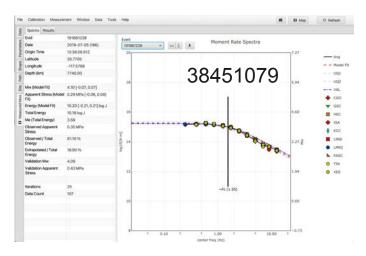


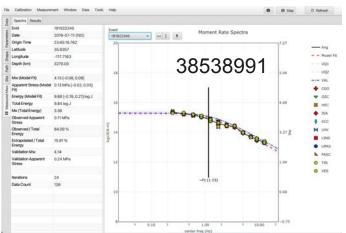


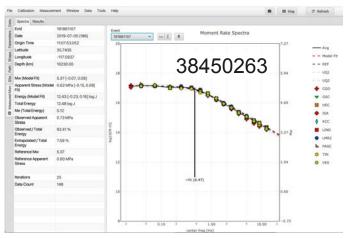


CCT Mw 4+





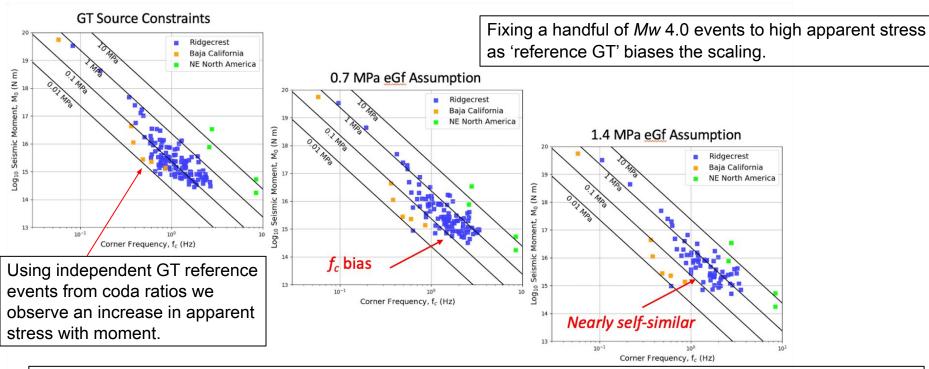




Summary

- Varying combinations of GT reference source spectra were tested with no significant changes in the final calibration parameters, source spectra, Mw's and source scaling.
- Mw's match those derived from moment-tensor solutions (Validation Mw's)
 3.4<Mw<6.9 and simultaneously agree with 6 independent apparent stress estimates.
- We can extend stable Mw well below what can be routinely waveform modeled.
- Minimal azimuthal variation observed in single-station source spectra.
- Event 38538991 (*Mw* 4.14) has a spectral hole at ~1.5-2.0-Hz for all stations.
- Event 38445975 has slightly lower apparent stress than event 38451079.
- For common events, apparent stress and corner frequency estimates derived from UC Berkeley's finite fault (3.3<Mw<5.53) are in good agreement, as well as those from UCSB (Mw>4.0) and GIT results (2.0<Mw<6.0).

Source Constraint Affects Site Terms and Scaling



To test the ECS approach, we chose six *Mw* ~4.0 'GT reference events' to have apparent stress of 0.7 MPa and 1.4 MPa, which under the *Brune* source shape assumption is roughly equal to a 3 MPa and 6 MPa *Brune* stress drop (center and right figure, respectively). When compared against the calibration that used GT spectral constraints (left), we see a bias at lower magnitudes and an overall increase in apparent stress over the entire magnitude range. (Note: Only the site terms change. Path correction and envelope shape remain the same but cannot match original GT ratios in previous slide.)

EXTRA SLIDES

ID logMo SElogMo fc SEfc stress_drop_vs_3D SEstdr stress_drop_vs_const

38445975 15.3299270379856 0.048093448591591 1.01800708086084 0.0709773196228018 0.701005826965392 0.153084738877273 0.582907428046382

38451079 15.2838751467793 0.0141603841348427 1.53798101398198 0.0340682640793185 1.40095798764162 0.086164369196776 1.80779101987969

 $38471103\ 13.9124399875432\ 0.0156534898192737\ 3.43279312034761\ 0.0965513082003814\ 0.623354726754051\ 0.0324365772827538\ 0.854679006011451$

38483215 13.6108637499221 0.0128252383736518 4.29091716382966 0.105832422723926 0.607163883877048 0.0458378593156059 0.833555068929524

38450263 17.6119424143883 0.0883602670144013 0.267276416901024 0.0287987412297874 1.56019827783463 0.592304288984253 2.01950294683978

38538991 15.3238473611207 0.0298362312433606 1.25085386492333 0.0561163992243787 1.22449686483017 0.163947443674182 1.06631922849795

38496551 12.8702041635577 0.0100399566274271 10.1845234583118 0.281673992958635 1.47276926322507 0.30875575880591 2.02509992377904

Spectral Ratios - Choice of EGF, Time Window etc.

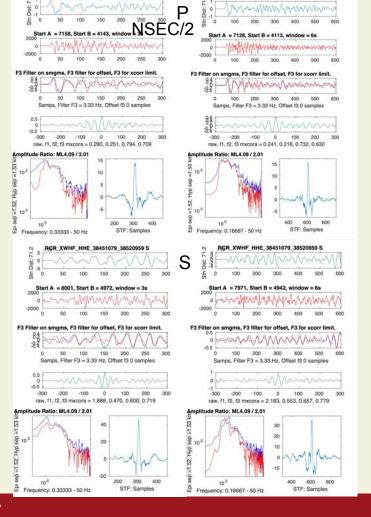
Apply Methods of Abercrombie et al., 2017; 2020 etc. to Ridgecrest Earthquakes

> Find all Small events within 2 km epicentral distance 1.2 and 2.5 M units < main (both ranges iteratively decreased IF >>200 EGFs found)

Define time window NSFC = constant * $M0^{1/3}$ Try 0.5NSEC to get more P waves for close stations

Try various EGF threshold depending on cross correlation and hypocentral depth difference

Rachel Abercrombie **Boston University** rea@bu.edu





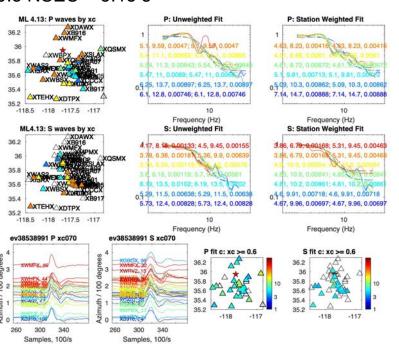
Example EGF Spectral Ratios: 38538991 M4.13

Minimum cross correlation = [0.5 0.5 0.6 0.7 0.7 0.8 0.9]

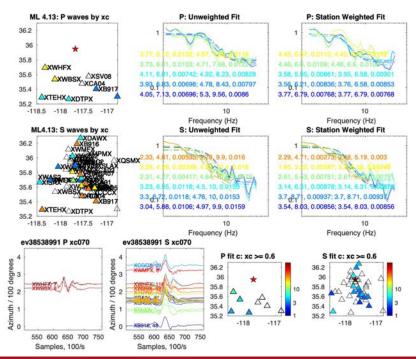
Max depth difference to main = [10 1 10 10 1 10]

Rachel Abercrombie

0.5 NSEC = 3.15 s



1 NSEC = 6.3 s



Example EGF Spectral Ratios: 38538991 M4.13

Minimum cross correlation = [0.5 0.5 0.6 0.7 0.7 0.8 0.9]

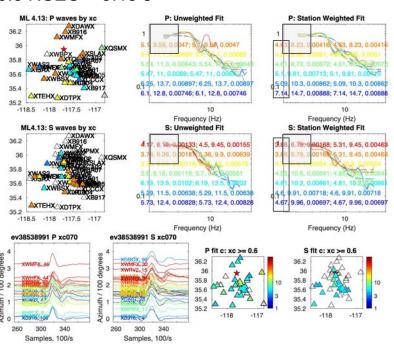
Max depth difference to main = [10 1 10 10 1 10 10]

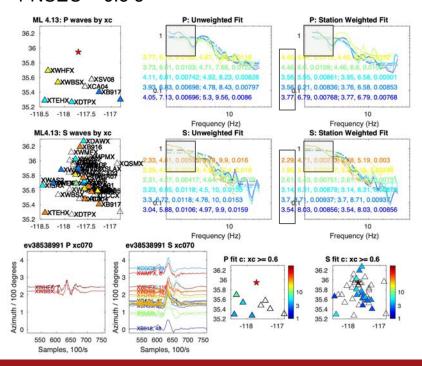
Rachel Abercrombie

0.5 NSEC = 3.15 s

fc1, fc2, variance (no constraints) fc1, fc2, variance (limits on fc2)

1 NSFC = 6.3 s

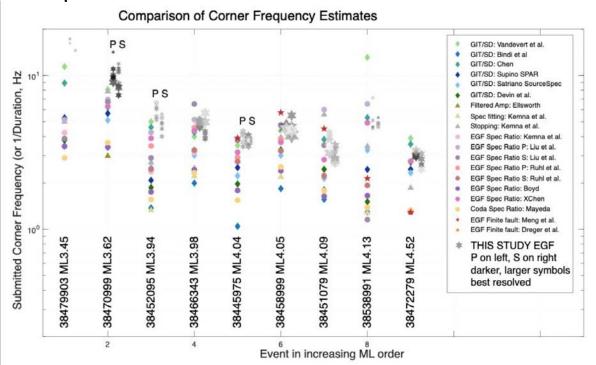


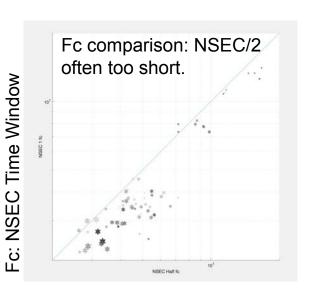


Look for consistency in Different Cross correlations? Time Window affects Frequency Range

Rachel Abercrombie

Comparison of EGF Variation Effects to Other Methods



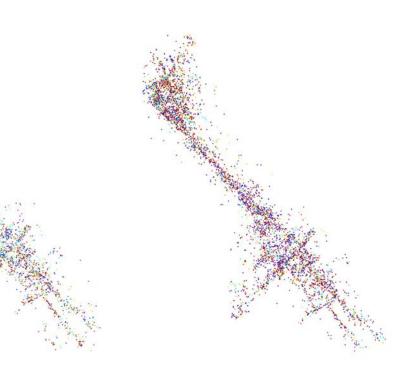


Fc: NSEC/2 Time Window

Ridgecrest P and S spectral decomposition comparisons

Peter Shearer, Ian Vandevert & Wenyuan Fan *IGPP/SIO/U.C. San Diego*





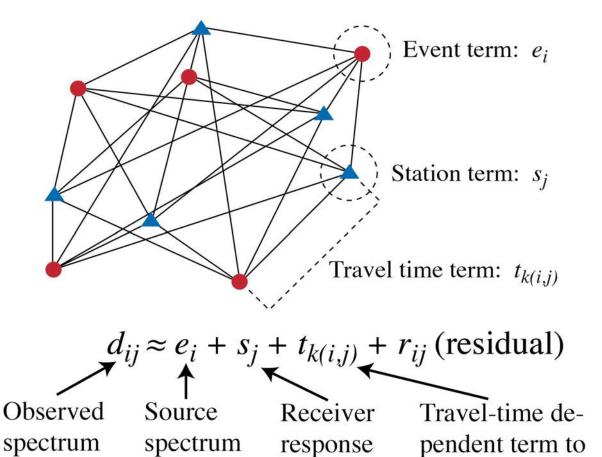








Spectral decomposition



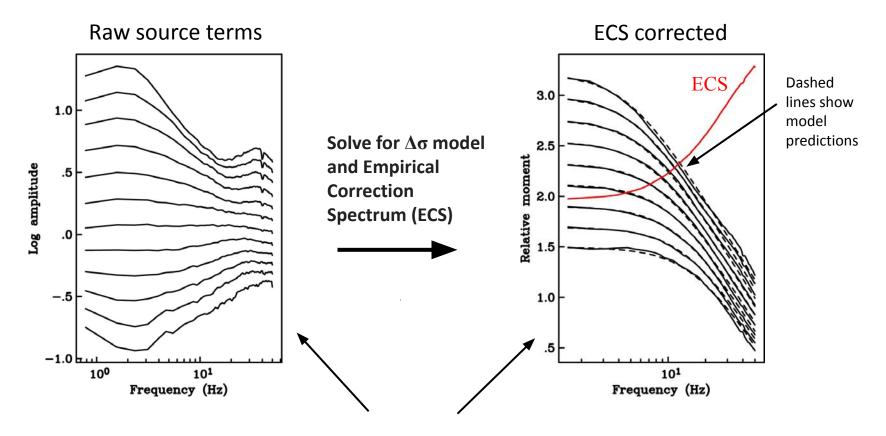
account for Q

Good for uniform processing of large data sets.

Variations of this approach used by *Shearer*, *Allmann*, *Chen*, *Trugman*, *Oth*, and others.

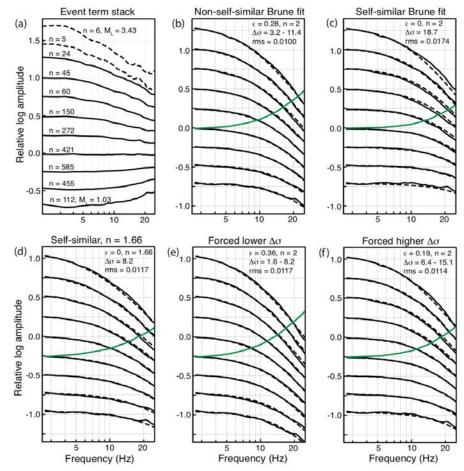
Completely empirical—a great advantage but one that has its limitations

Old approach: Solve for global ECS function and fc model that best fit entire data set



Source spectra binned by relative moment

Results indicate method has too many free parameters!



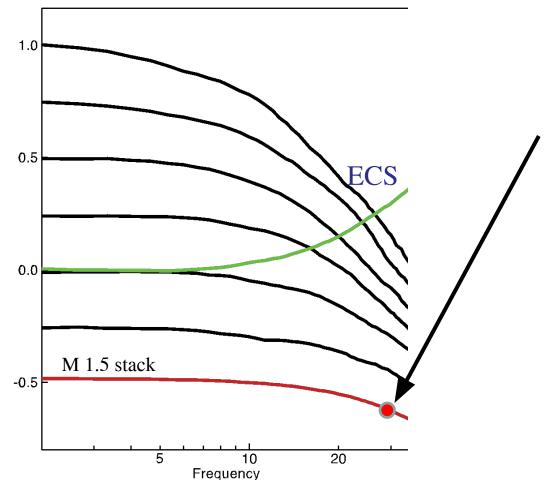
from Shearer, Abercrombie, Trugman & Wang (2020)

For Landers cluster, many different models can fit the data, given tradeoffs among the model parameters and the global EGF function.

Thus, we cannot be sure of the absolute level of stress drop, the high-frequency fall off rate, or non-self-similar scaling factors. These issues require **independent** constraints on the EGF function.

However, relative differences in stress drop among events in the same cluster are well-resolved, when estimated using the same EGF function.

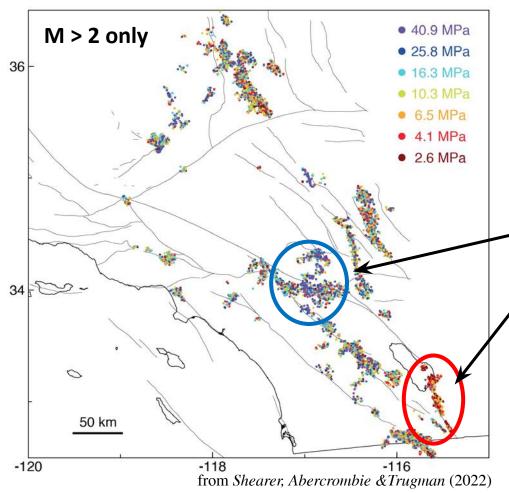
New approach: Locally fix small earthquake average corner frequency



Force the Brune corner frequency (fc) of nearby M 1.5 earthquakes to 30 Hz in estimating the ECS for each target event.

This directly determines the ECS at each location and ensures that any spectral differences seen in M > 1.5 earthquakes are caused by source variations rather than inaccurate path corrections.

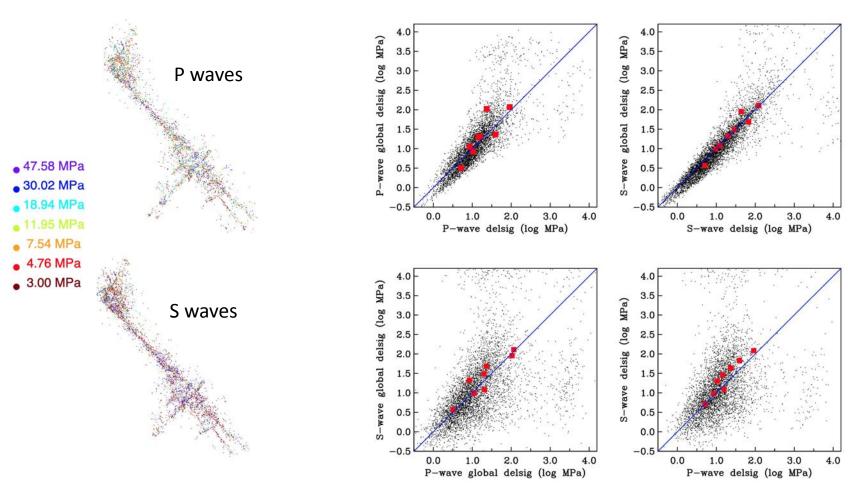
Our results for southern California: robust spatial variations

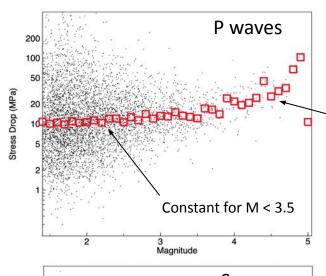


Stress drop estimates for 28,685 M 2 to 4 earthquakes (1996–2019). Each earthquake has at least 10 M ≤ 1.6 calibration events (assumed to have Brune fc = 30 Hz) within 5 km in horizontal distance and 2 km in depth.

These lateral variations in average stress drop must be real because they are derived from the relative behavior of M > 2 quakes with respect to M 1.5 quakes in each local region, i.e., any propagation path differences are removed.

Applications to Ridgecrest test dataset

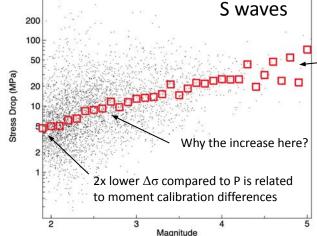




Median stress drop vs. magnitude shows increase—is it real?

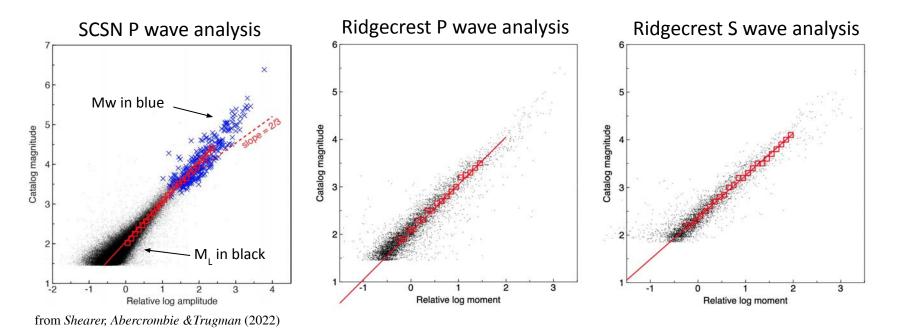
Increase at M > 3.5 is likely mostly artifact caused by unresolved low-frequency part of spectrum and/or HF falloff rate < 2

Somewhat more stable results for M > 4.5 quakes than P results

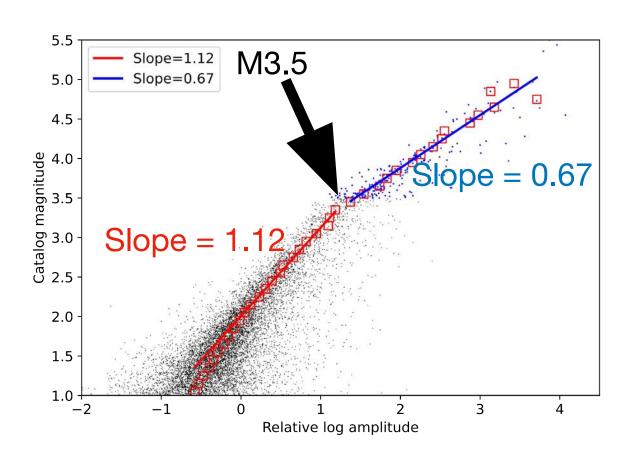


Resolving the low-frequency (< 1 Hz) part of the spectrum is key for getting reliable results for larger (M > 4) quakes.

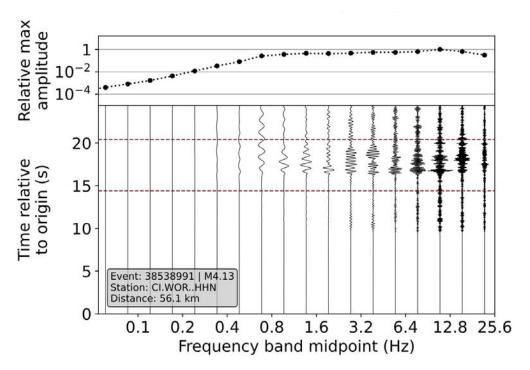
Problem: relative moments vs catalog magnitude do not show expected change in slope as M_I changes to Mw for M > 3.5



Ridgecrest S-wave amplitude decomposition and comparison to S-wave spectral decomposition

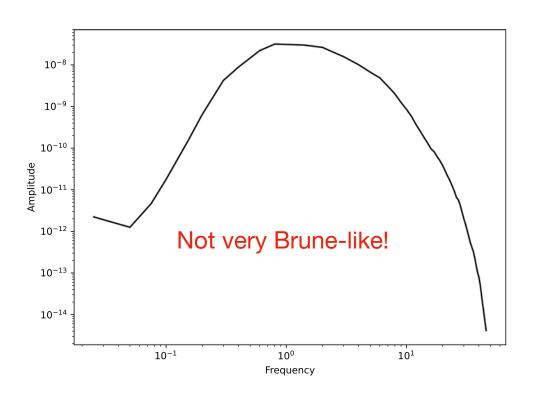


New method: S-wave amplitude decomposition in the time domain

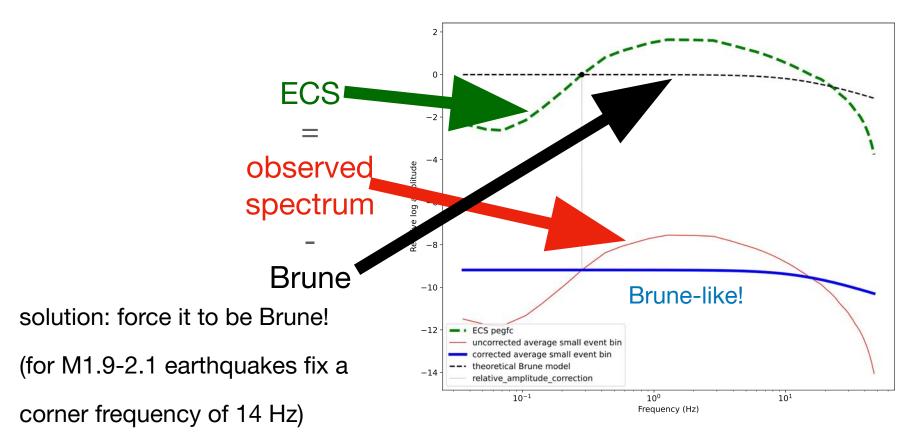


- Estimate S-wave arrival
- 2. Filter entire trace at different frequency bands, measure peak amplitude
- 3. Assemble observed spectrum
- 4. Invert to get event terms

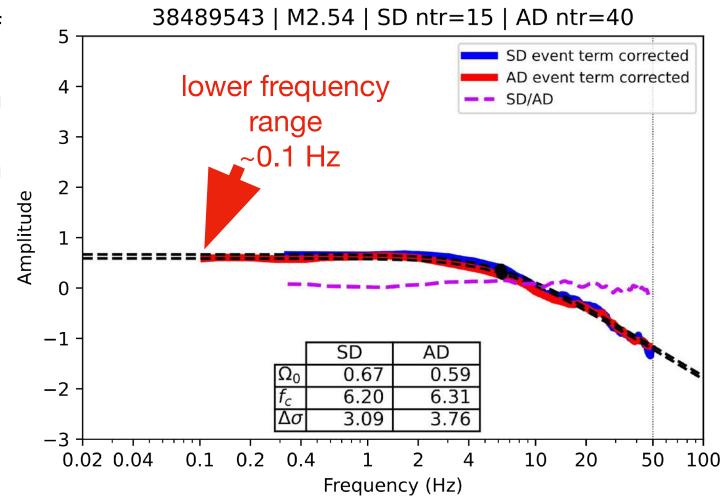
Example event term (S-wave displacement)

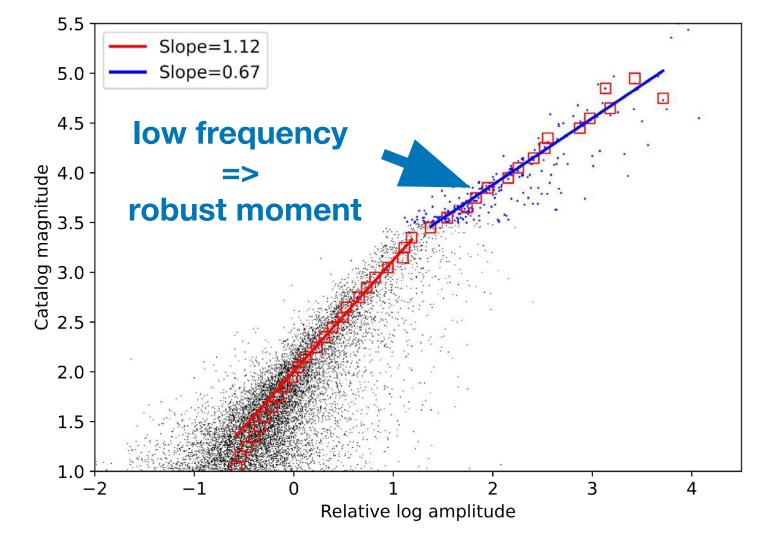


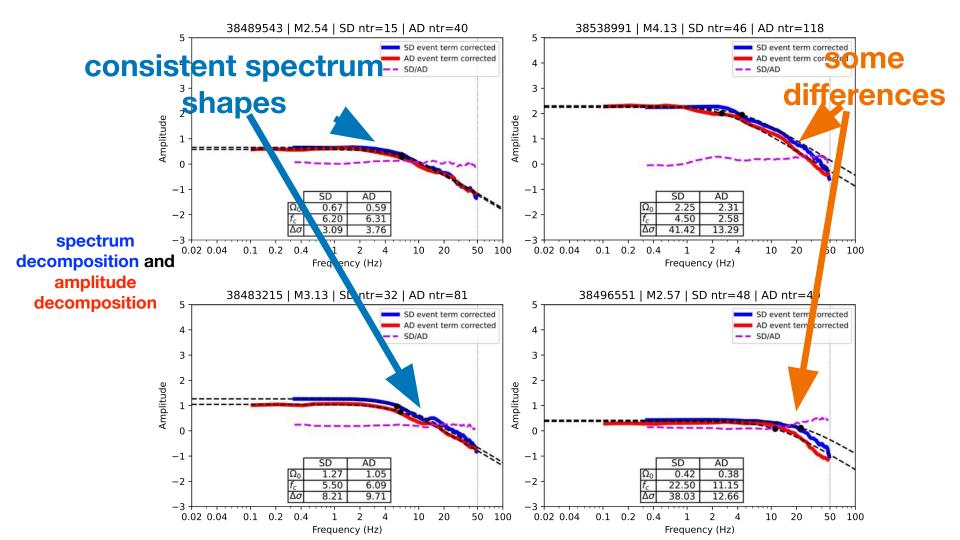
Empirical correction spectrum (ECS)

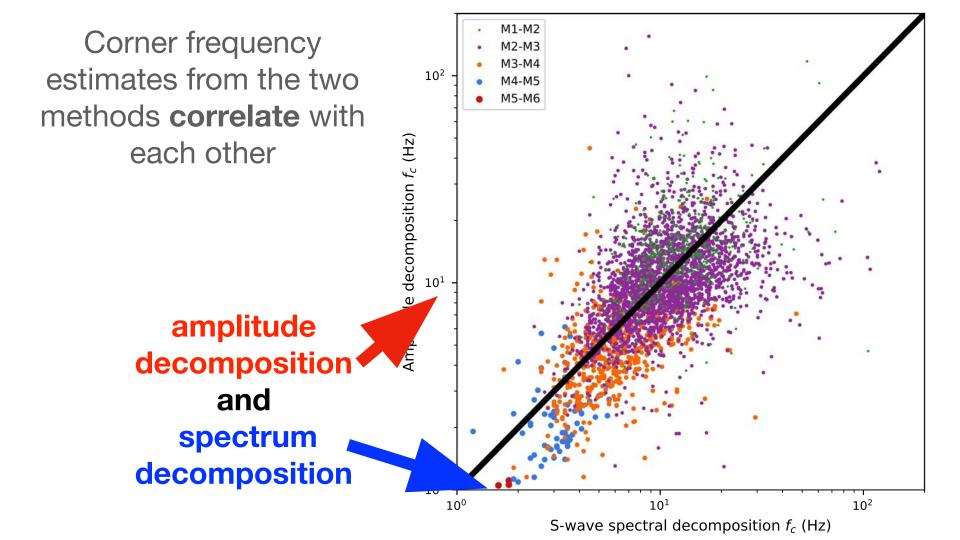


Comparison of spectral decomposition to amplitude decomposition

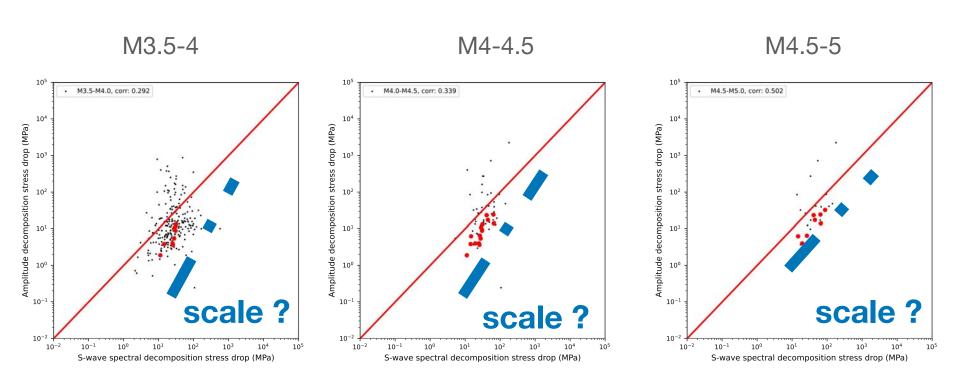








stress drop estimates from the two methods **correlate** with each other (M>3.5)



stress drop estimates from the two methods **show scatters for** M<3.5 ??

