# The SCEC Community Stress Model (CSM) Project

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# **Community Stress Model (CSM):**

- New project starting in SCEC4.
- Long-term goal: a model or set of models of stress and stressing rate in the southern California lithosphere.
- The CSM will probably not be a single model, but a (hopefully small) set of models.
- Any branches in the CSM will be based on clearly-defined differences in data or assumptions.

#### Who are the users of the CSM and what do they need?

**External users:** Rupture dynamics, geodynamics, seismic hazard, stress triggering, others?

**Needs:** You tell us! (Invite representatives of external user communities to all workshops.)

**Provide:** One or more reference stress and stressing rate models, accessible through an interface developed jointly with the user communities.

**Internal users:** Researchers working on problems directly related to stress.

**Needs:** Access to existing data and models, easier ways to integrate and compare models and observations.

**Provide:** A modeling environment with tools that will enable researchers to develop and test candidate models against suites of data and/or quantitatively compare their models with other models.

## **Progress:**

- First workshop, September 2011, at SCEC Annual Meeting. Large wide-ranging group discussion.
- Second workshop, October 2012: compiled and compared existing stress and stressing rate models from the SCEC community.
  - All models in common format on pre-defined 3D grid.
  - Most models were submitted as full 6-component stress or stressing rate tensors.
- Third workshop planned for May 29-30 2013: focus on reconciling stressing rate models, and validating models with data.

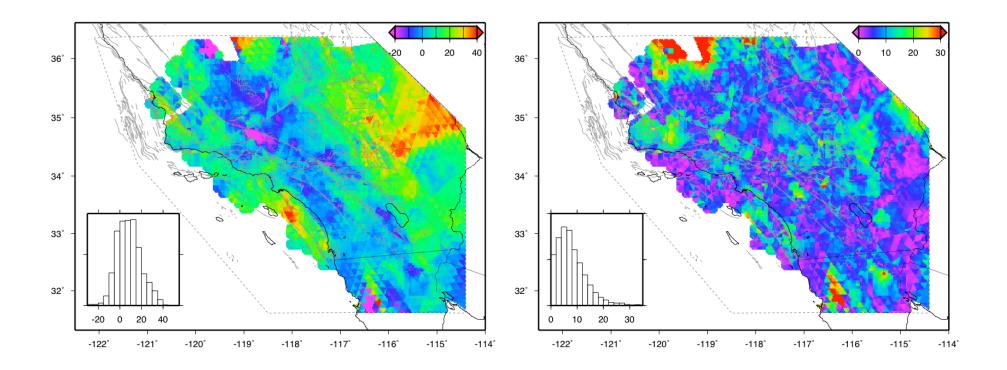
#### Contributed Models:

#### Stress:

- 1) Inversion of focal mechanisms for stress orientation. Wenzheng Yang and Egill Hauksson (Caltech).
- 2) Finite element model including topography, depth-dependent rheology, frictional faults, and long-term deformation model. *Peter Bird (UCLA)*.
- 3) Inversion for stress field that fits topography, fault loading from dislocation model, tectonic loading, and focal mechanisms. *Karen Luttrell (USGS), Bridget Smith-Konter (Texas), and David Sandwell (UC San Diego).*
- 4) Global model from density-driven mantle flow, plus lithosphere gravitational potential energy, fit to geoid and global plate motions. *Attreyee Ghosh and Thorsten Becker (USC).*

SHmax trend (degrees); depth=5 km

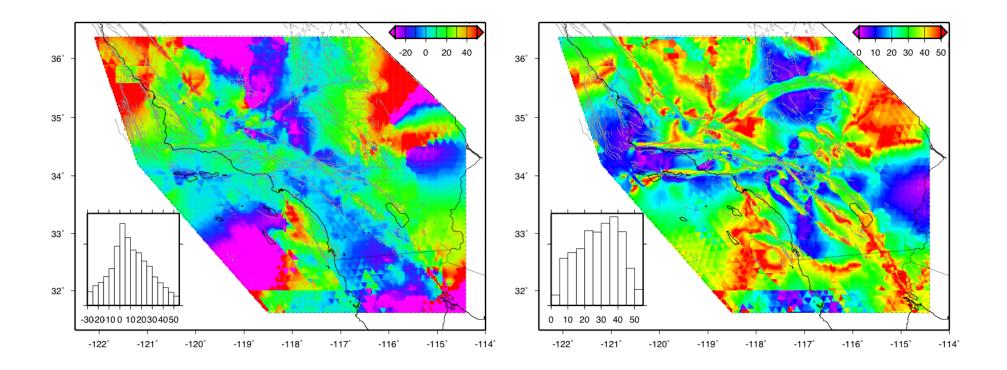
SHmax RMS (degrees); depth=5 km



<sup>\*</sup> Average of Bird; Luttrell, Smith-Konter & Sandwell; and Yang & Hauksson models, everywhere at least two of these models are defined.

SHmax trend (degrees); depth=19 km

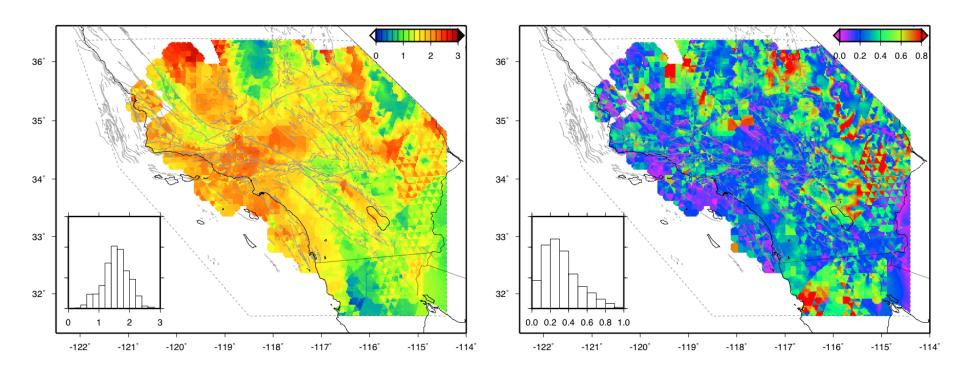
SHmax RMS (degrees); depth=19 km



<sup>\*</sup> Average of Bird; Luttrell, Smith-Konter & Sandwell; and Yang & Hauksson models, everywhere at least two of these models are defined.

A\_phi; depth=5 km

A\_phi RMS; depth=5 km



phi= $(\sigma 2-\sigma 3)/(\sigma 1-\sigma 3)$ 

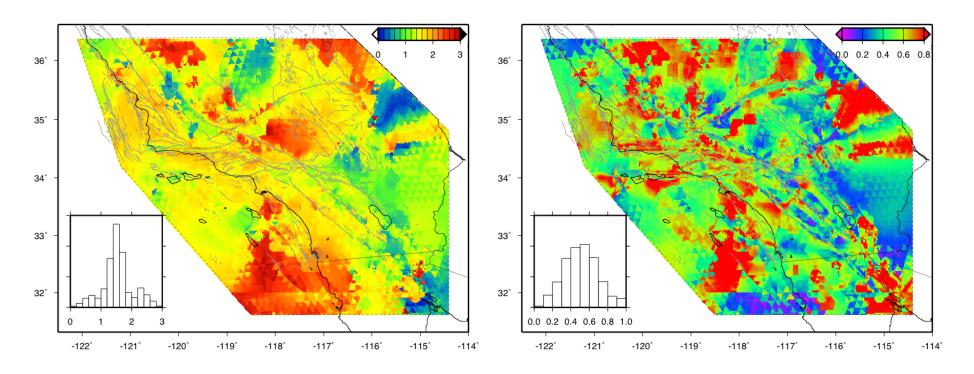
Aphi=phi 0-1 : normal faulting (σ1 most vertical)

Aphi=2-phi 1-2 : strike-slip faulting ( $\sigma$ 2 most vertical)

Aphi=2+phi 2-3 : reverse faulting ( $\sigma$ 3 most vertical)

A\_phi; depth=19 km

A\_phi RMS; depth=19 km

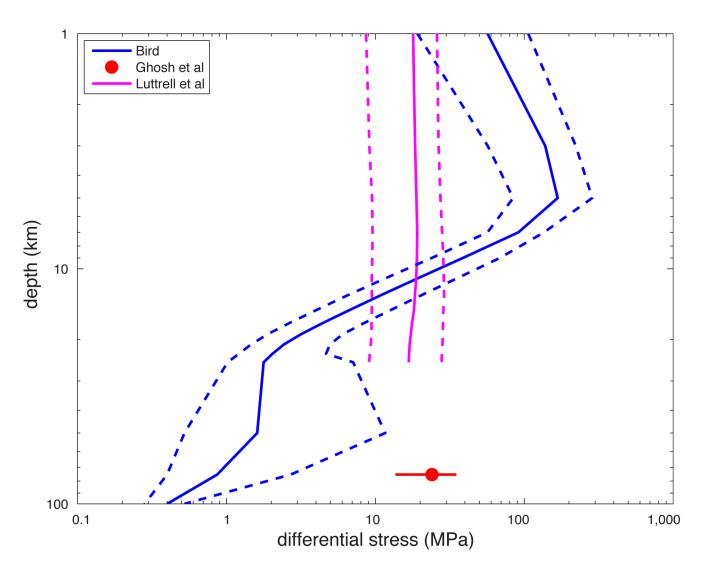


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Stress Models: differential stress ( $\sigma$ 1- $\sigma$ 3) versus depth.



Solid line/symbol: median. Dashed line: middle 68%.

#### Contributed Models:

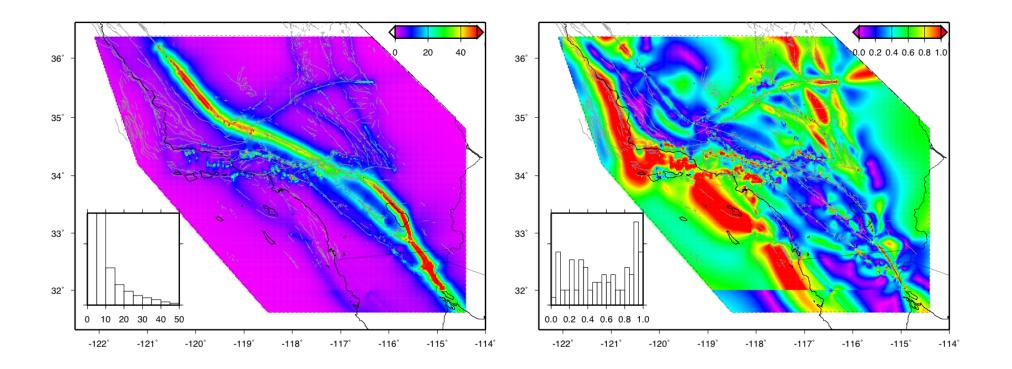
#### Stressing Rate:

- 1) Block model fit to geodetic data. *Jack Loveless (Smith) and Brendan Meade (Harvard).*
- 2) Fault loading from dislocation model using geologic and geodetic slip rates. *Bridget Smith-Konter (Texas), and David Sandwell (UC San Diego).*
- 3) Fault loading from dislocation model plus static stress changes from earthquakes. *Anne Strader and David Jackson (UCLA).*
- 4) Local boundary element model fit to slip rates. *Michele Cooke* (UMass) and Scott Marshall (Appalachain State).

## Average Stress Rate Model: average differential stressing rate: $\delta(\sigma 1-\sigma 3)/\delta t$ .

diff stressing rate (kPa/yr); depth=5 km

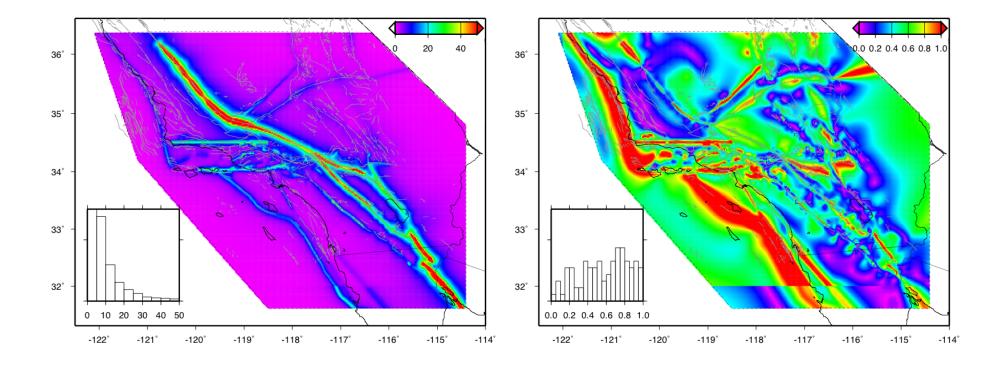
diff stressing rate RMS (fraction); depth=5 km



## Average Stress Rate Model: average differential stressing rate: $\delta(\sigma 1-\sigma 3)/\delta t$ .

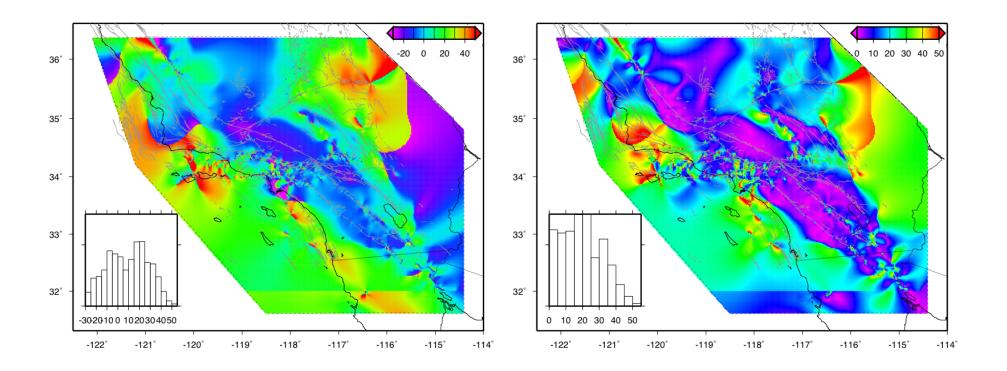
diff stressing rate (kPa/yr); depth=17 km

diff stressing rate RMS (fraction); depth=17 km



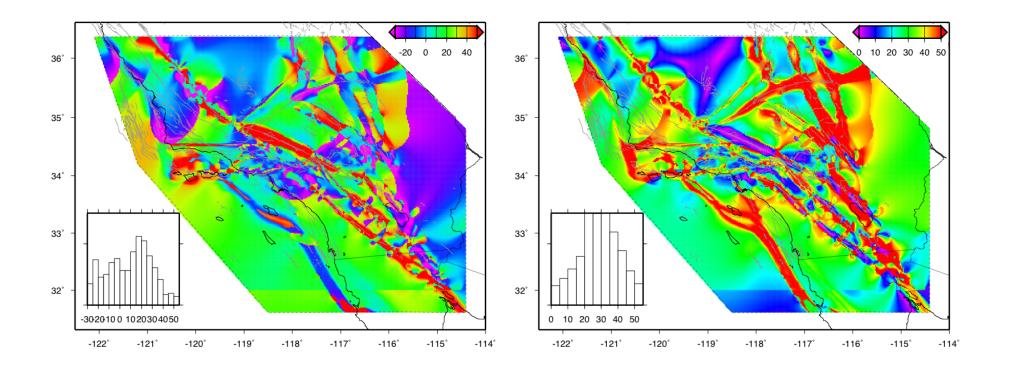
SHmax trend (degrees); depth=5 km

SHmax RMS (degrees); depth=5 km



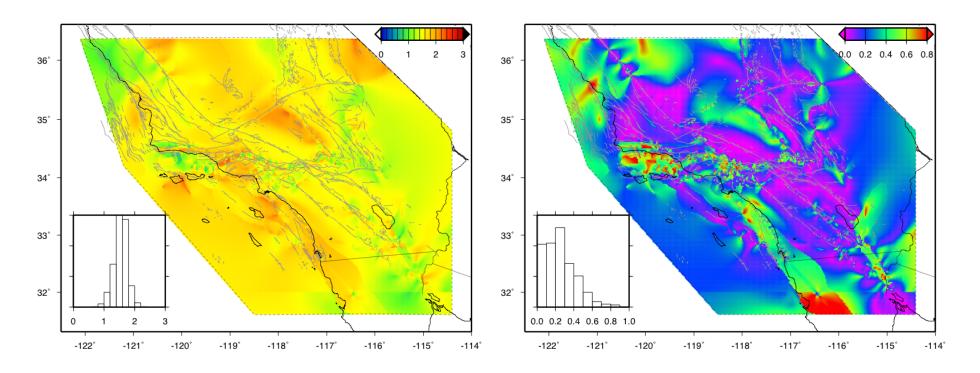
SHmax trend (degrees); depth=19 km

SHmax RMS (degrees); depth=19 km



A\_phi; depth=5 km

A\_phi RMS; depth=5 km

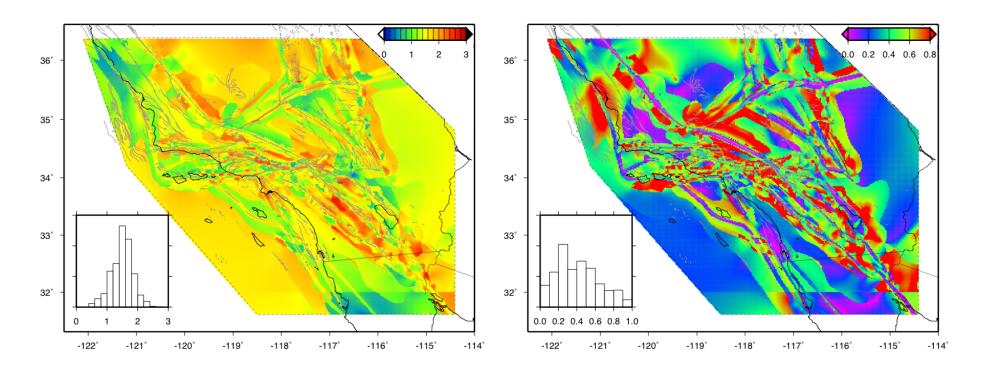


phi= $(\sigma 2-\sigma 3)/(\sigma 1-\sigma 3)$ 

Aphi=phi 0-1 : normal faulting ( $\sigma$ 1 most vertical) Aphi=2-phi 1-2 : strike-slip faulting ( $\sigma$ 2 most vertical) Aphi=2+phi 2-3 : reverse faulting ( $\sigma$ 3 most vertical)

A\_phi; depth=19 km

A\_phi RMS; depth=19 km



phi= $(\sigma 2-\sigma 3)/(\sigma 1-\sigma 3)$ 

Aphi=phi 0-1 : normal faulting ( $\sigma$ 1 most vertical) Aphi=2-phi 1-2 : strike-slip faulting ( $\sigma$ 2 most vertical) Aphi=2+phi 2-3 : reverse faulting ( $\sigma$ 3 most vertical) How can the CSM benefit from a better understanding of ductile rheology?

- Physics-based stress and stressing rate models rely on assumptions about rheology.
- Stress and stressing rate models are generally in good agreement in the upper crust where elastic and brittle deformation dominate.
- Models become more different near the base of the seismogenic zone. Much of this disagreement is due to differences in assumed fault locking depth – better understanding of the brittle-ductile transition could reduce this source of uncertainty.
- Very poor agreement of models below seismogenic depths. Therefore, the depths where we are in most need better constraints are the depths where ductile rheology is important.
- ➢ If you have a physics-based model of stress and/or stressing rate in the lithosphere, please consider contributing it to the CSM project.