

***Is there a role for stress change modeling (fault interactions) to forecast aftershocks?***

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***Answer: Yes, in hybrid mode with empirical methods (like ETAS).***

***Results from:***

*Parsons, T., Y. Ogata, J. Zhuang, and E. L. Geist (2012), Evaluation of static stress change forecasting with prospective and blind tests, Geophysical Journal International, v. 188, 1425–1440, doi: 10.1111/j.1365-246X.2011.05343.x.*

*Segou, M., T. Parsons, and W. Ellsworth (2012), Rate/state friction model implementation for earthquake forecasts in northern California, Seismological Research Letters, v. 82. (manuscript in preparation).*

<http://earthquake.usgs.gov/regional/nca/seminars/2012-05-23/>

Or <http://on.doi.gov/segou-seminar-2012>

Or Google “USGS earthquake seminars”

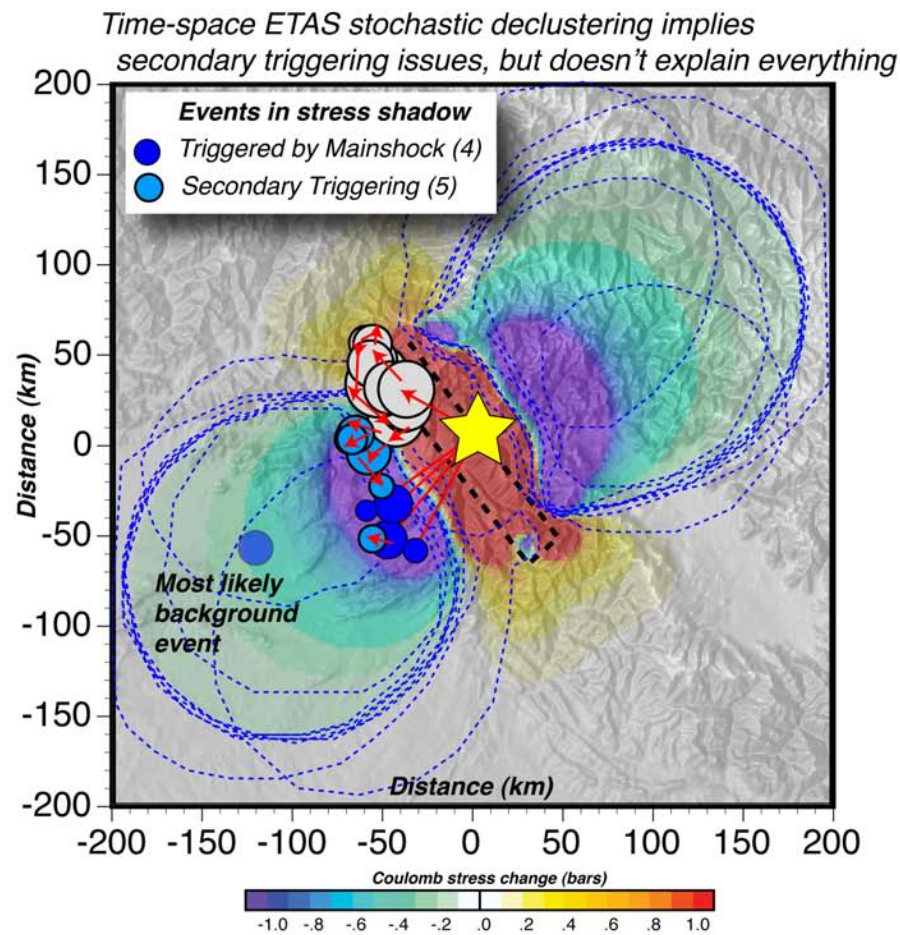
***Issue: Prospective testing of rapid Coulomb stress calculations routinely shows violations of calculated stress shadows (i.e., 8 October, 2005 M~7.6 Kashmir earthquake).***

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Causes include: •Rupture source complexity

•Unmodeled dynamic triggering

•Temporal stress evolution (secondary triggering)





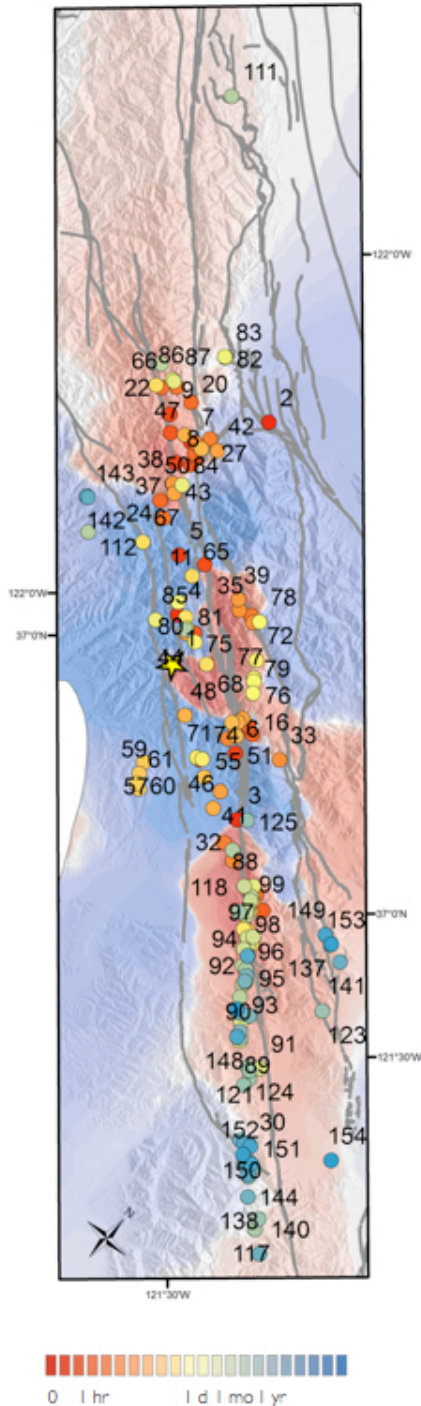
# HOW IMPORTANT IS SECONDARY TRIGGERING FOR FORECASTING?

Motivation: Recent studies have supported that incorporation of secondary triggering is a critical aspect in operational forecasting [Parsons et al., 2012] and past studies refer to this as an possible source of uncertainty leading to poor performance of CRS with time [Toda et al., 2005].

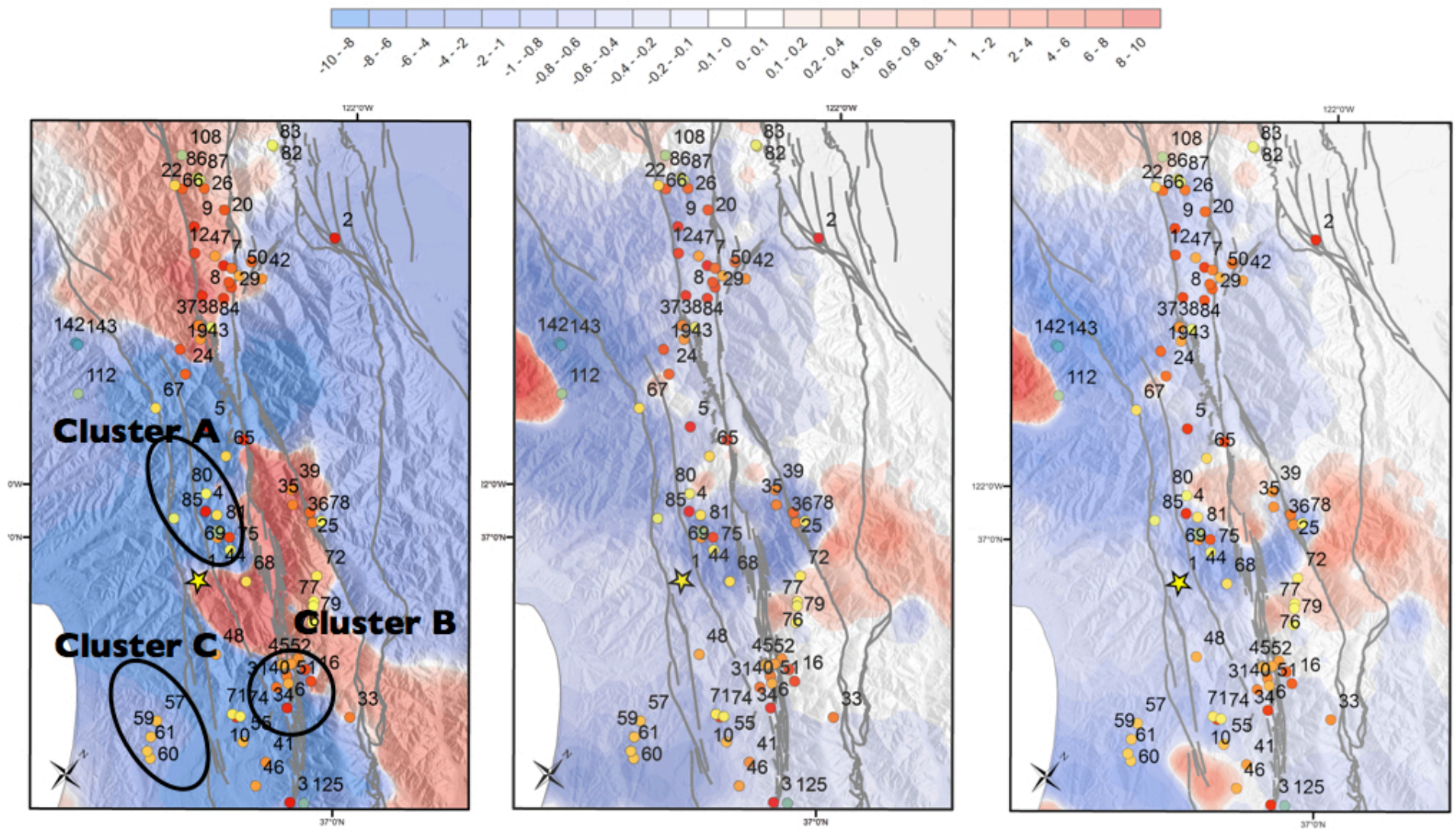
Can we improve CRS-models by including stress perturbations of smaller events ?

Designing the test: Loma Prieta aftershocks with  $M > 3.5$  lying in the stress shadow of the mainshock

Goal: Not to study how efficient CRS forecast models are, but that incorporating stress changes from  $3.5 < M < 5.0$  makes a difference







$$\mu = 0.2$$

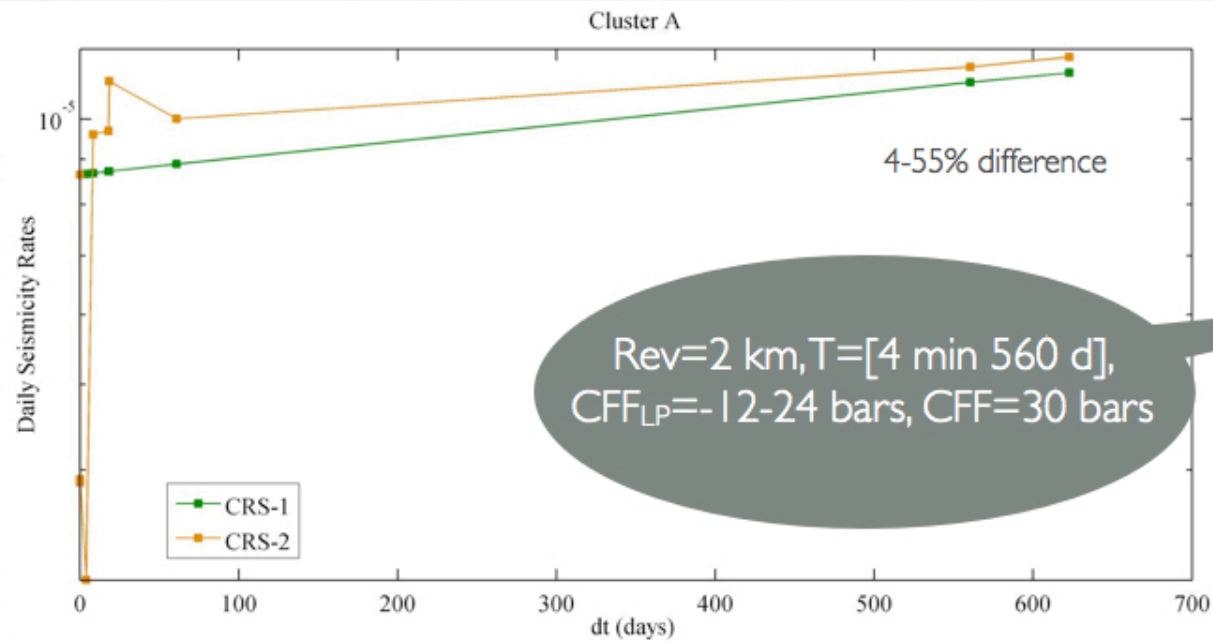
$$\mu = 0.4$$

$$\mu = 0.6$$

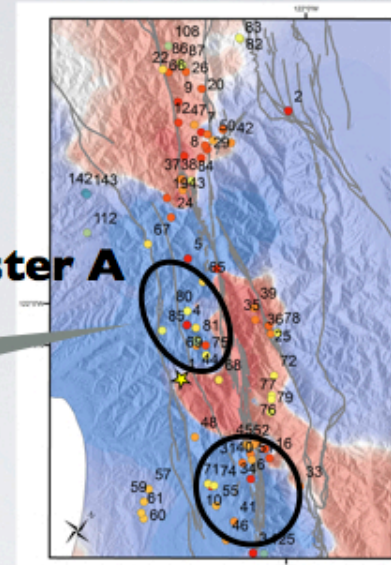
SENSITIVITY COULOMB STRESS CHANGE CALCULATION



# CRS-1 VS CRS-2



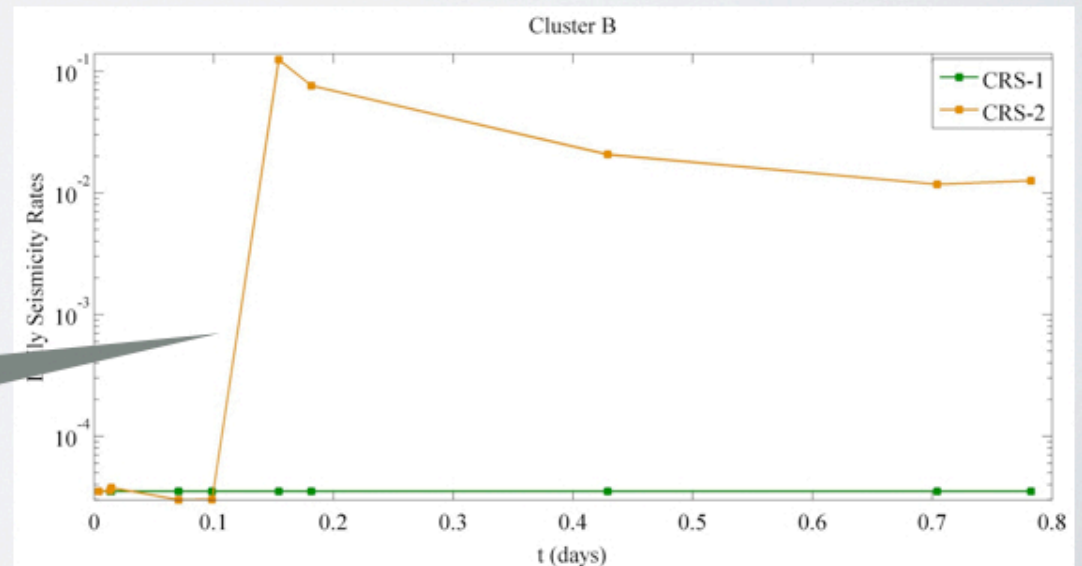
**Cluster A**



**Cluster B**

CFF calculation grid 200m, depth 10 km

Rev=1.2 km, T=[1 d],  
CFF<sub>LP</sub>=-30bars, CFF=30 bars



***Tests are showing that ~50% of shadow violations can be explained by secondary static triggering***

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***However, in many cases the initiation of the first shadow violator cannot be explained with simple Coulomb calculations, meaning that:***

***(1) Optimal fault orientation calculations are suspect, and that detailed geological fault assignments are needed***

***and/or***

***(2) True mainshock rupture complexity is not accounted for***

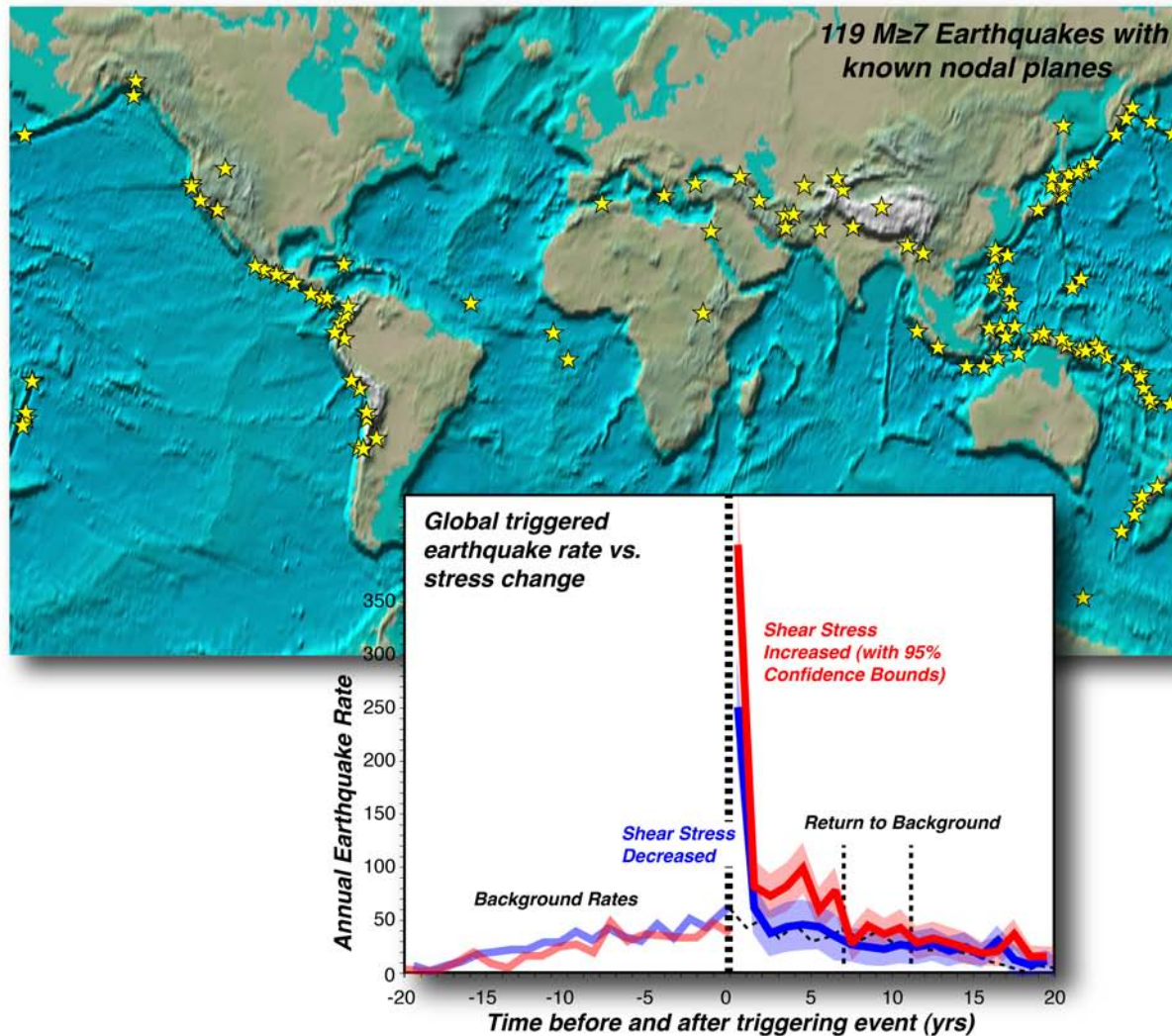
***and/or***

***(3) Near source dynamic triggering happens***



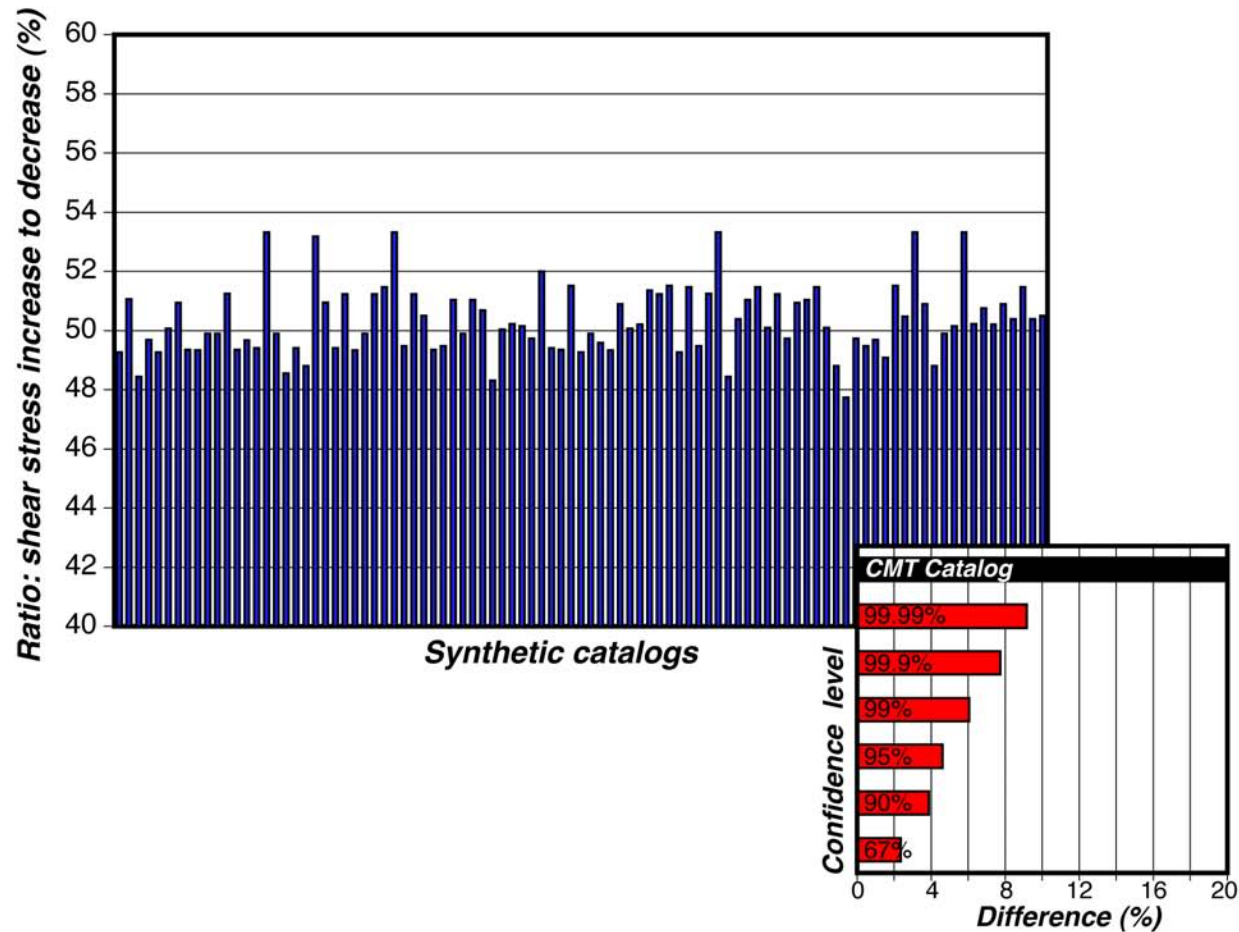
## ***Revisiting global database of stress changes from centroid moment tensor planes of 119 $M \geq 7$ EQ's onto $>1300$ $M \geq 4.5$ planes:***

*Raw result: 61% of events within  $\sim 10$ -250 km range associated with increased shear stress over 20 yr*



## Revisiting global database of stress changes from centroid moment tensor planes of 119 $M \geq 7$ EQ's onto $>1300$ $M \geq 4.5$ planes:

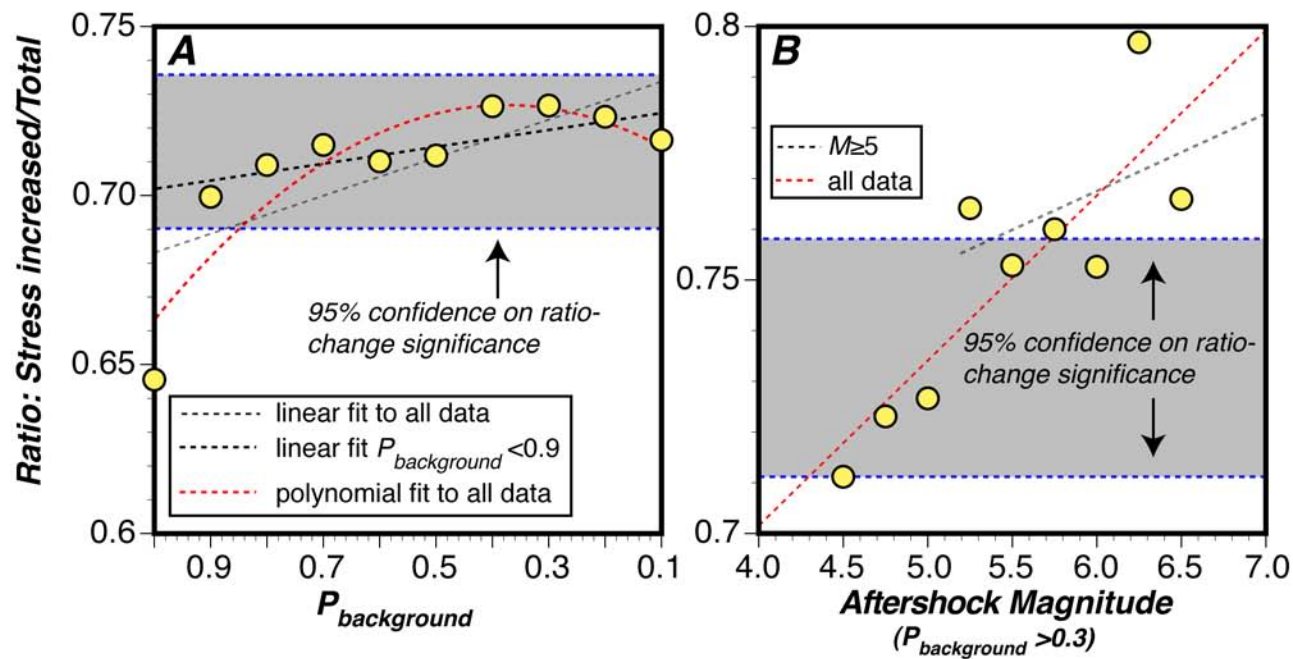
Shear stress changes calculated using 100 synthetic catalogs show expected variability





## Revisiting global database of stress changes from centroid moment tensor planes of 119 $M \geq 7$ EQ's onto $>1300$ $M \geq 4.5$ planes:

- Stress-increased ratio grows with decreasing background probability
- Stress-increased ratio grows with increasing magnitude

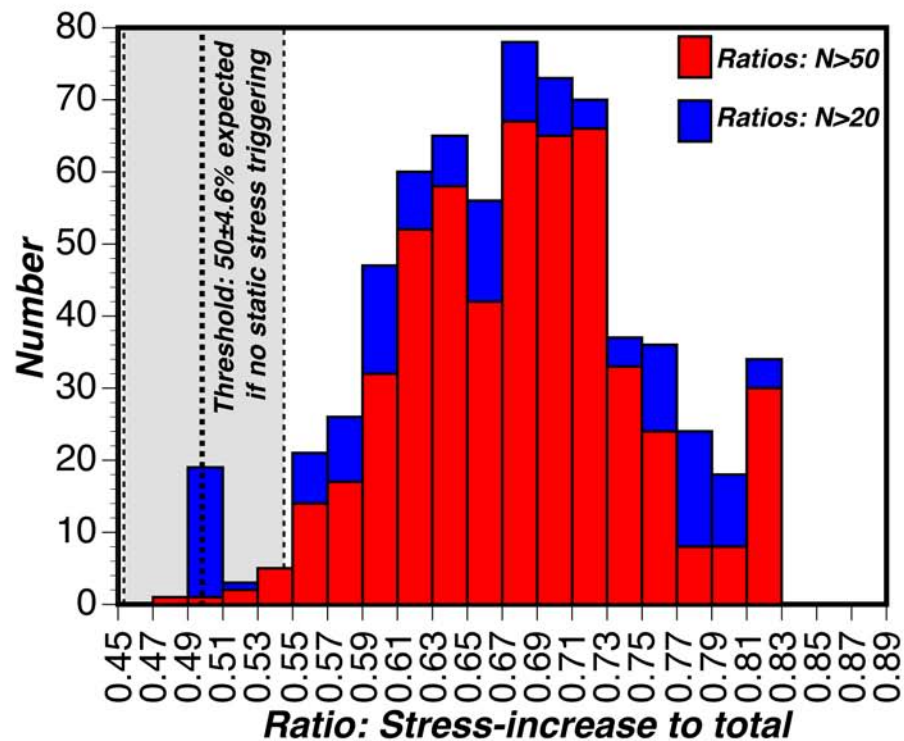


## Revisiting global database of stress changes from centroid moment tensor planes of 119 $M \geq 7$ EQ's onto >1300 $M \geq 4.5$ planes:

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Distribution of shear-stress-increase ratios as functions of:

- Aftershock magnitude
- Mainshock magnitude
- Background probability





***Examining the the global catalog suggests that the largest aftershocks are most consistent with static stress calculations***

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***We think this means that when the aftershock planes are well resolved, then the static stress change calculations are decent forecasters (consistent ~75-80% of the time above  $M=6$ ).***

***Margarita Segou tests this idea with a retrospective forecast test in Northern California that compares a geologically-based Rate/State and an empirical ETAS forecast.***

# FORECASTING EARTHQUAKES IN NORTHERN CALIFORNIA USING PHYSICS-BASED AND STATISTICAL MODELS

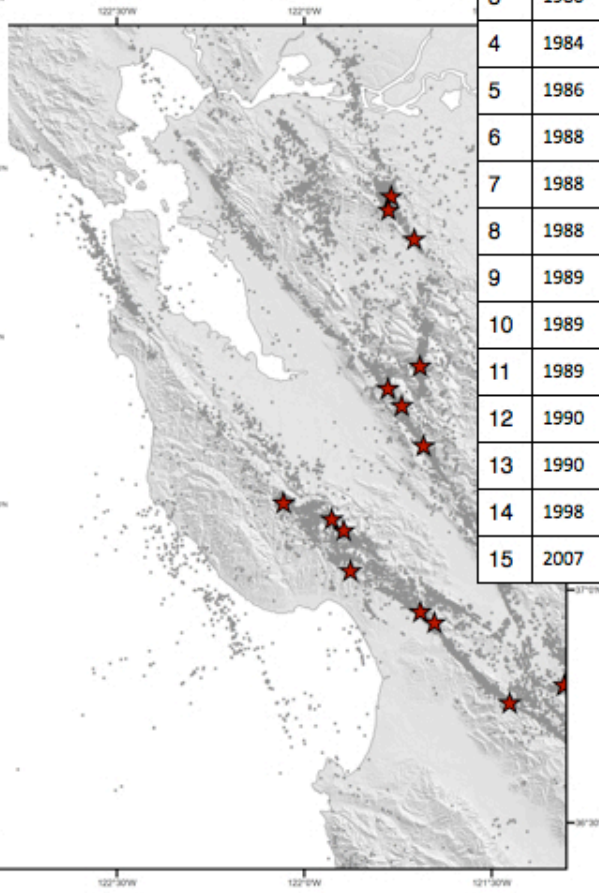
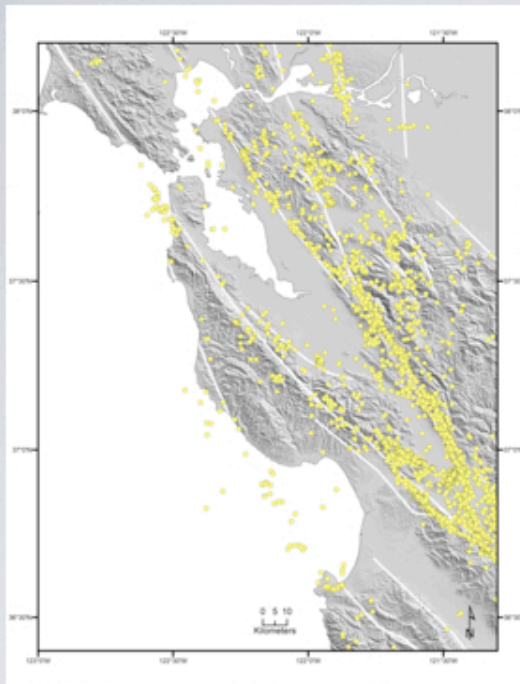
Margarita Segou

[msegou@usgs.gov](mailto:msegou@usgs.gov), [msegou@gmail.com](mailto:msegou@gmail.com)

Stanford Geophysics 5/22/12  
USGS 5/23/12

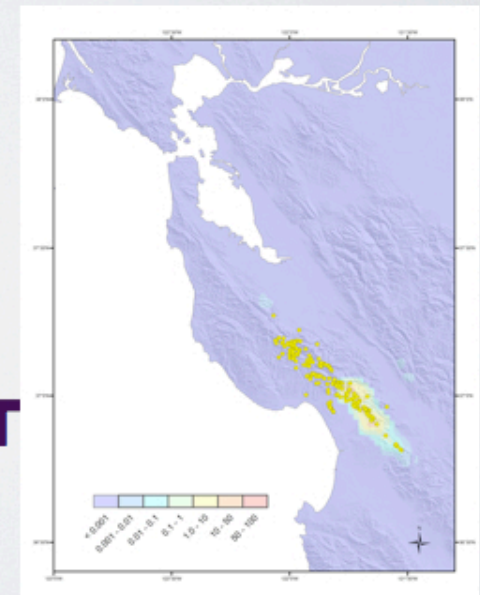


# CRS components



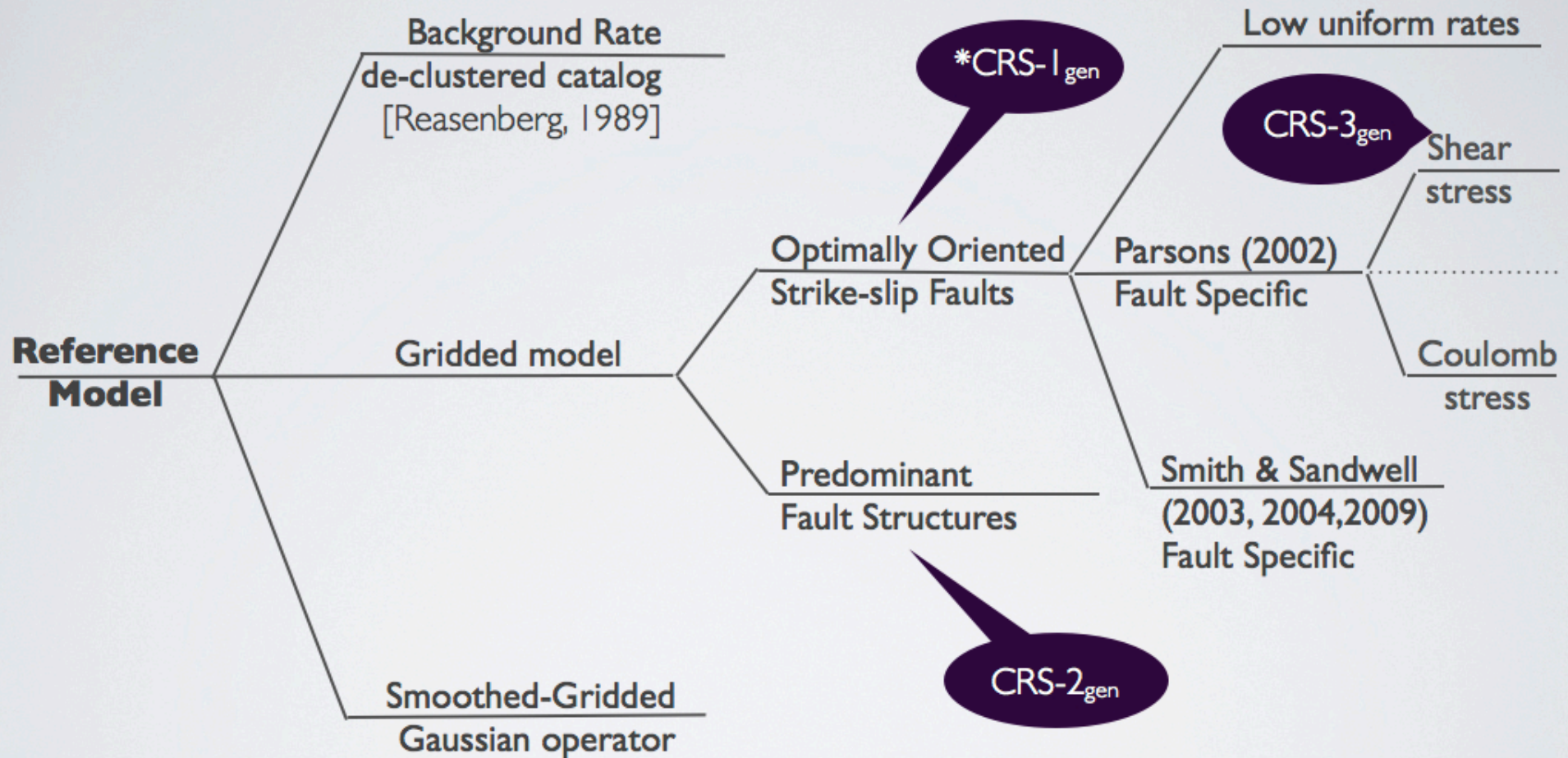
No	Y	mo	D	h	mi n	M	d	EQ name
1	1980	1	24	19	0	5.8	14.79	Livermore I
2	1980	1	24	19	1	5.1	6.9	Livermore
3	1980	1	27	2	33	5.4	14.43	Livermore II
4	1984	4	24	21	15	6.2	7.968	Morgan Hill
5	1986	3	31	11	55	5.7	8.394	Mt. Lewis
6	1988	2	20	8	39	5.1	8.222	Alum Rock 1988
7	1988	6	13	1	45	5.3	8.869	Hollister
8	1988	6	27	18	43	5.3	11.54	Ellsman Lake I
9	1989	8	8	8	13	5.4	12.59	Ellsman Lake II
10	1989	10	18	0	4	7	16.41	Loma Prieta mainshock
11	1989	10	18	0	41	5.1	13.86	Loma Prieta aftershock-St Cruz Mts
12	1990	4	18	13	53	5.4	4.611	Watsonville I
13	1990	4	18	15	46	5.1	6.338	Watsonville II
14	1998	8	12	14	10	5.1	7.746	SJ Bautista
15	2007	10	31	3	4	5.4	7.486	Alum Rock 1988

$$\begin{array}{l}
 \text{Reference} \\
 \text{Rate} \\
 \text{ANSS Mc3.0} \\
 \text{1974-1980}
 \end{array}
 +
 \begin{array}{l}
 \text{DCFF} \\
 \text{M>5.0} \\
 \text{1980-2009}
 \end{array}
 =
 \begin{array}{l}
 \text{RETRO} \\
 \text{FORECAST} \\
 \text{Mc3.0}
 \end{array}$$



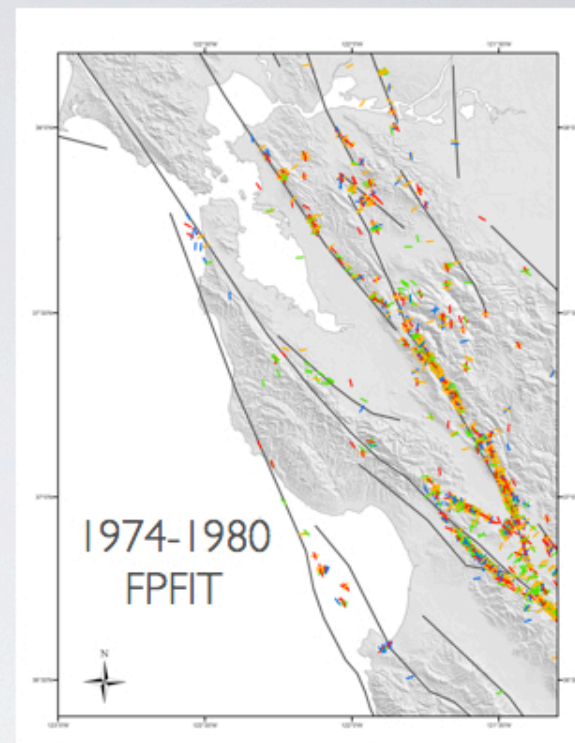
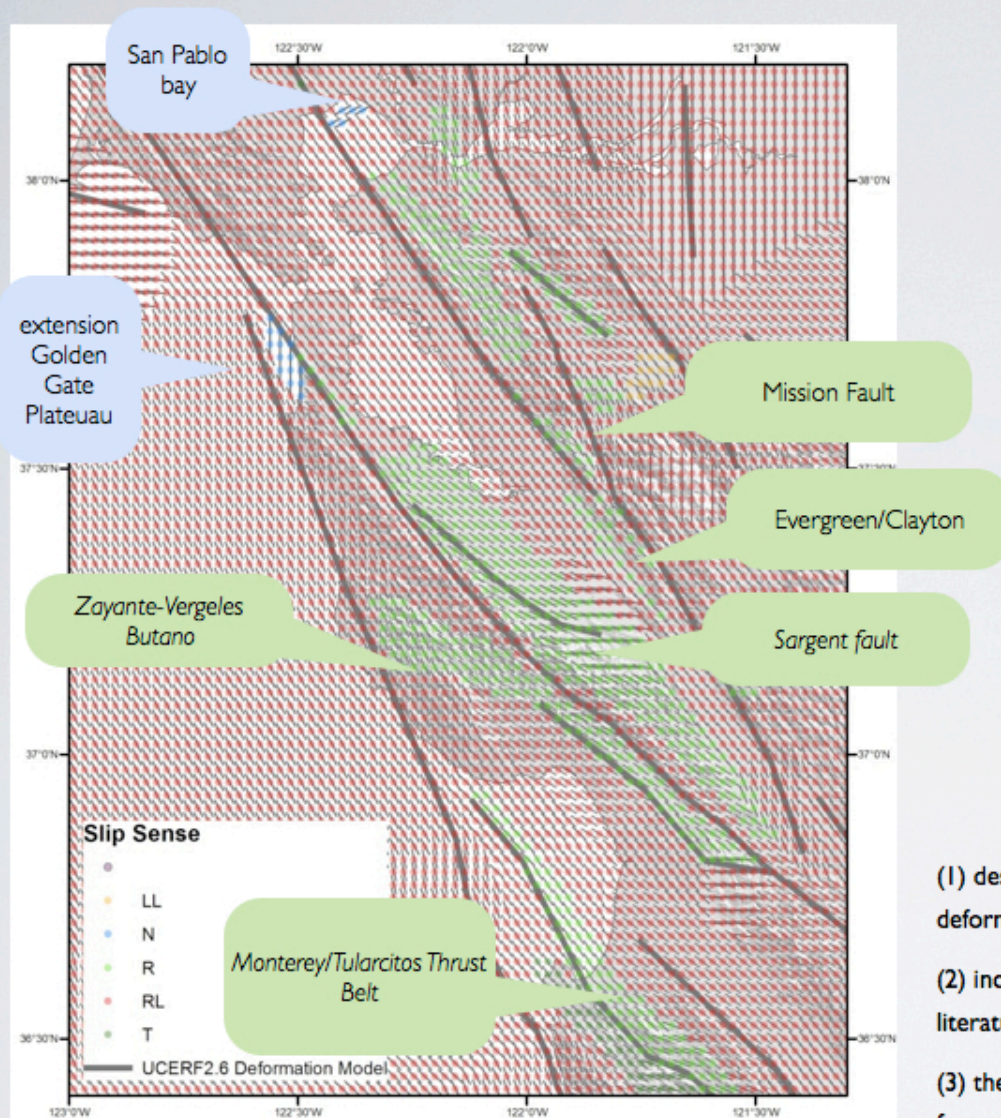


# Diagram for calculation sets for each CRS forecast model



\* increases the number of free parameters & refers to optimally oriented fault planes to the pre-mainshock stress field





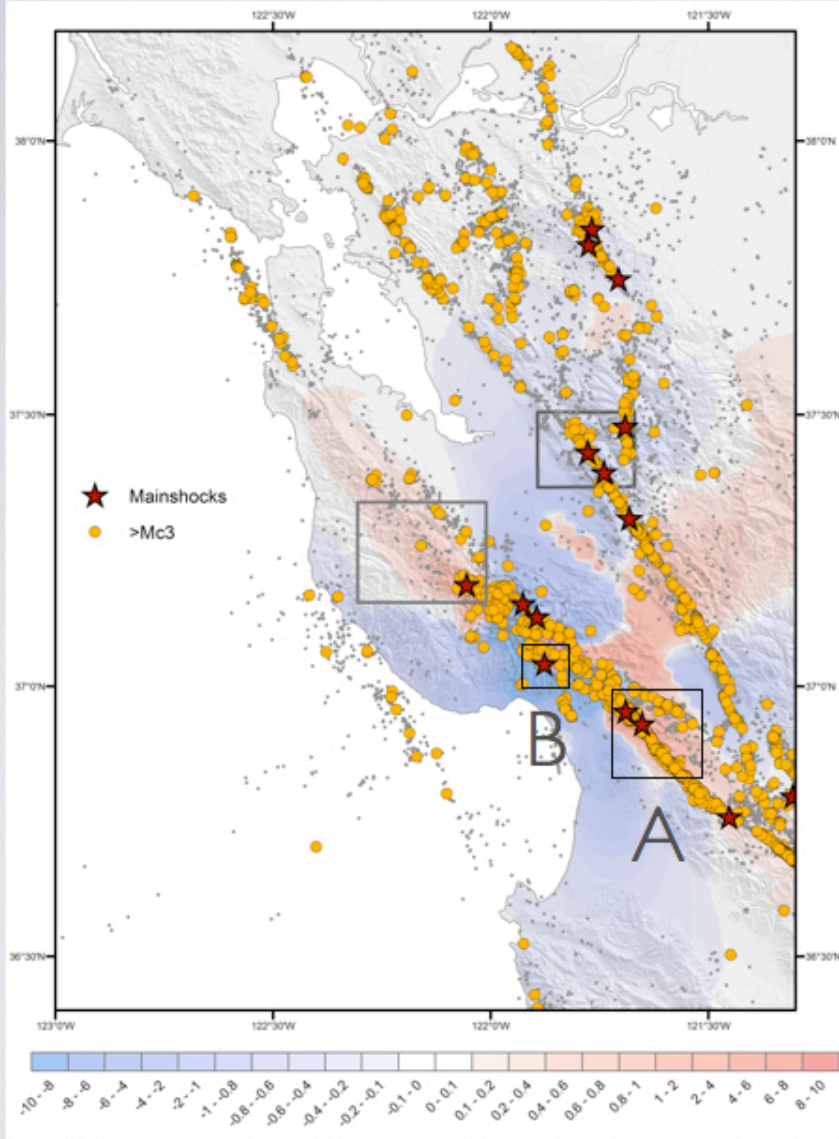
## Requirements

- (1) describe main geological features which are involved in the long term deformation history of Northern California and
- (2) include, whichever the scale, important structures for which we now from literature survey that there are spatio-temporal implications for their activity
- (3) the use of seismological data should be constrained to data available inside the forecast's learning phase (1974-1980) to avoid conflicts arising from any prior knowledge, which would jeopardize our evaluation

# 3D-GRID DISCRETE FAULT PLANES

PREDOMINANT GEOLOGY





today we are focusing  
on areas A & B

### *Area A advantages*

- study models' performance at the off-fault zones
- time-dependency of performance since Watsonville events occur some months after Loma Prieta

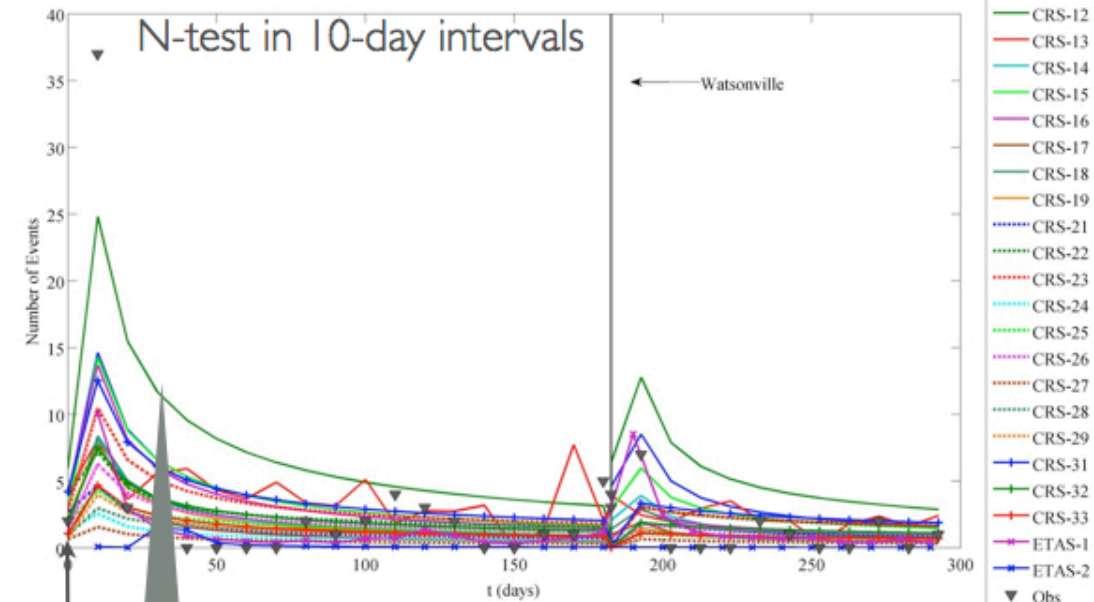
### *Area B advantages*

- which is the best choice for the near-source region?  
CRS or ETAS?

Different Testing Areas? What we hope to find out?



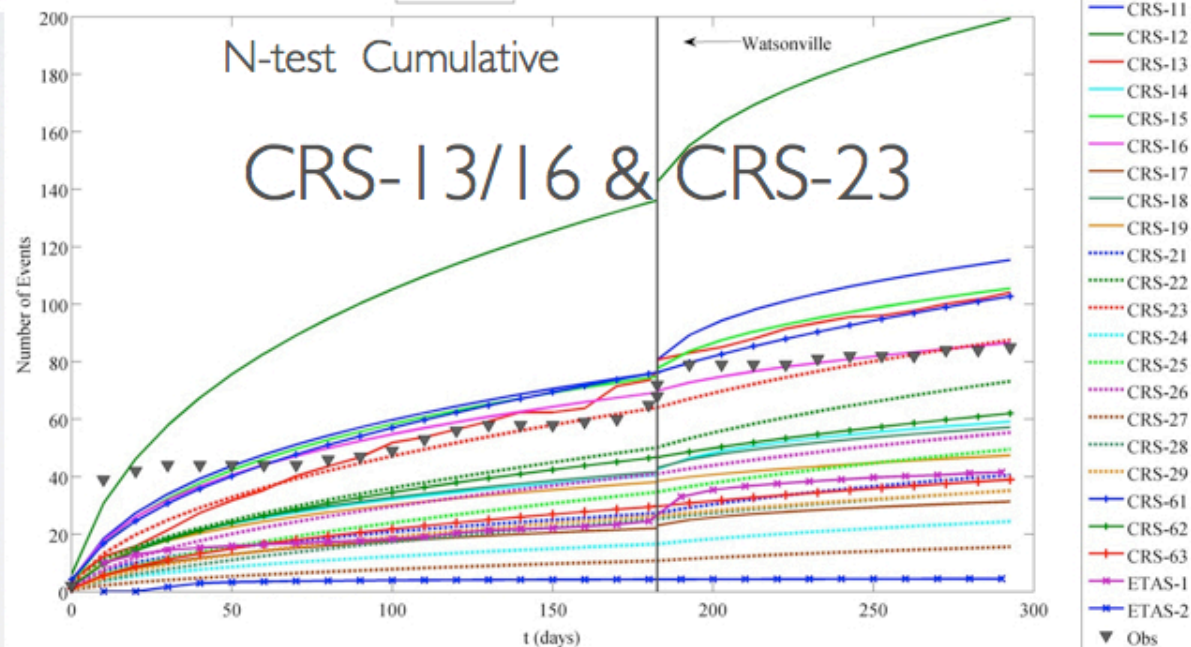
@off-fault



given sufficient time even the most slowly stressed 'faults' will return to reference state

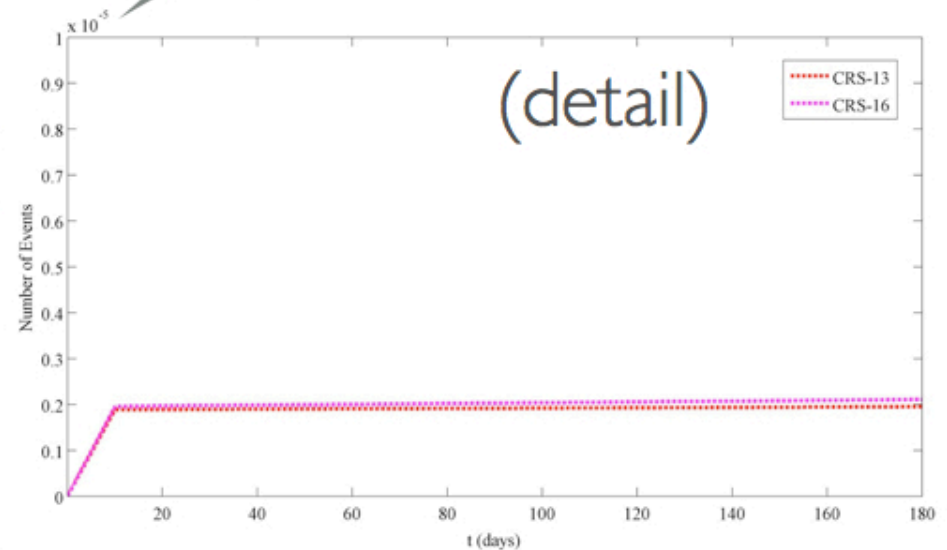
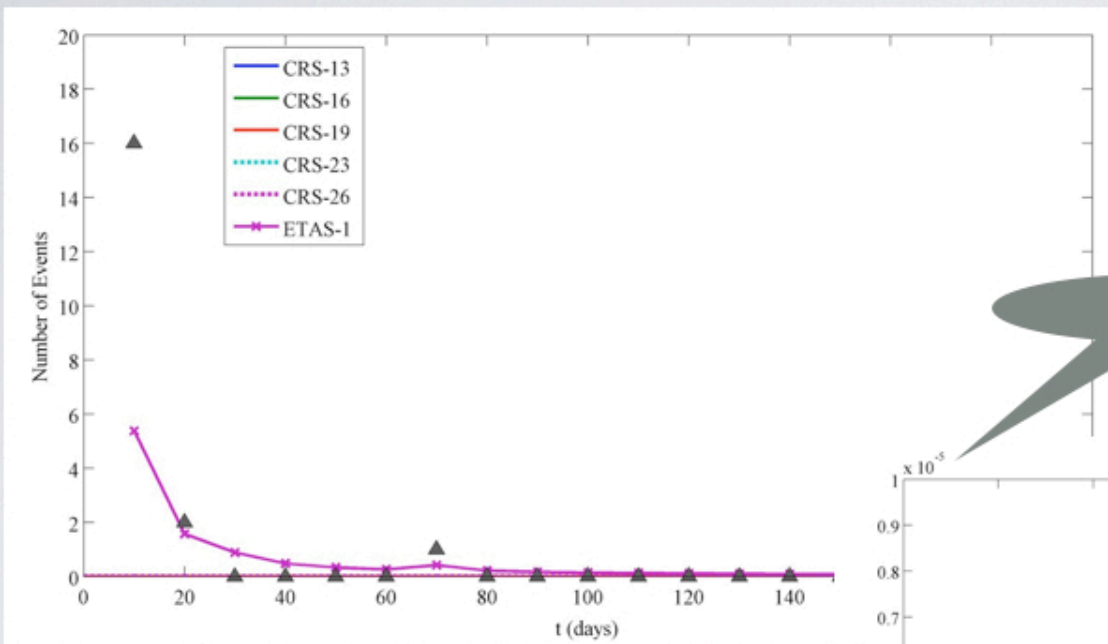
ETAS under-predicts in the beginning since a small number of ancestor events  
CRS under-predicts due to low background rates

first 40 min



NUMBER OF EVENTS

@near source



10 km from Loma Prieta mainshock inside DCFF<0  
crash test between the best performing forecast models



# PERFORMANCE EVALUATION RESULTS

Physics-based models @short-term outperform ETAS @off-fault region. Formulations based on smoothed-gridded Coulomb-stress based on optimally-oriented with loading rates on SAF [0.067-0.3 bar/yr] & receivers based on predominant geology [0.3 bar/yr] result in better spatial consistency

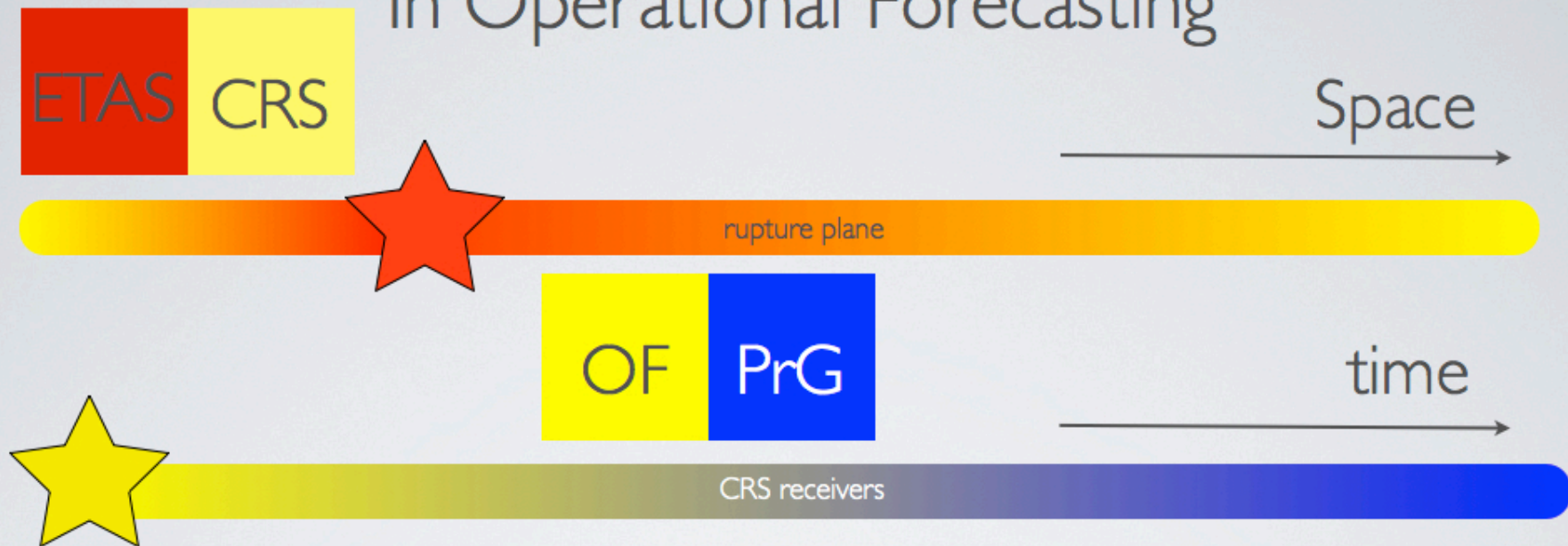
CRS forecast models are represented @long-term @off-fault region by models resolved on predominant geology receivers

What happens @near-source region ? ETAS models outperform CRS models due to low reference rates & stress shadow zones

Critical to incorporate previous important ruptures



# In Operational Forecasting



rely not on an optimum model but in a  
combination of best-performing models

we need pre-definition of best-performing CRS models & requires update of the state variable of the system

CRS models covers successfully the off-fault area where prediction is critical, no immediate need for variable slip distributions

introduction of stochasticity for simulations of early (1<sup>st</sup> day) aftershocks, then revert to observed aftershocks as ETAS ancestor events, depending on network's detectability